

Paper review

mainly refer to

On the Correlation between Total Condensate and Moist Heating in Tropical Cyclones and Applications for Diagnosing Intensity

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Outline

☀ Introduction

☀ Simulated tropical cyclones

☁ *Simulations modeled after the hurricane nature run*

☀ To Use numerical simulations of TCs:

☁ Total condensate or rain rate

☀ To → Diabatic heating

☁ → Intensity (change)

☁ → Evaluate narrow-swath measurements

☁ *Some results with different microphysics schemes*

☁ *TC intensity and rain rate*

☀ Results with narrow-swath measurements

☁ *Motivation and framework*

☁ *Results*

☁ *Analogous estimates with real CloudSat profiles*

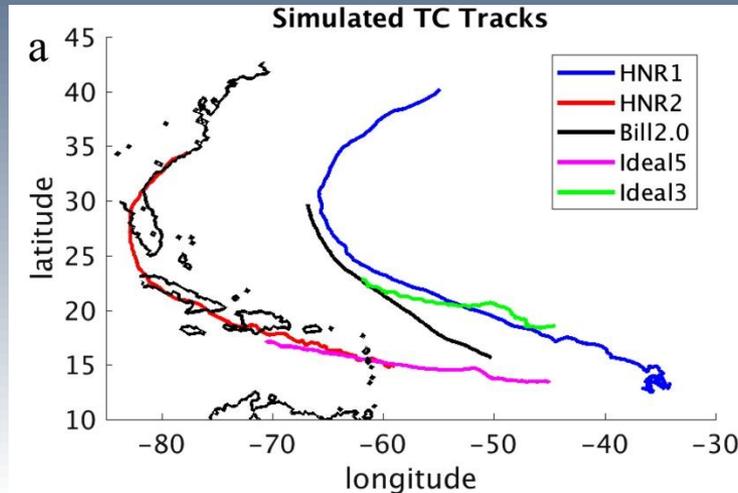
☀ Conclusions

Introduction

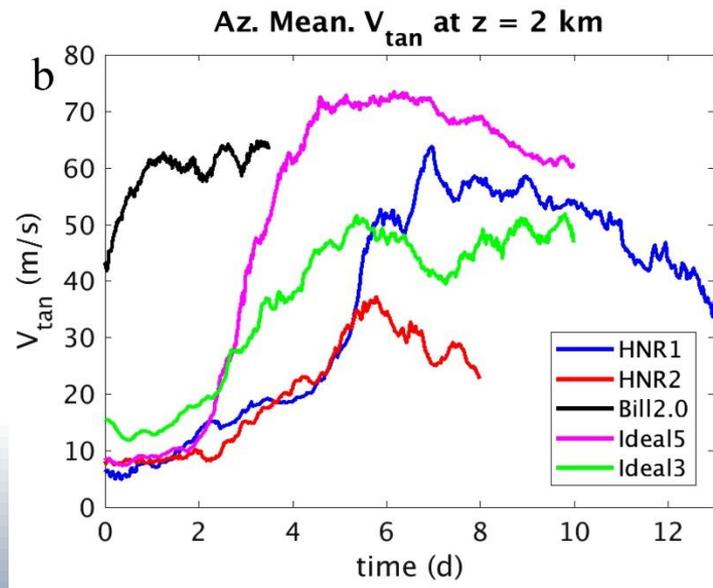
- ☀ Close relationships between the **production of condensate**, the **release of latent heat**, and **subsequent precipitation** are the physical basis for a long history of efforts to diagnose **TC intensity and intensity change** from satellites.
- ☀ Many previous studies have found that **reduced brightness temperatures** observed by microwave imagers from precipitating clouds can be **correlated with precipitation rates and thus with TC intensity** (Adler and Rodgers 1977; Rodgers et al. 1998, 2000; Cecil and Zipser 1999)
- ☀ The **vertical distribution of heat release near the TC center** could provide additional diagnostic and predictive power for remote sensing of TC intensity (Hack and Schubert 1986; Nolan et al. 2007; Vigh and Schubert 2009).
- ☀ **“Hot towers”** or **“convective bursts”** either precede or occur concurrently with periods of **significant intensification** (Kelley et al. 2004; Kelley and Halverson 2011; Guimond et al. 2010, 2016; Rogers et al. 2015, 2016; Hazelton et al. 2017; Wadler et al. 2018).
- ☀ Broader, more symmetric, and **more persistent convection** around the storm center is **better correlated with intensification** (Kaplan et al. 2010; Zagrodnik and Jiang 2014 and Shimada et al. 2017).

Simulated tropical cyclones

3 real cases and 2 ideal cases

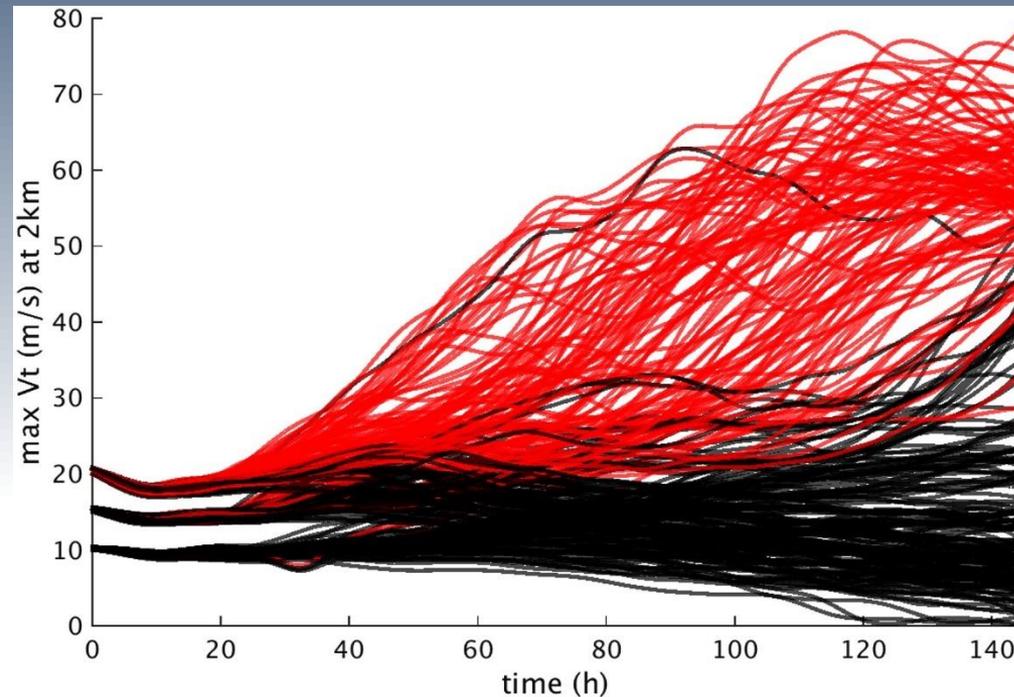


- WRF 3.2.1 and 3.4.1
- Nested grids with 27, 9, 3, and 1 km grid spacing
- WDM6 microphysics scheme
- YSU planetary boundary layer scheme
- RRTMG longwave and shortwave radiation schemes



