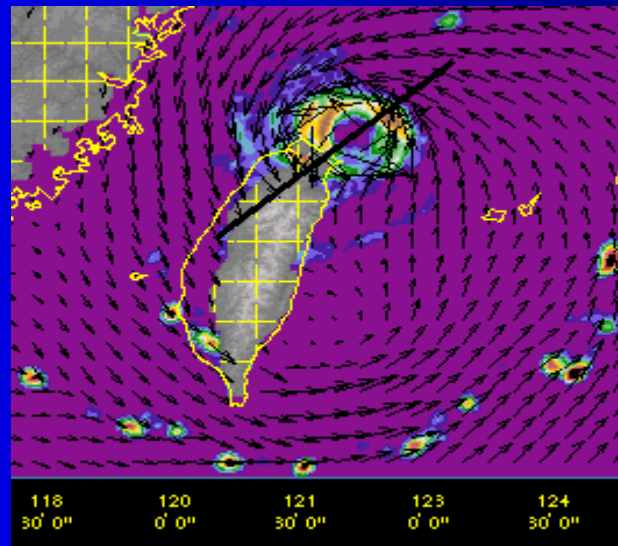


Terrain-Induced Asymmetric Structures of Typhoon Nari (2001)

Ming-Jen Yang 楊明仁

*Dept. of Atmospheric Sciences, Inst. of Hydrological & Oceanic Sciences
National Central University*



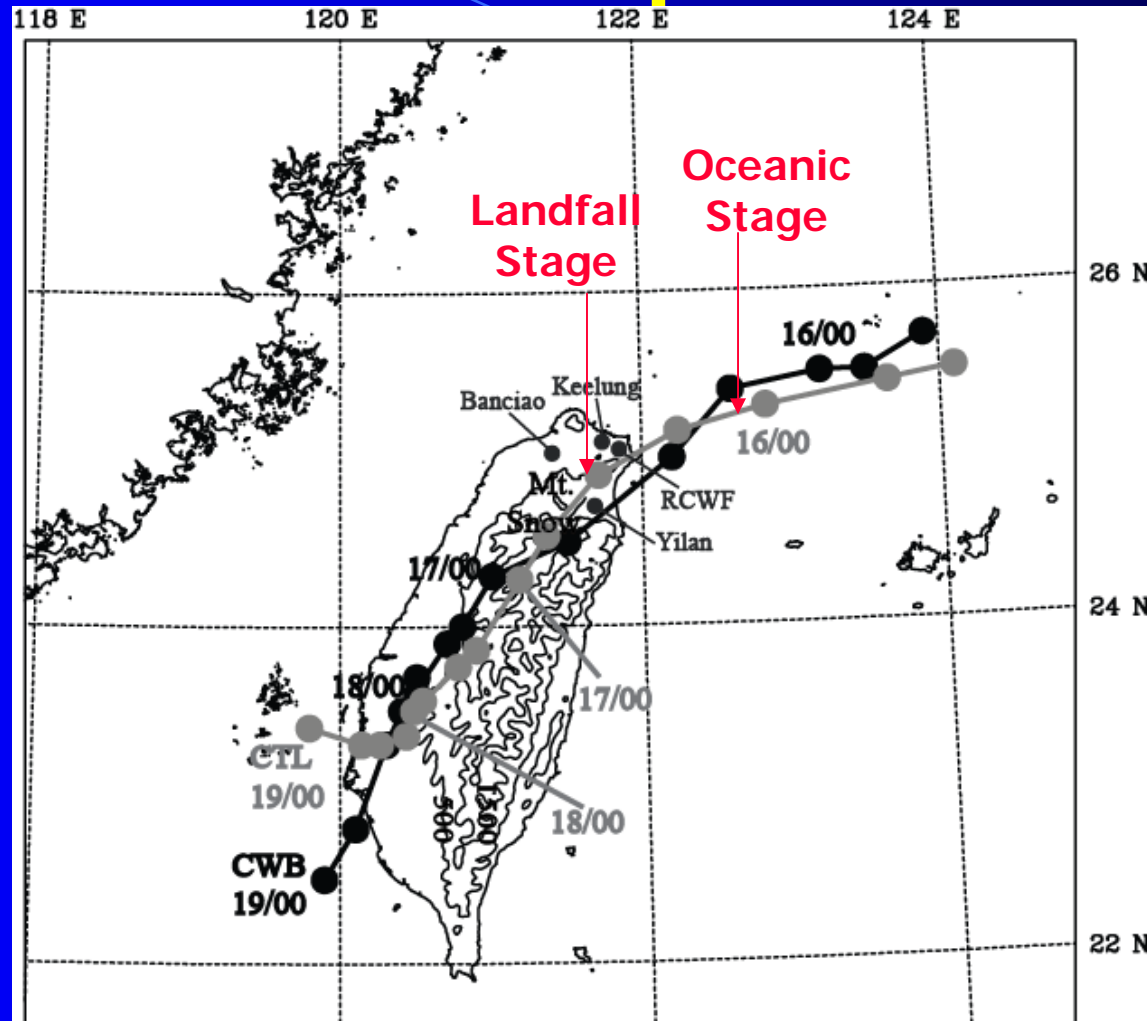
Collaborator: Da-Lin Zhang, Chi-Hsin Liao, and Xiao-Dong Tang

Seminar at NTU, 2009/10/06

Objectives

- 1) To document the evolution of Nari's precipitation, kinematic, and microphysics structures during its landfall on Taiwan
- 2) To investigate the terrain-induced asymmetric structures of Nari
- 3) To examine the physical mechanism responsible of the sloping mid-level radial outflow over the topography

