Downwind Development in a Stationary Band Complex Leading to the Secondary Eyewall Formation in the Simulated Typhoon Soudelor (2015)

Xue-Song Zhu, Hui Yu, and Yuqing Wang

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Introduction

- Intense TCs often develop a secondary eyewall.
- Mechanisms leaded to SEF:
 - VRWs: VRWs propagate radially outward to a stagnation radius. The wave kinetic energy is converted to the tangential wind by wave-mean flow interaction. Then, the wind enhance surface enthalpy flux and initiate convection. (Montgomery and Kallenbach 1997; Mong and Emanuel 2003)
 - Unbalanced dynamics: BL inflow induced by outer rainband can sharpen the radial wind gradient and create strong supergradient wind. Broadening of the tangential wind induces the unbalanced outflow just above the BL and then develops the axisymmetric secondary eyewall. (Qiu and Tan 2013; Huang et al. 2012; Abarca and Montgomery 2013, 2014)
 - Balanced dynamics: Strong diabatic heating in outer rainbands and high environmental RH can broaden the tangential wind outside the primary eyewall. Diabatic heating from stratiform precipitation is more effective in producing a secondary wind maximum than that form convective precipitation. (Wang 2009; Rozoff et al. 201; Moon and Nolan 2010)
 - Large scale interaction: Interaction between a TC and a midlatitude westerly jet can produce stratiform precipitation with embedded deep convection and evolve into the secondary eyewall finally. (Dai et al. 2017)

Introduction

- Asymmetric rainbands in SEF:
 - Downward and inward development of an intense outer rainband and associated inflow (MDI) led to the associated strong tangential wind above BL. (Wang et al. 2019)
 - When reaching the outer edge of the rapid filamentation zone, the tangential wind jet is effective axisymmetrized. The secondary tangential wind maximum and convective ring formed sequentially.
 - MDI initiated in the stratiform precipitation region can descend into the BL and trigger an intense updraft. The associated secondary circulation can accelerate the tangential wind. (Didlake et al. 2018; Yu et al. 2021; Wang and Tan 2020)

Introduction

- Not all outer rainbands propagate radially inward and lead to SEF.
- Stationary band complex (SBC): wavenumver-1 asymmetry and quasi-stationary relative to the TC.
- In a statistical analysis of 5-year microwave satellite observations:
 - 6 hours prior to SEF: 79% of the 84 SEF cases had SBC.
 - 12 hours prior to SEF: SBC becomes more tangential and axisymmetrization. (Vaughan et al. 2020)
- Objective:
 - The evolution from spiral rainband to a SBC
 - How the SBC contributes to the convective onset in SEF?

WRF Configuration

Version	3.8.1	
Domains	3, d02 and d03 are moving	
Grid size	18, 6, 2 (km)	
Start time	2015-08-01 1200 Z	
End time	2019-08-07 1200 Z	
Eta levels	36 (8 levels in z < 1.5 km)	
Model top	20 hPa	
Cumulus Scheme	Kain-Fritsch (d01 only)	
Microphysics Scheme	NSSL two-moment	
Longwave Scheme	RRTM	
Shortwave Scheme	RRTM	
PBL Scheme	YSU	
Land surface scheme	Noah	
DATASET resolution	0.25° x 0.25°	
DATASET	NCEP GFS FNL	

All horizontal wind in this study are storm-relative.



Simulation results



Plan view of simulated Soudelor



Z = 3 kmColored: reflectivity (dBZ) Arrow: VWS (m/s) Circle: 30, 90, 150 km from TC center

60

55

50

45

40

35

30

25

20

dBZ

Vertical wind shear (VWS): Difference between 200 and 850 hPa of averaged wind speed in r = 200-800 km annular area.

Plan view of simulated Soudelor



Colored: Brightness temperature (K) Arrow: VWS (m/s)

190

200

210

220

230

240

250

260

270

140E

Vertical wind shear (VWS): Difference between 200 and 850 hPa of averaged wind speed in r = 200-800 km annular area.

Azimuthal mean

Azimuthal-mean Colored: w (m/s) Contour: v (m/s) Arrow: u, w (m/s)



Azimuthal mean

Azimuthal-mean Colored: w (m/s) at z = 3 – 7 km Contour: v (m/s) at z = 0.6 km

Middle-level updraft > 0.5 m/s outside the eyewall implies the strengthening of rainband convection. (Rozoff et al. 2012)









Cross section of SBC

Line-averaged Color: dBZ Vector: u, w Contour: w Contour: v'



Cross section of SBC

Line-averaged Color: θ_e Vector: u, w Contour: qr Contour: qv Dashed Contour: RH



Supergradient force

Line-averaged Color: divergence Vector: u, w Contour: w Contour: AF

$$\mathbf{AF} = f\boldsymbol{v} + \frac{\boldsymbol{v}^2}{r} - \frac{1}{\rho}\frac{\partial p}{\partial r},$$

Agradient force (AF) AF > 0 supergradient wind AF < 0 subgradient wind



Propagation speed of IR



Average radial propagation speed of IR: 2.4 m/s 65 km -> 73 km -> 82 km Average tangential propagation speed of IR: 17.3 m/s (43% of tangential wind) Low-level outflow: 1.6 m/s

Tangential propagation speed associated with VRWs: 60 ~ 80% of tangential wind

Budget analysis of AF

$$AF = fv + \frac{v^2}{r} - \frac{1}{\rho} \frac{\partial p}{\partial r},$$
$$AF_O \qquad AF_I$$

Contour: w at z = 1 km Red dot: max dBZ



Budget analysis of AF





Cold pool formation in the downwind propagation

Average between z = 1 - 3 km Color: 5x hourly-change θ_e Circles: 30, 90, 150 km





Average between $\theta = 210 - 250^{\circ}$ (cross rainband) Color: θ_e

Cold pool formation in the downwind propagation



Color: Asy - PGF Contour: u Contour: inward PGF Vector: u, w

Cold pool dynamics in convection enhancement

Line-averaged Color: θ difference Contour: w

Vector: cross-band, vertical wind

 $\begin{aligned} \Delta\theta &= \theta_{min} - \theta_0 \\ \theta_0: \theta \text{ before cold pool reached} \\ (average within 25-km radial \\ distance in the outward side of \\ rainband) \end{aligned}$





Cold pool formation in the downwind propagation

Cold pool radial speed:

$$V_{\rm cp} = k \sqrt{\frac{gH(\rho_c - \rho_e)}{\rho_e}} + bV_c$$

 ρ_c : cold pool density ρ_e : environmental density k: 0.9 b: 0.6



	SR_54	SBC_D	
<i>H</i> (km)	4.0	5.5	(4)
$ ho_c$ (kg/m³)	1.132	1.128	Timo
$ ho_e$ (kg/m ³)	1.129	1.123	¹
<i>V_c</i> (m/s)	2.88	-13.26	
mean V_{cp} (m/s)	-5.19		



(c) Organization of SBC with convection enhanced in the downwind sector

Stationary band complex (SBC)



(d) A dynamical balance for convection enhancement