Simulated tropical cyclones

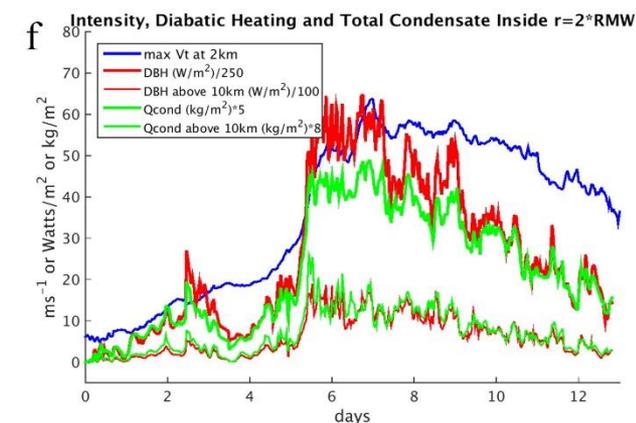
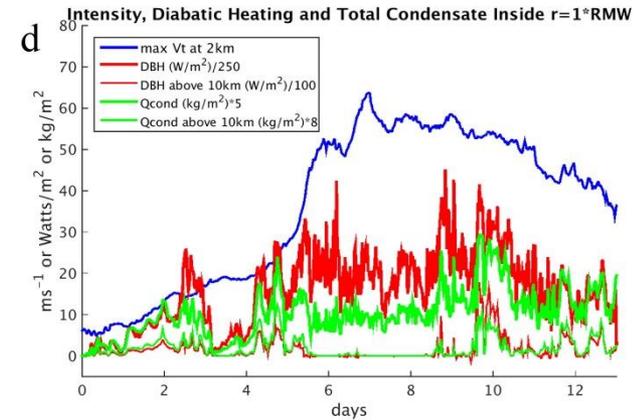
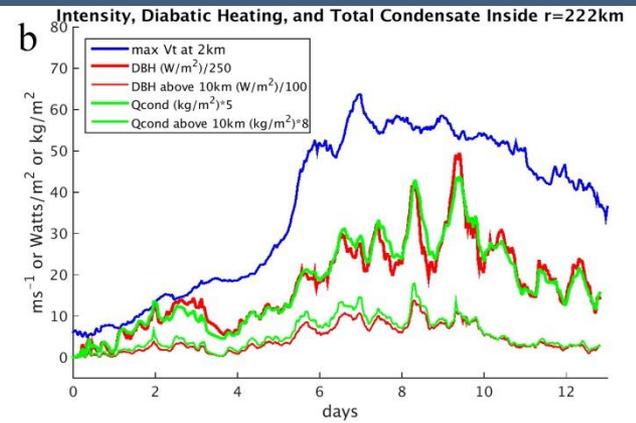
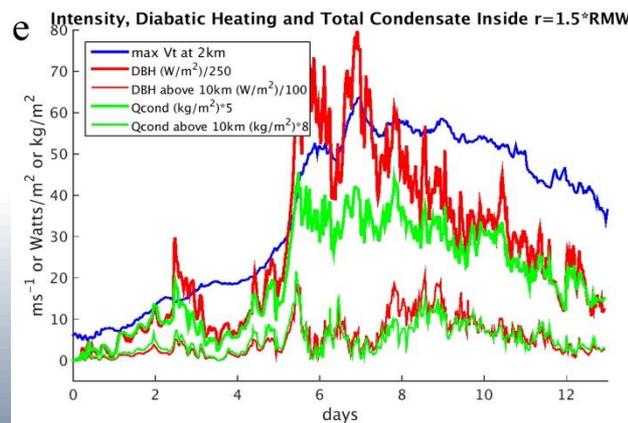
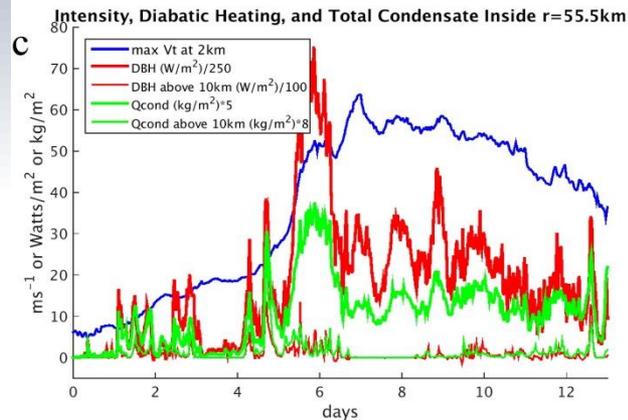
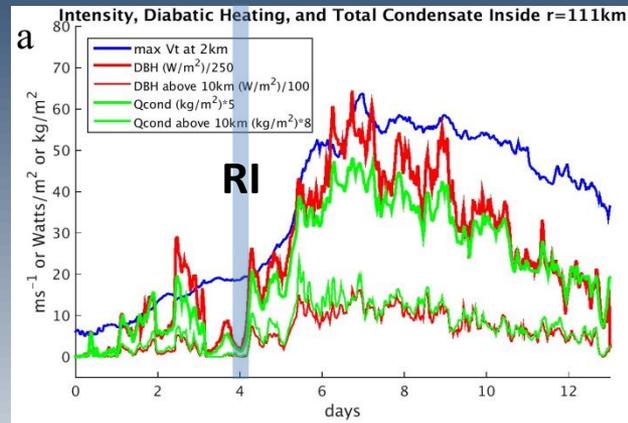
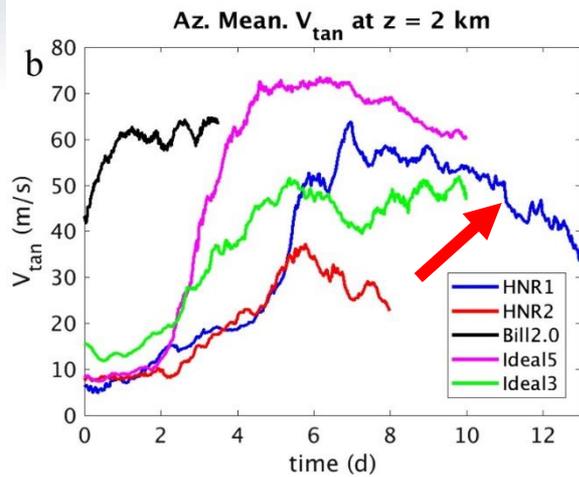
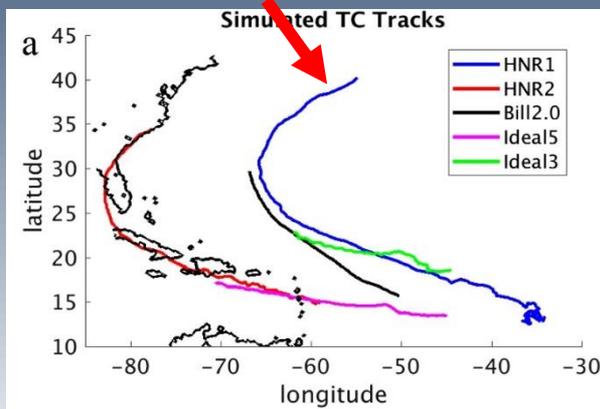
The Miyamoto and Nolan ensemble



- WRF 3.7.1
- Nested grids with 18, 6, 2 km grid spacing
- 40 vertical levels up to 20 km altitude
- WSM6 microphysics scheme
- YSU planetary boundary layer scheme