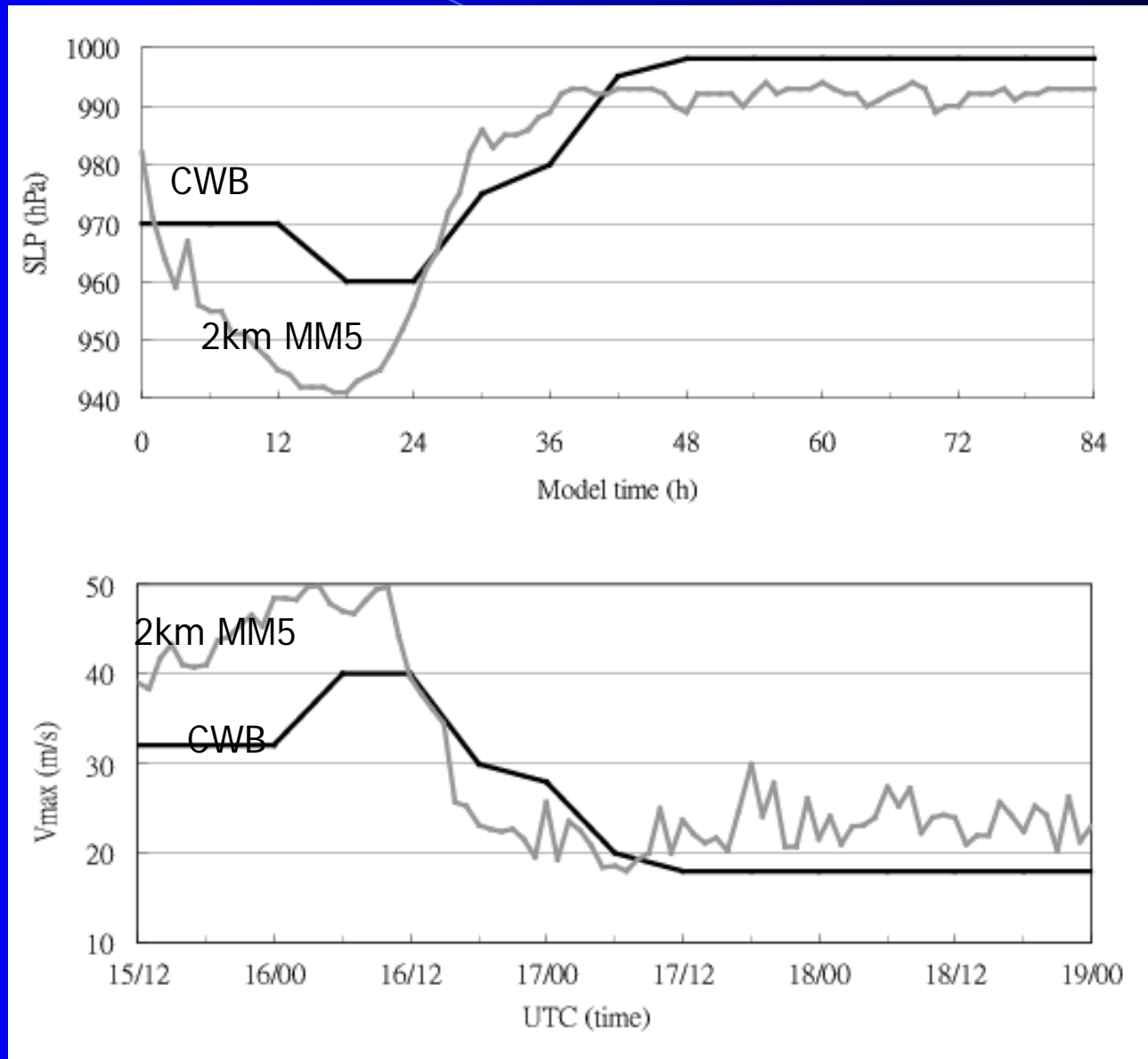
Track Comparison



Yang et al.
(2008; JAS)

Simulation time (hr)	12	24	36	48	60	72	84
Track error (km)	43.3	61.2	26.8	13.4	12	8.5	104.8

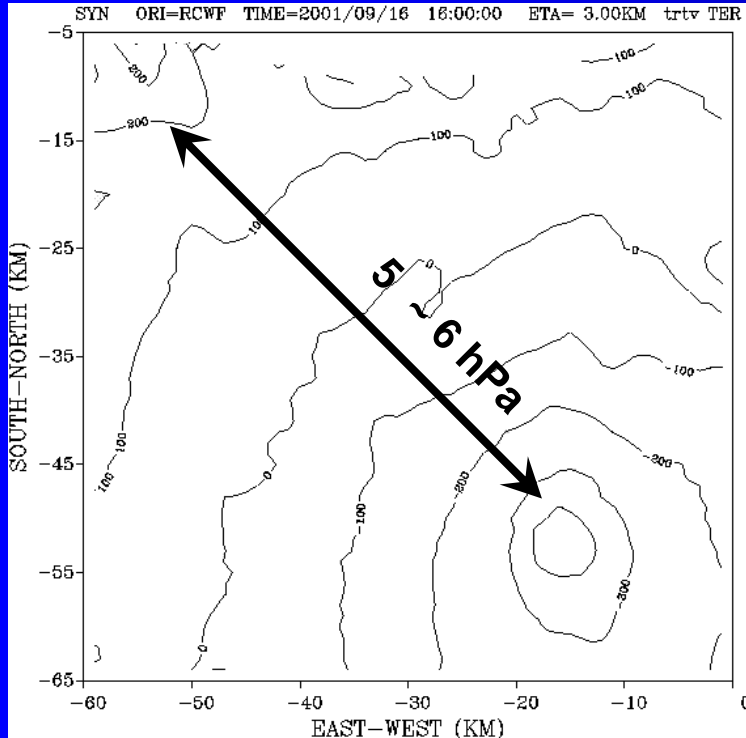
Time Series of SLP and Vmax



Yang et al.
(2008; JAS)

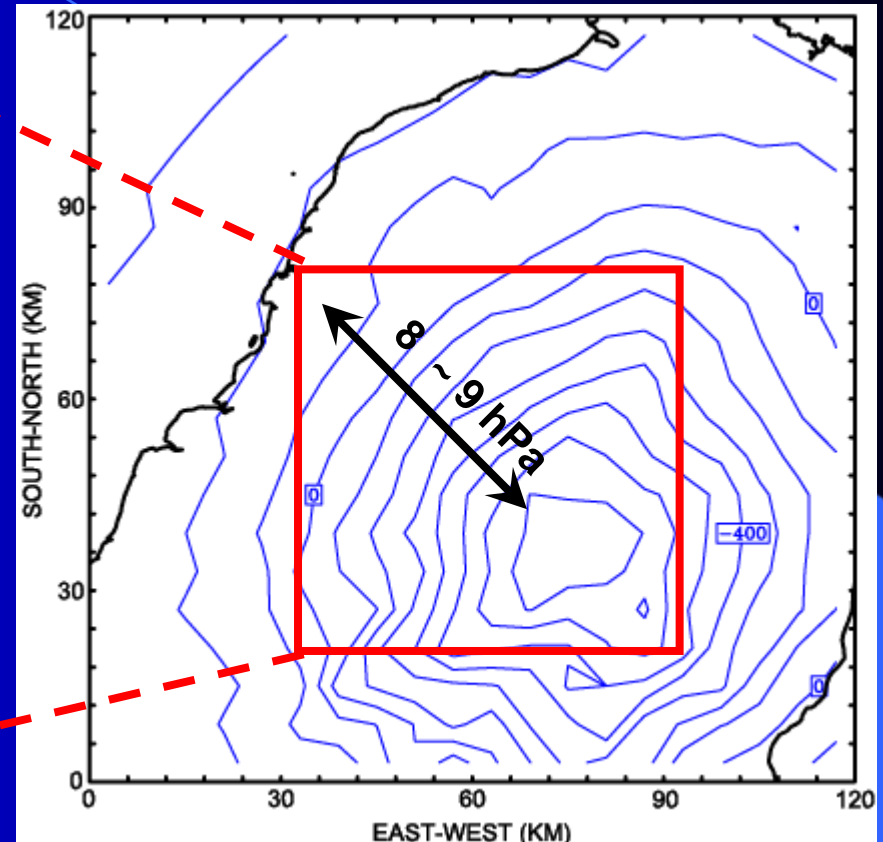
Horizontal Cross Section of Pressure Perturbations

0916_1400 UTC



Radar Retrieval (wrt. a Station Sounding)

Courtesy of T.-C. Chen and Y.-C. Liou

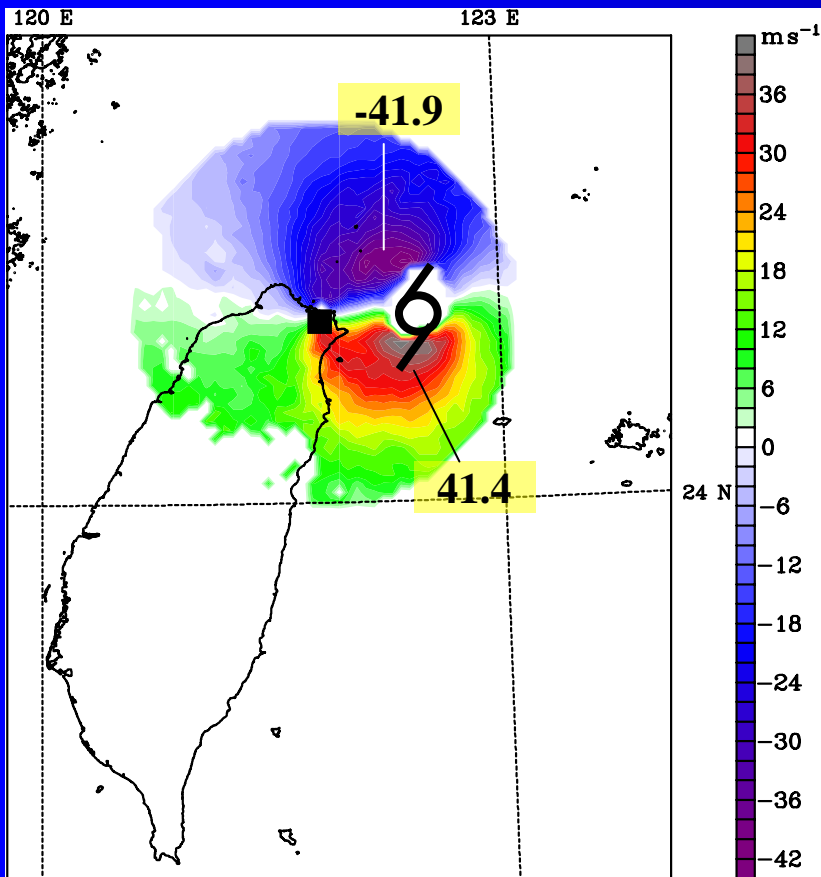


MM5 Simulation (wrt. a Horizontal Area Mean)