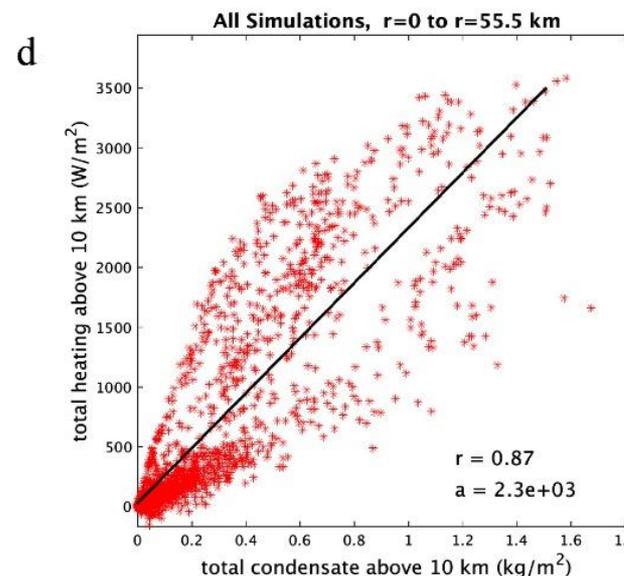
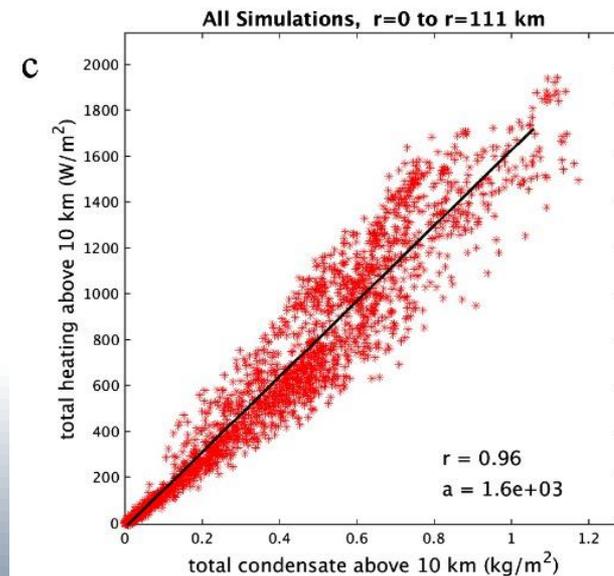
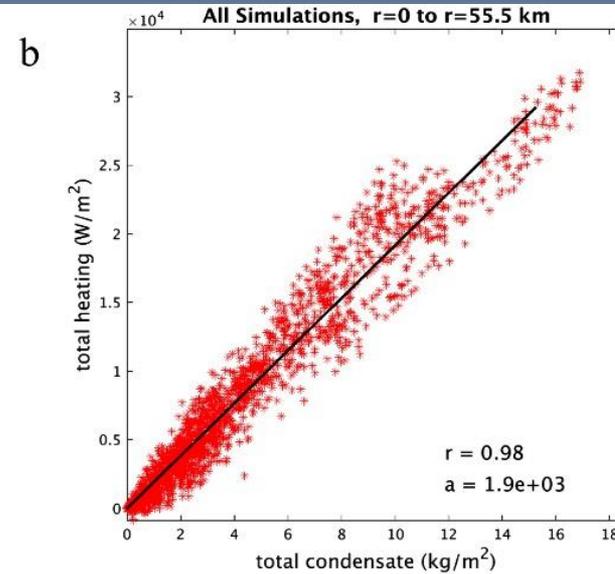
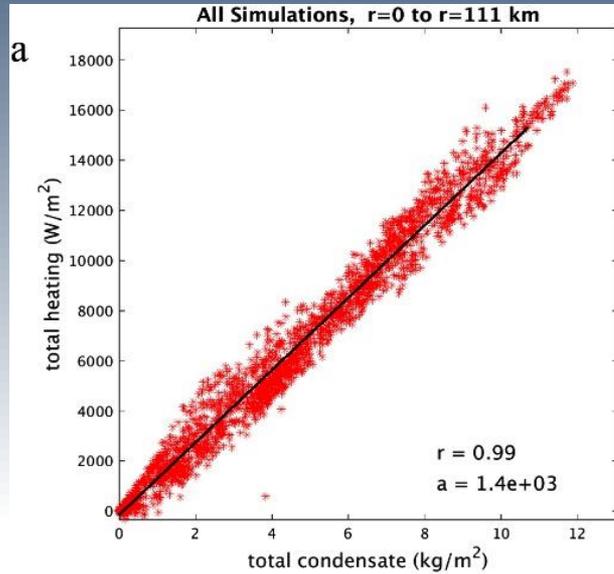
- surface wind speed (5)
 - wind shear (6)
 - Initial vortex size (3)
 - initial vortex strength (3)
- Total: 270

Integrated heating and condensate



Total condensate and total heating

ALL

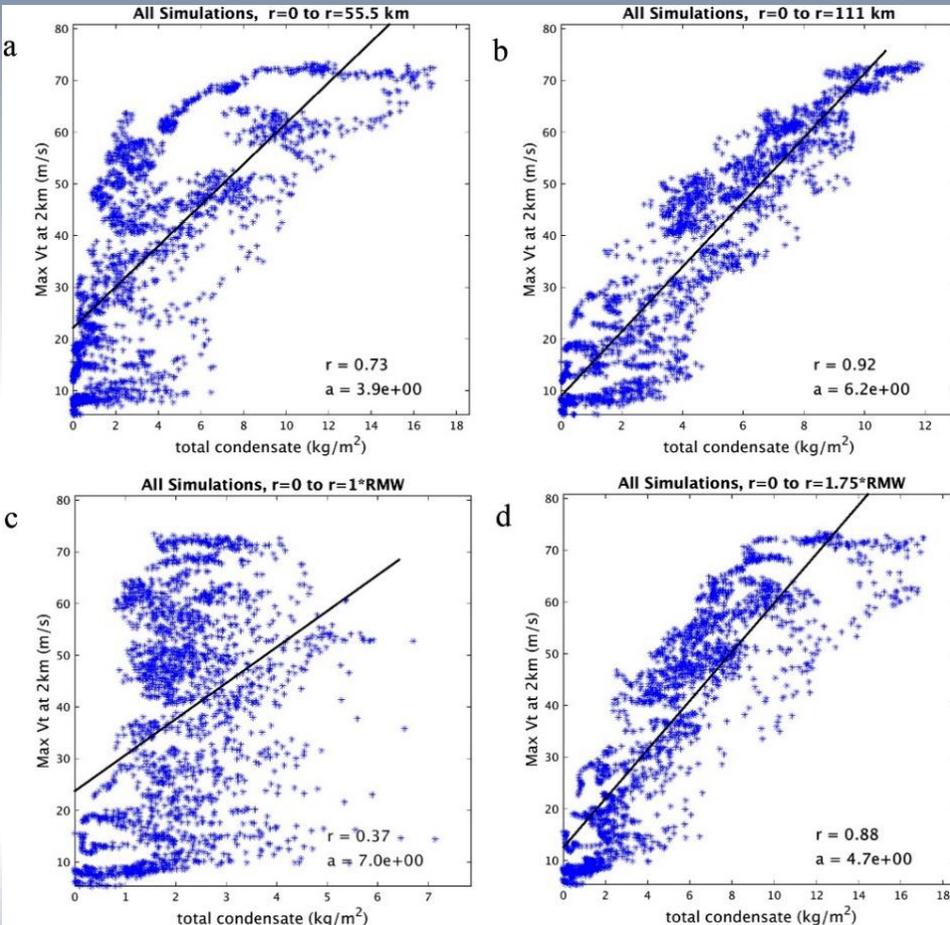


Above 10km

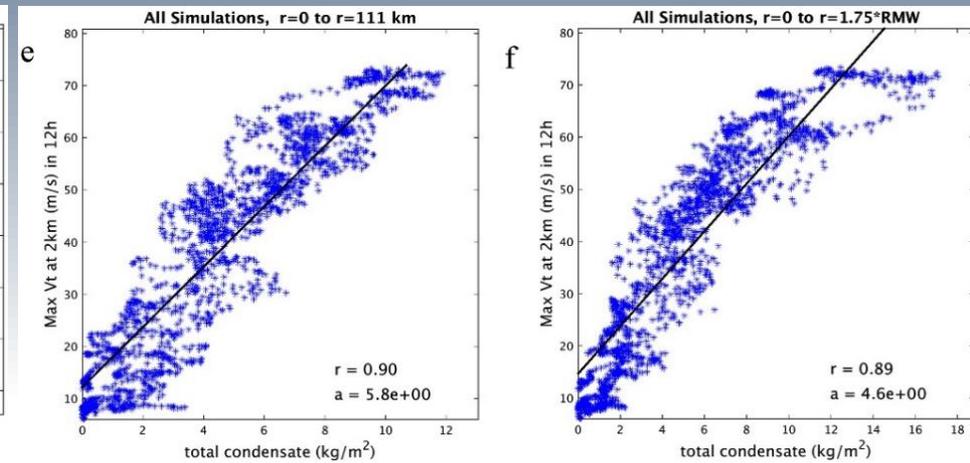
Total condensate and intensity

Intensity

Intensity (current)