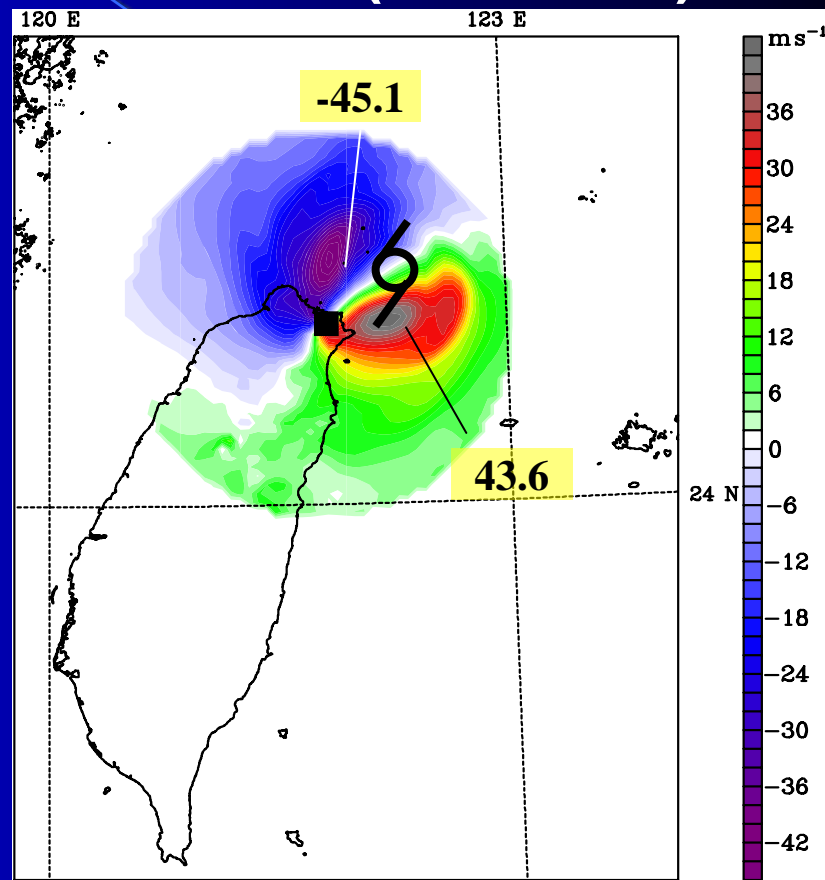
Radial Wind wrt RCWF Radar

@ 3 km Height

Obs Vr (6 km pixel)



MM5 Vr (dx = 6 km)