Intensity (12 h later)

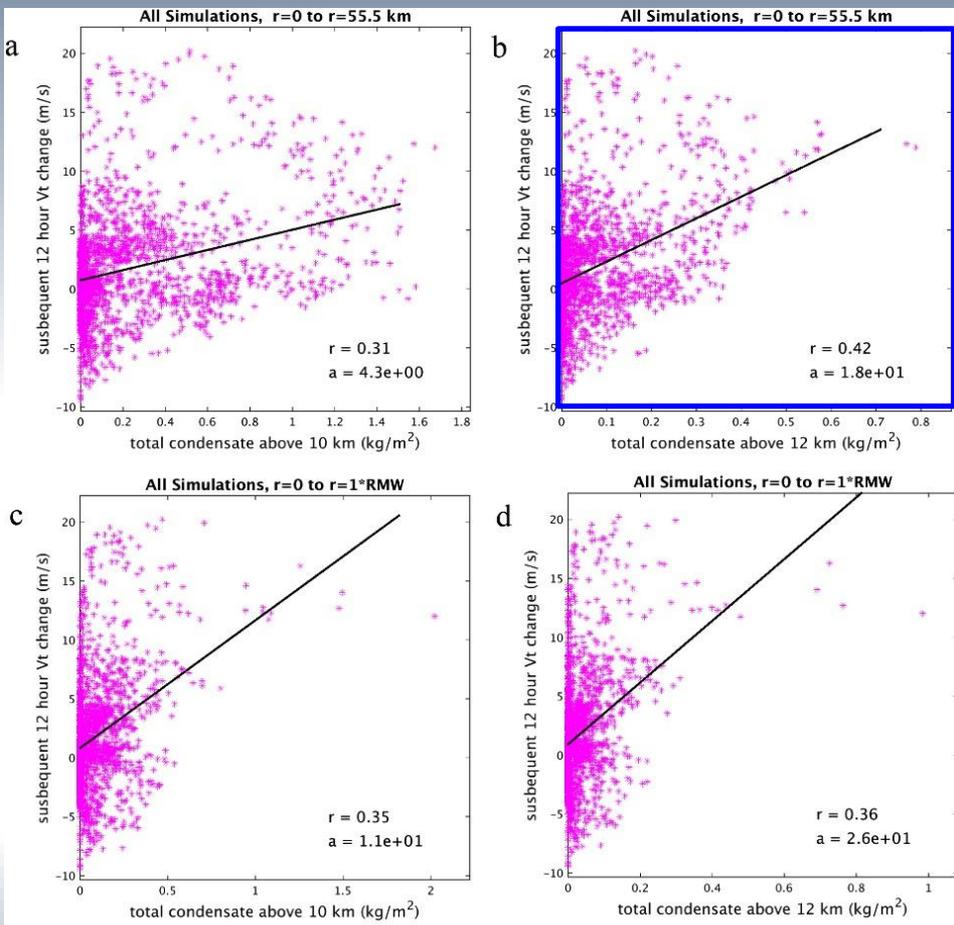


One might expect **considerable scatter** about this relationship, it does not account for differences in **storm size**, or factors that contribute to a **reduced efficiency** in the conversion of heat to wind energy.

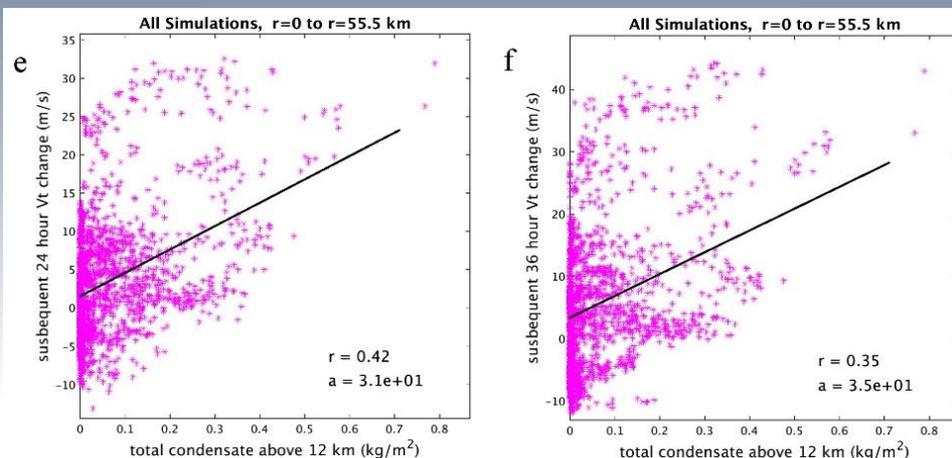
Total condensate and intensity

Intensity change

Intensity change (12 hr)

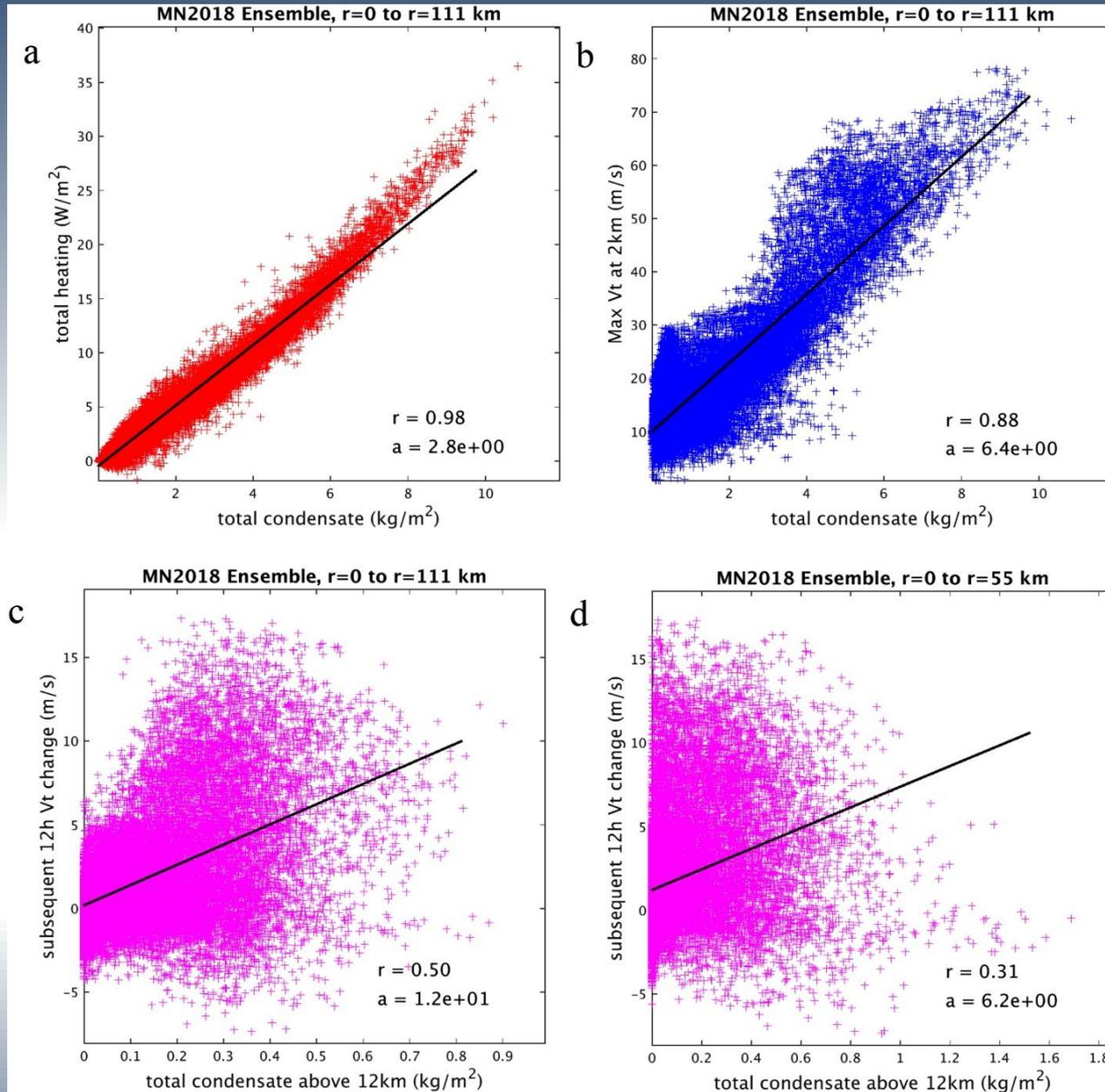


Intensity change (24 h later)



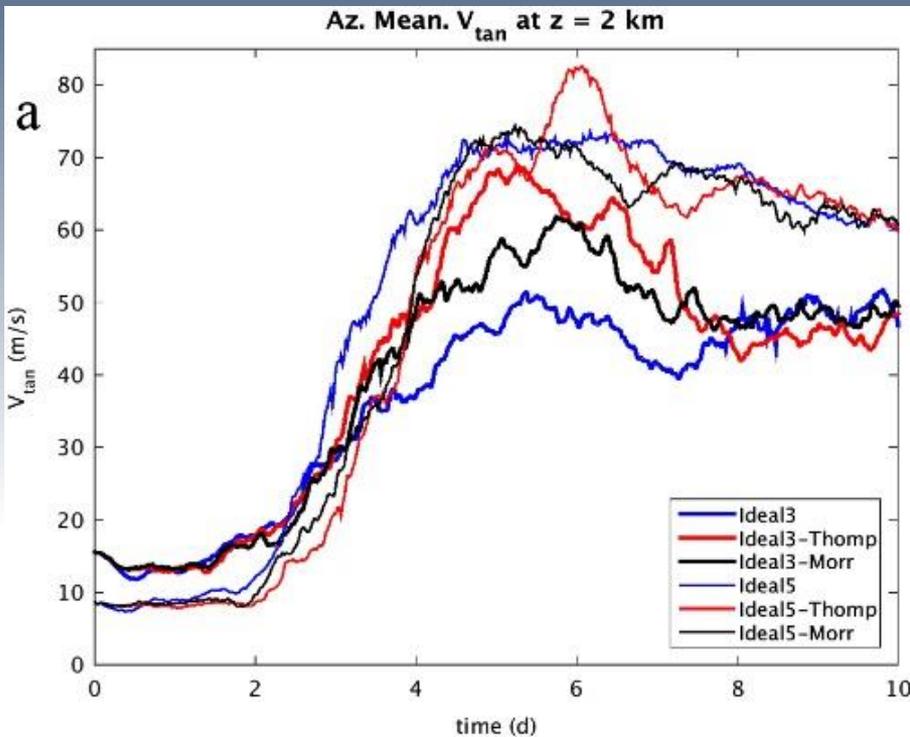
The strongest correlation is found for condensate above 12 km within a radius of 55 km, with $r = 0.42$

Results for the Miyamoto and Nolan ensemble

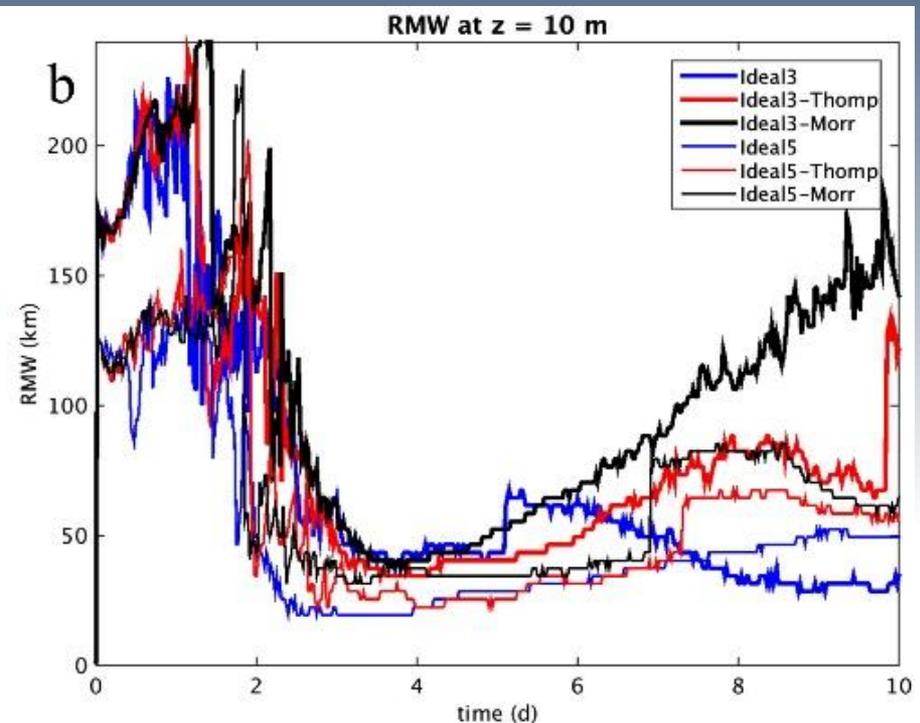


Different microphysics schemes

WDM6, Thompson and Morrison



Both schemes make **Ideal3** and **Ideal5** considerably stronger after their initial intensification phases, but then bring them back closer to the WDM6 results for the last few days of each simulation.



Both Ideal3 and Ideal5, the Thompson and Morrison schemes lead to steadily increasing TC sizes, with the RMW of the **Ideal3-Morrison case reaching 150 km** by $t = 10$ day, and **Ideal3-Thompson** showing an eyewall replacement cycle just before day 10, leading to **nearly equal size**.