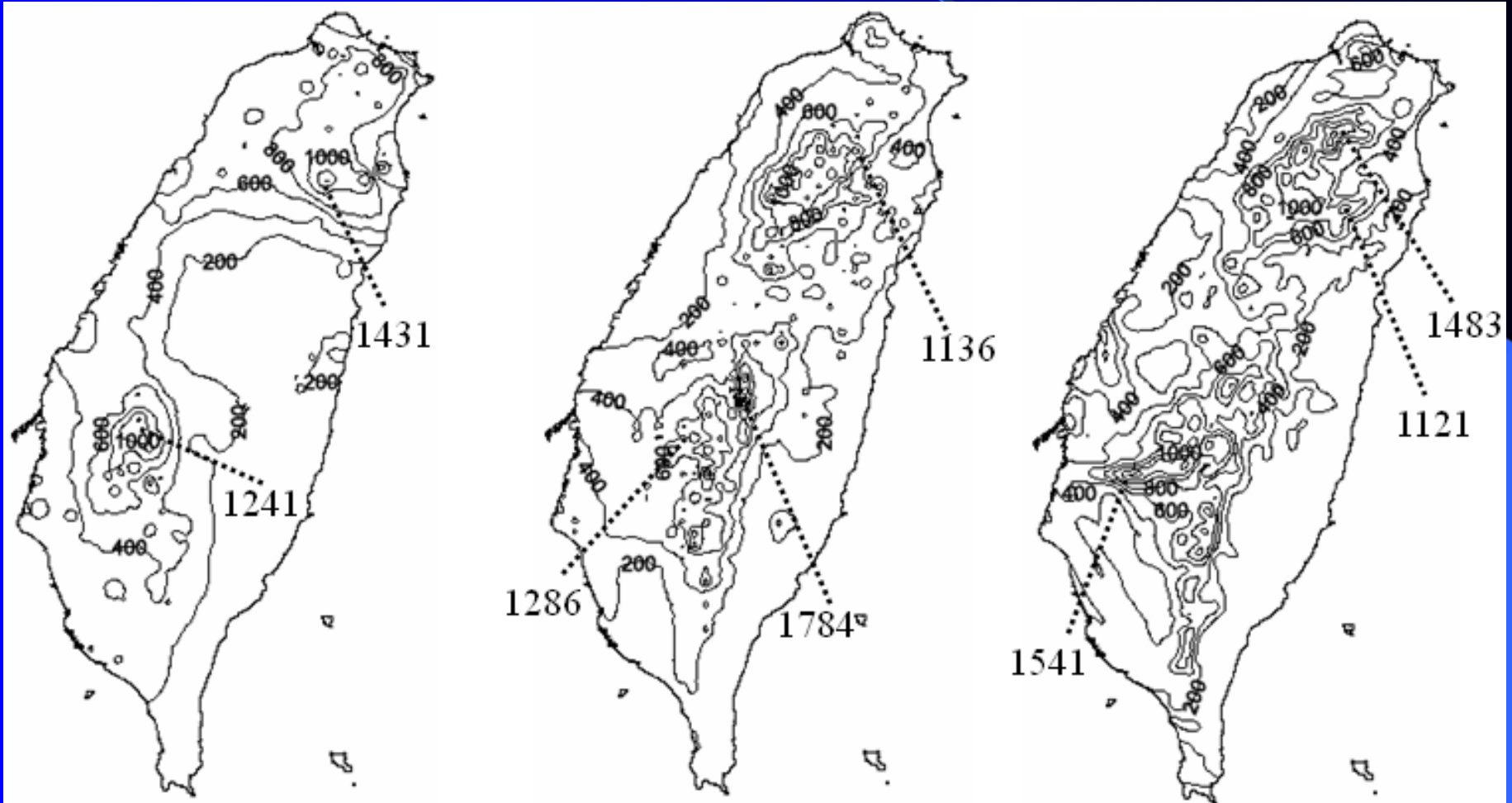
Courtesy of T.-C. Chen and Y.-C. Liou

3-day rainfall (09/16~09/18)

OBS