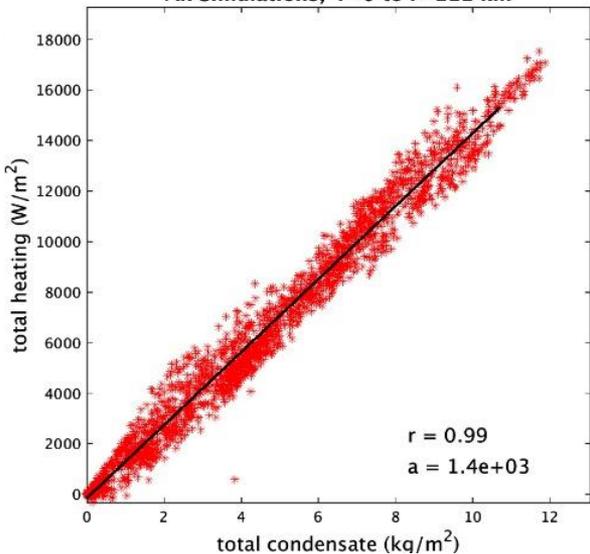
Different microphysics schemes

WDM6

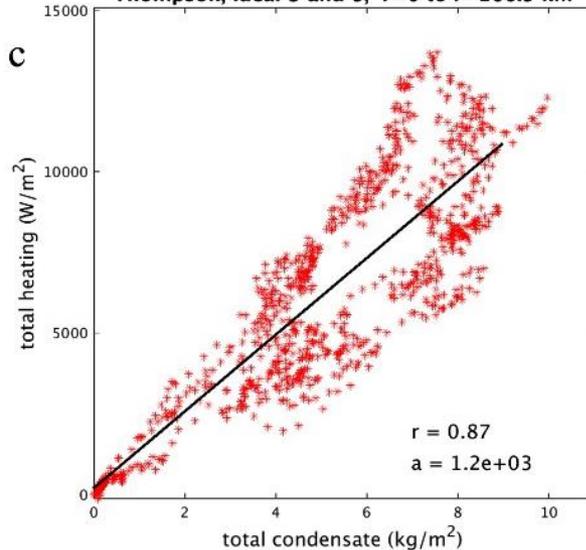
Thompson

Morison

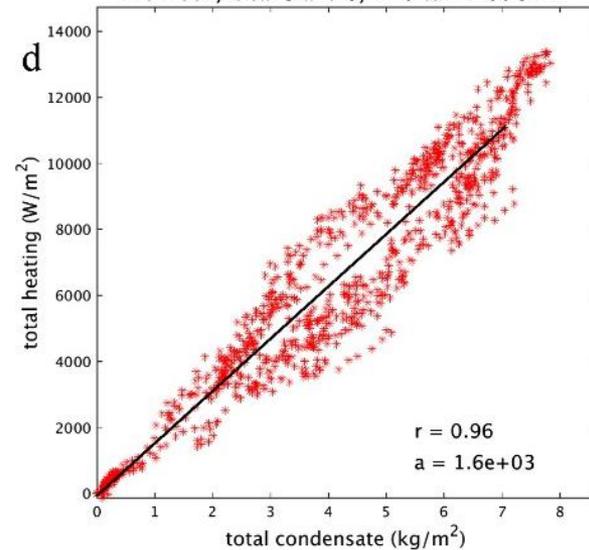
All Simulations, $r=0$ to $r=111$ km



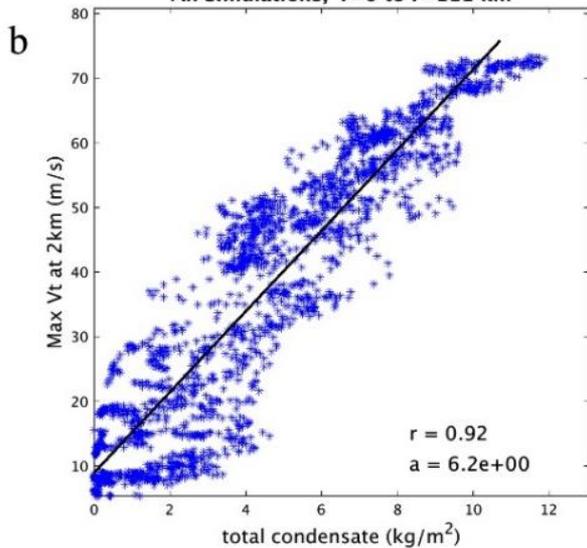
Thompson, Ideal 3 and 5, $r=0$ to $r=166.5$ km



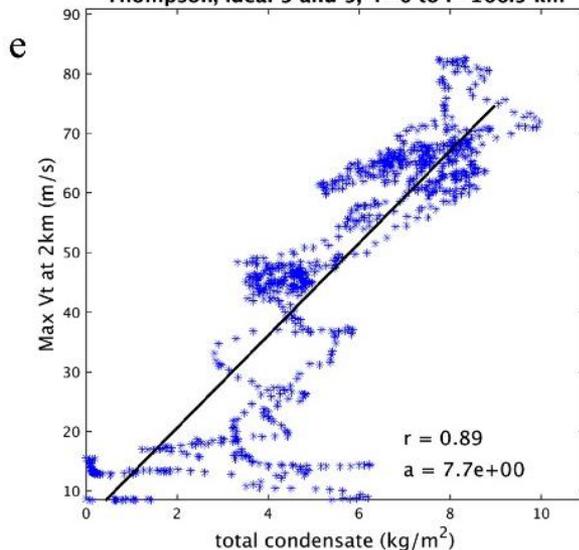
Morrison, Ideal 3 and 5, $r=0$ to $r=166.5$ km



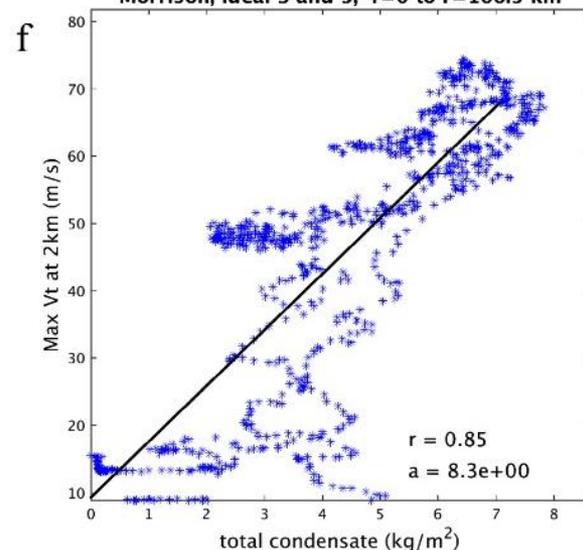
All Simulations, $r=0$ to $r=111$ km



Thompson, Ideal 3 and 5, $r=0$ to $r=166.5$ km



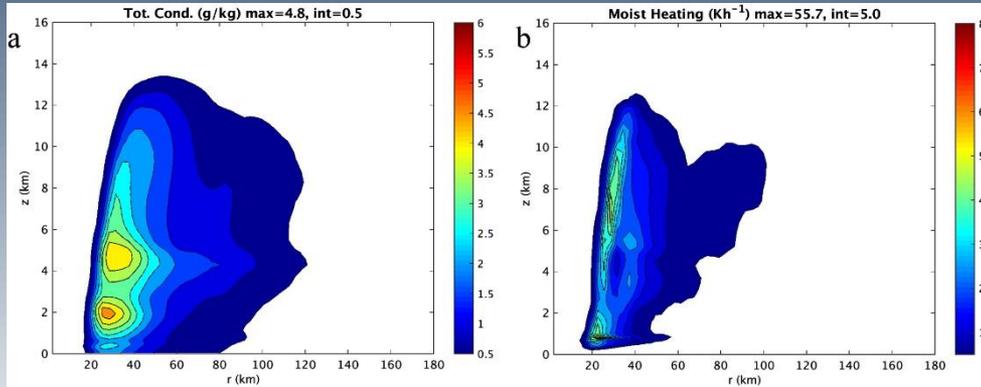
Morrison, Ideal 3 and 5, $r=0$ to $r=166.5$ km