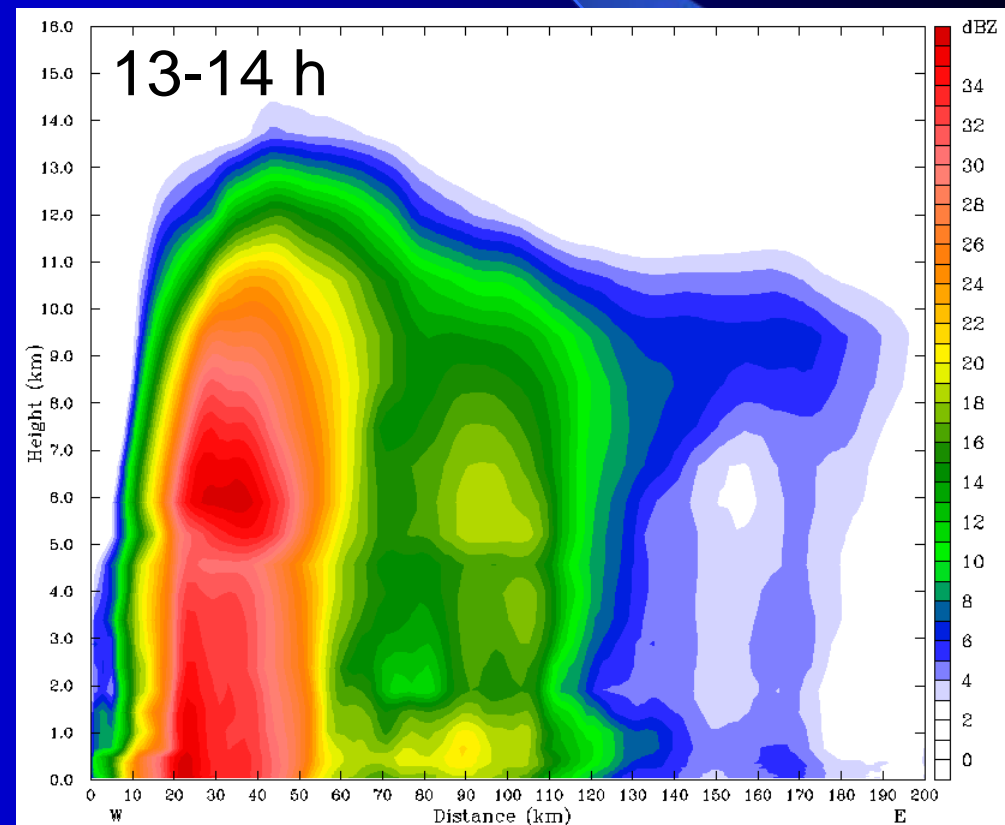
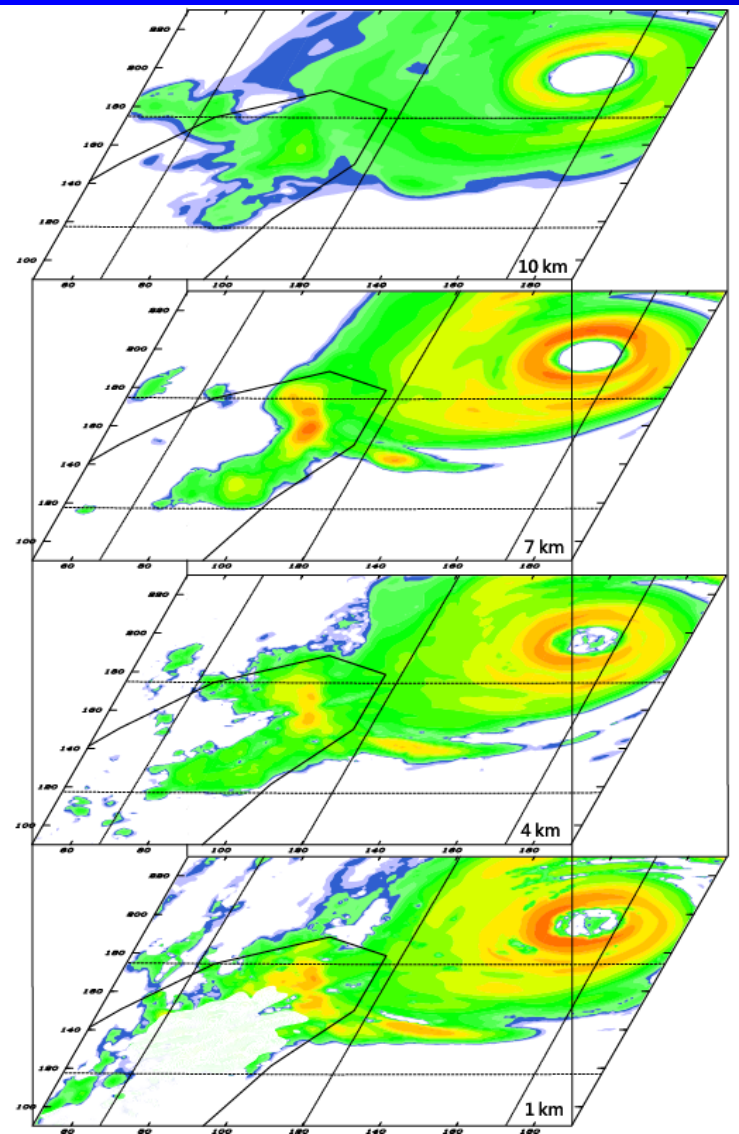
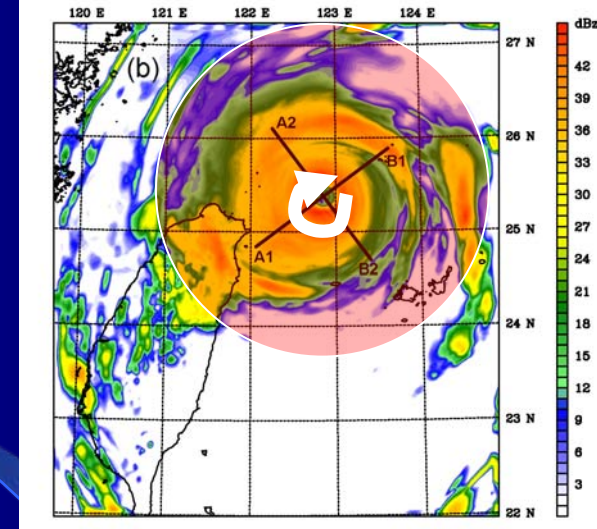
6km MM5

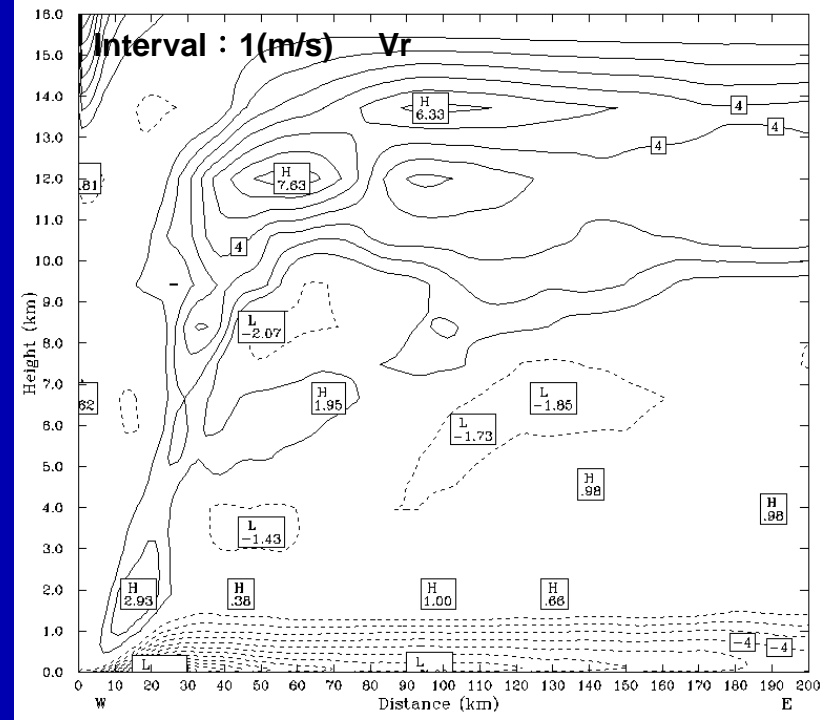
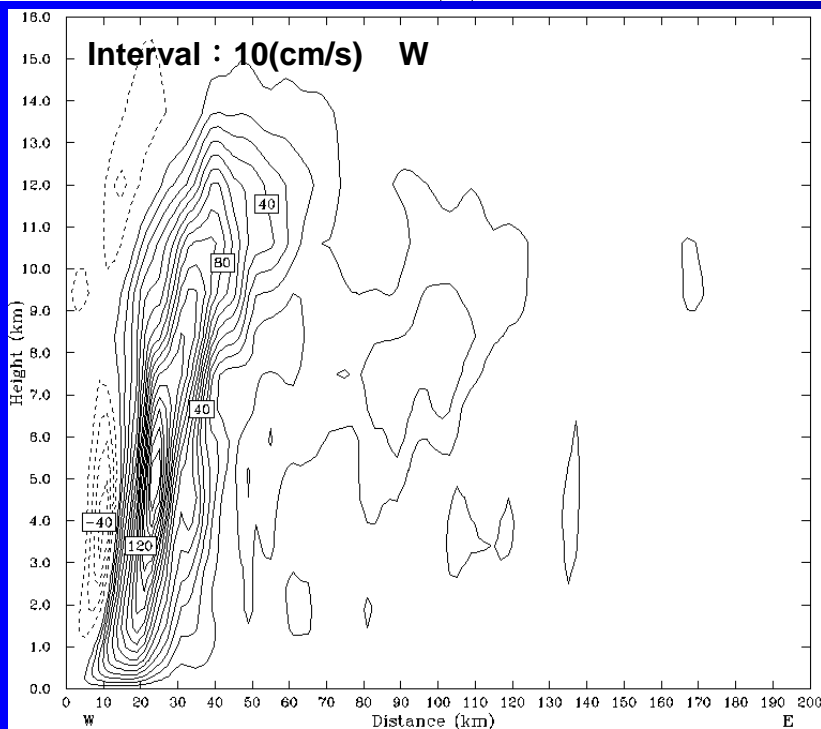
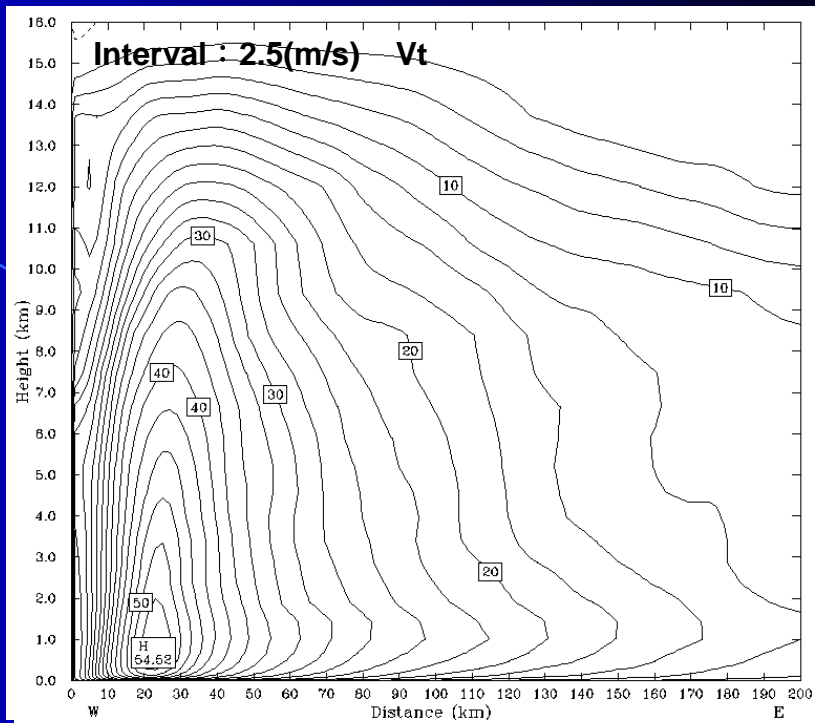
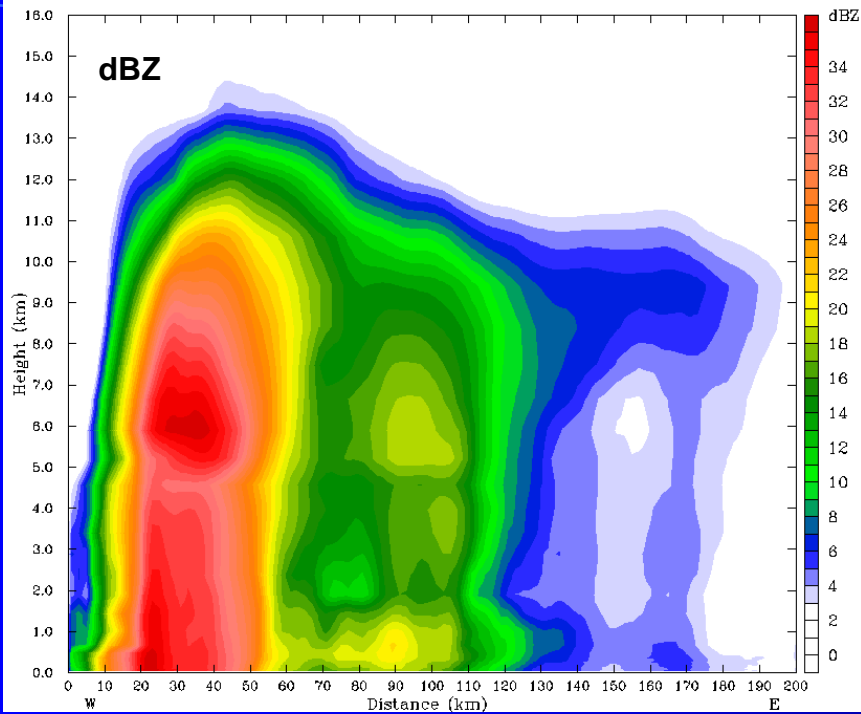
2km MM5