Different microphysics schemes

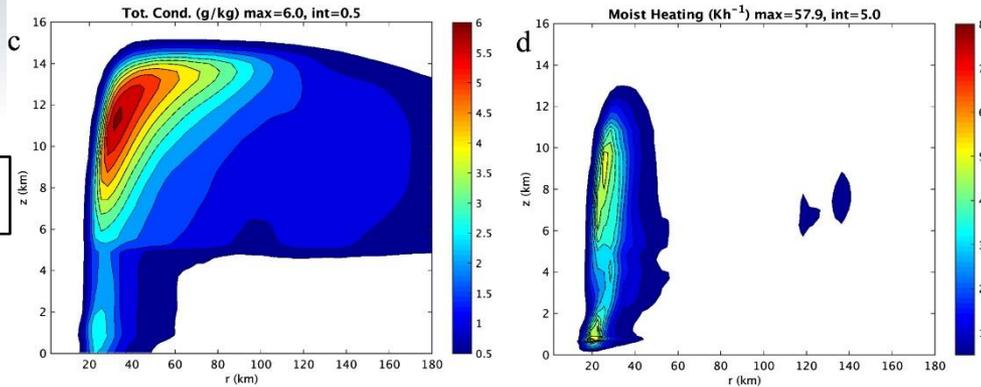
Moist diabatic heating rates

WDM6



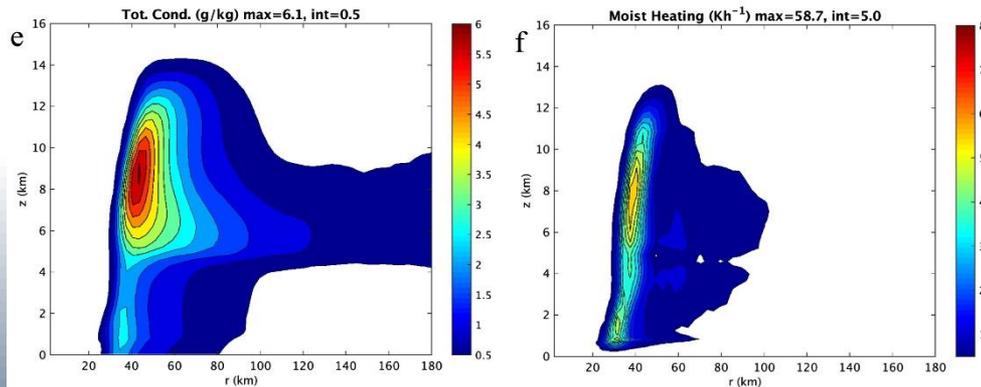
WDM6 has a fairly limited region of condensate, with values exceeding 4 g kg^{-1} only in the eyewall and below 6 km altitude.

Thompson



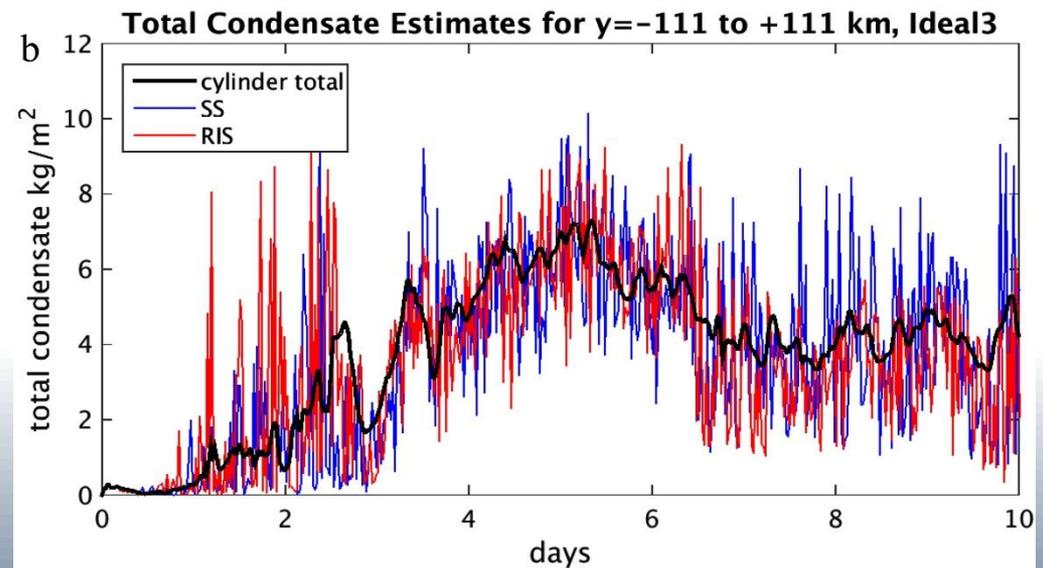
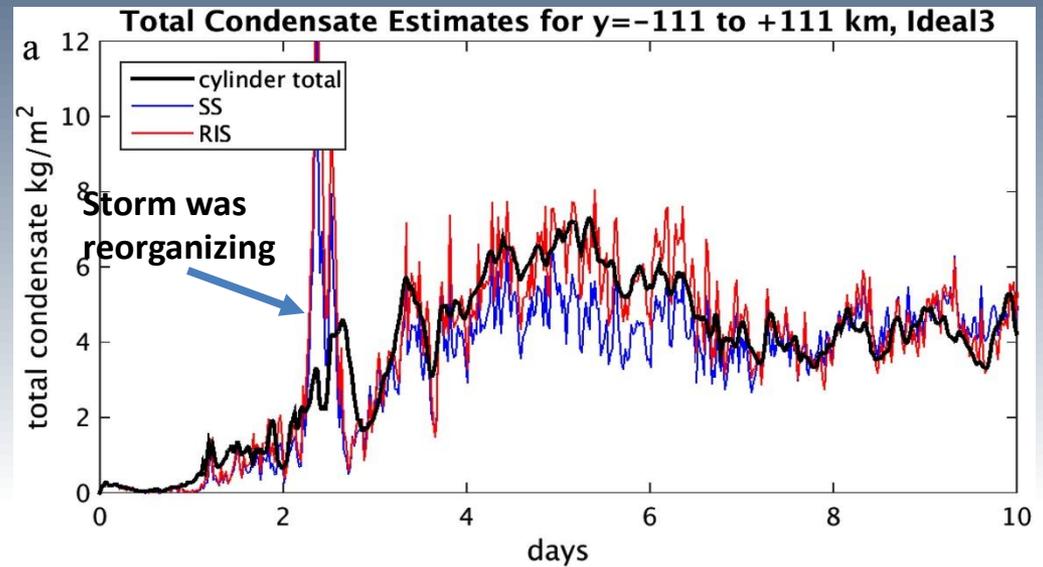
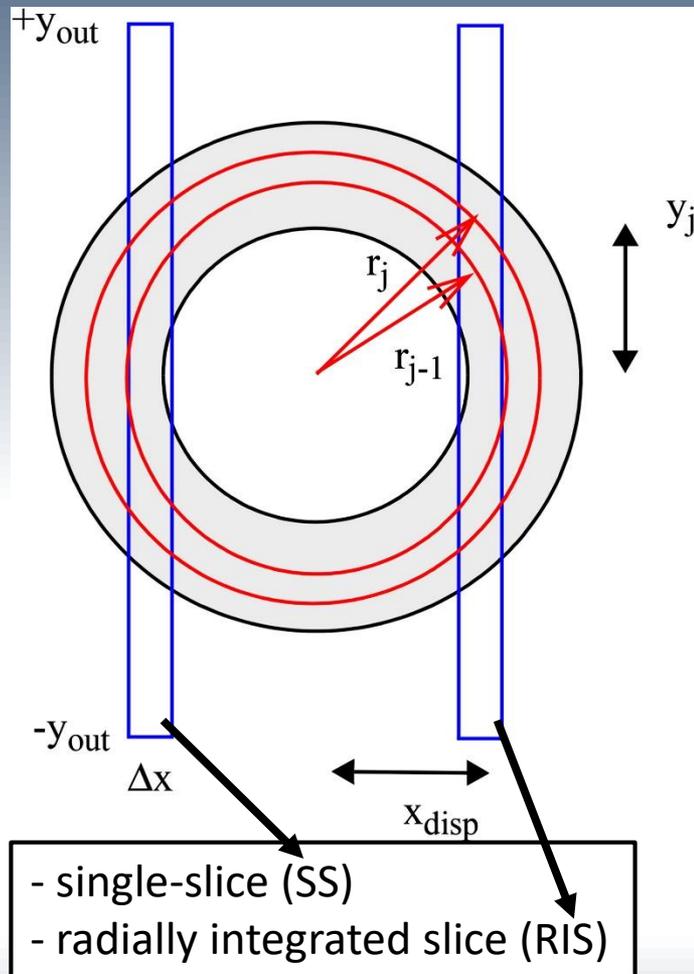
The Thompson scheme has a vast region of higher concentrations, all above 8 km height and extending out beyond 60 km radius.

Morison



The distribution for the Morrison scheme is somewhat between the two, but closer to the Thompson scheme.

Narrow-swath measurements



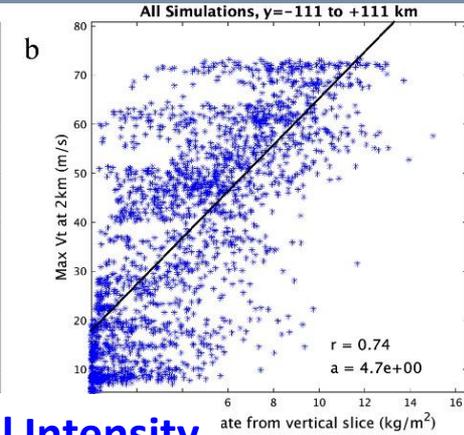
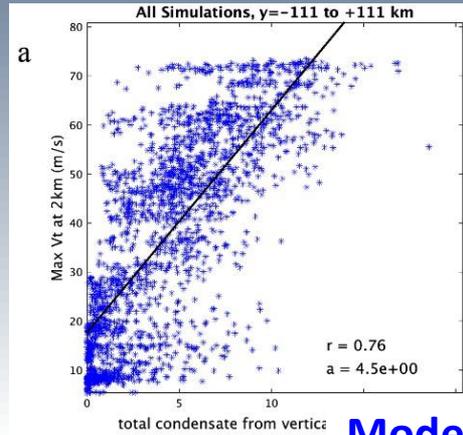
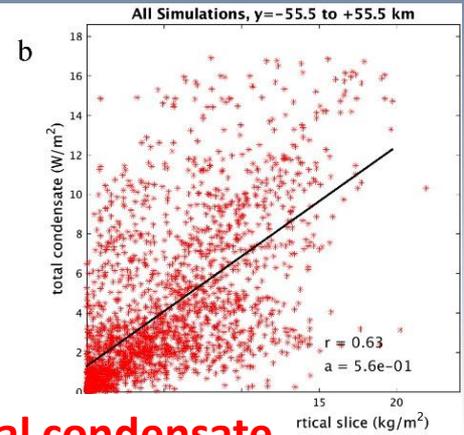
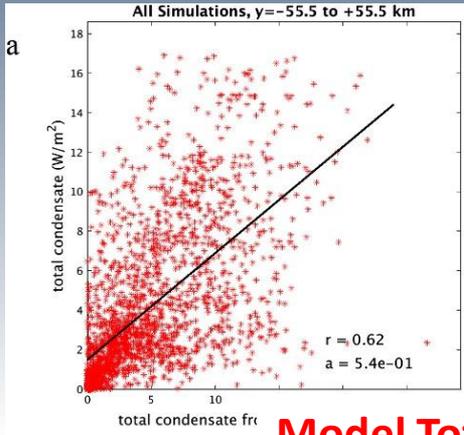
SS and RIS retrieved total condensate

SS

RIS

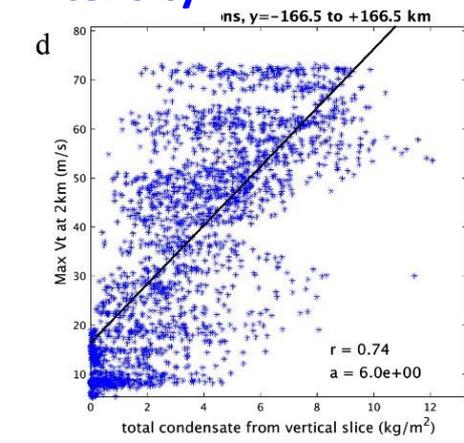
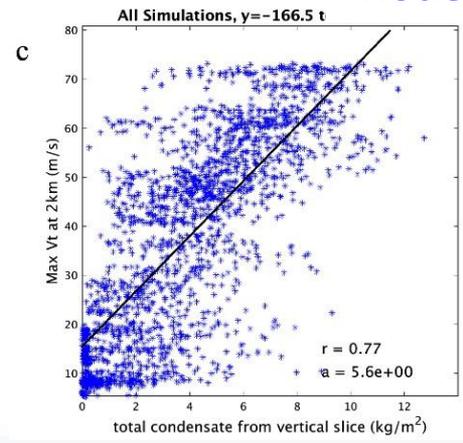
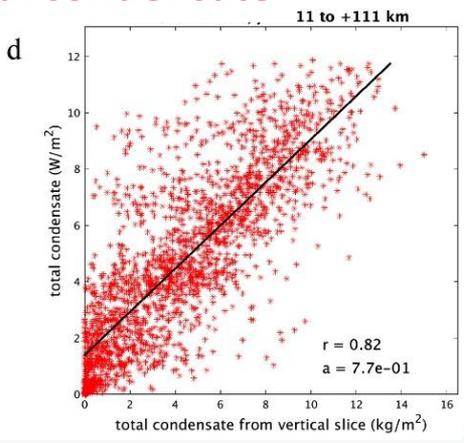
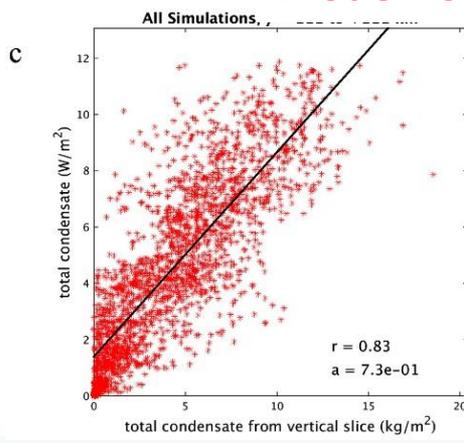
SS

RIS



Model Total condensate

Model Intensity



Conclusions

- ☀️ The numerical simulations used in this study show a very close correlation between **total condensate**, **total moist diabatic heating** and **current TC intensity** in the same volume.
- ☀️ **Moist heating above certain altitudes** is **not a reliable indicator** of imminent intensity change
- ☀️ The results were **not particularly sensitive to the radius of the cylinder** over which the correlations were made (a radius of 111 km generally showed the strongest correlations).
- ☀️ These weak dependencies suggest **compensating factors** in the relationship between total heating within a given radius and TC intensity.

Thanks