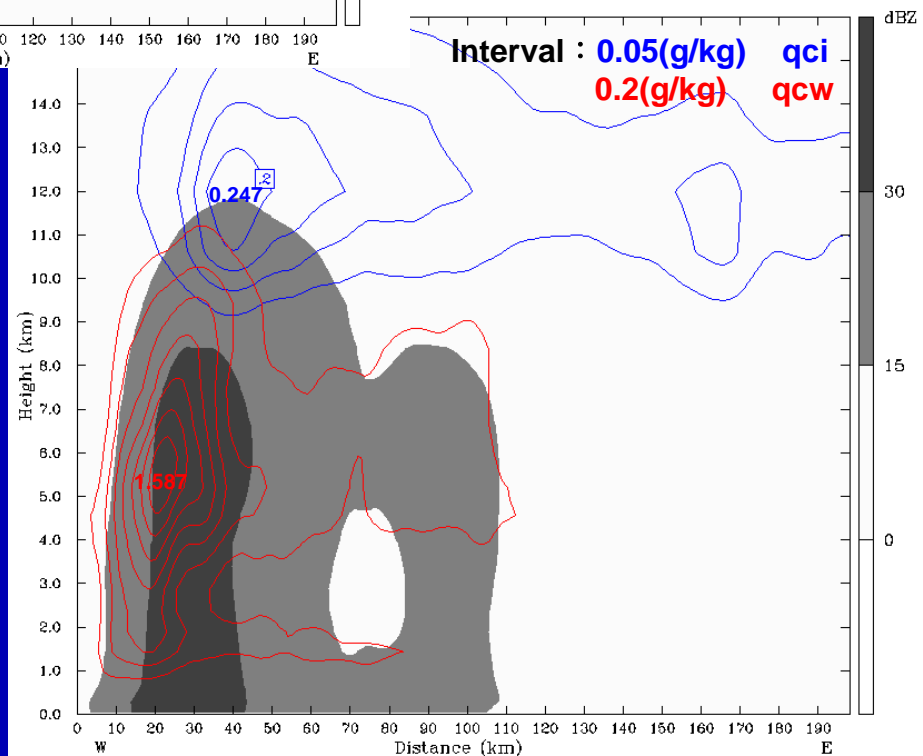
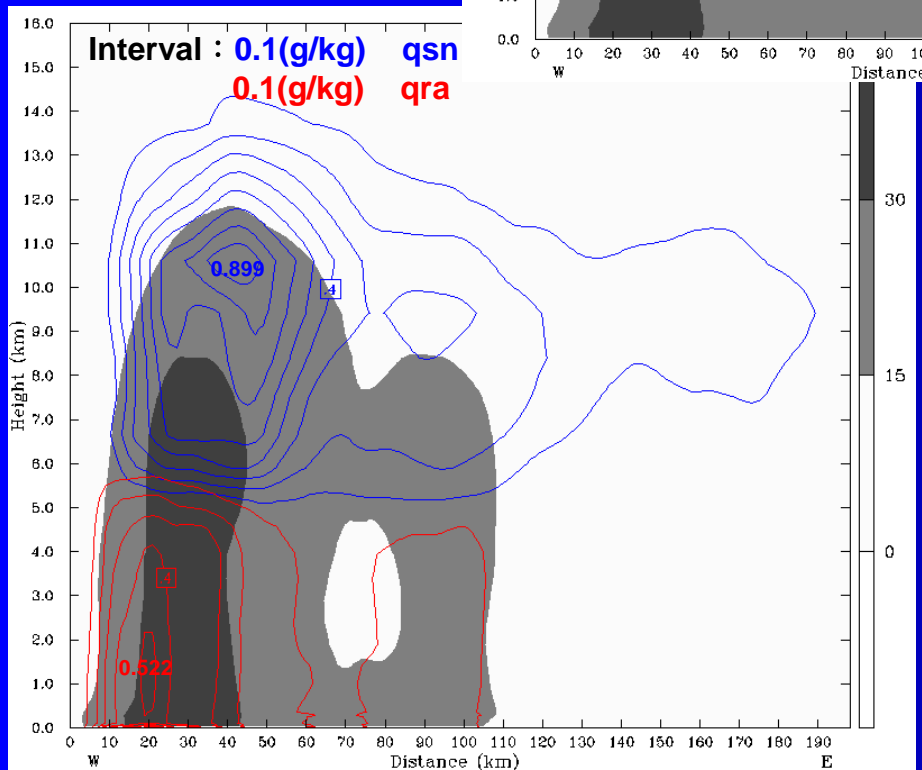
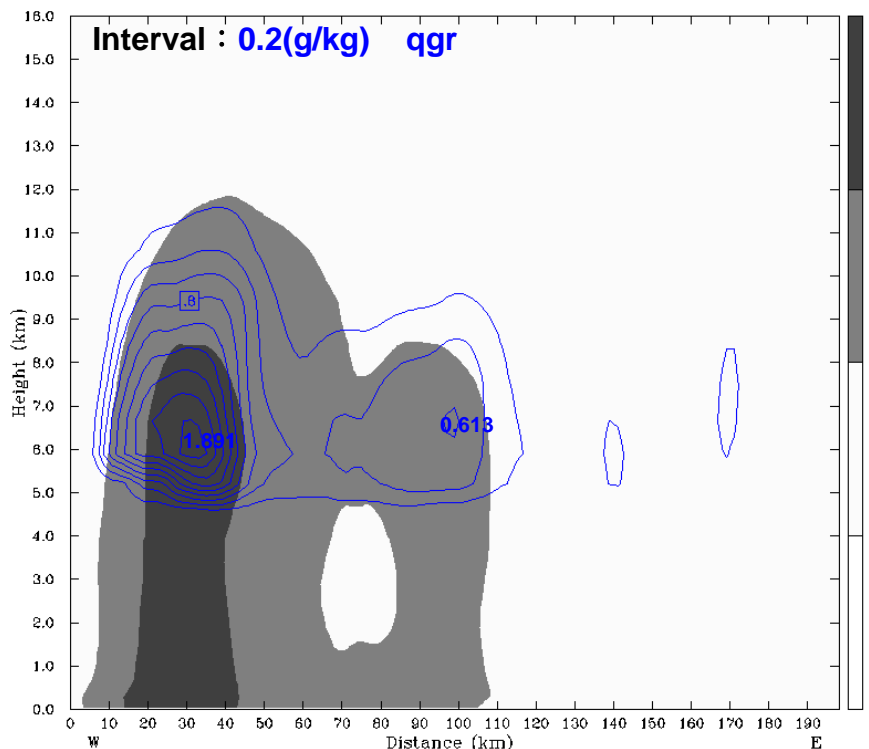


Oceanic Stage (9 Hours Before Landfall)

Azimuthal-avg. structure ($r=200$ km) while Nari is over ocean



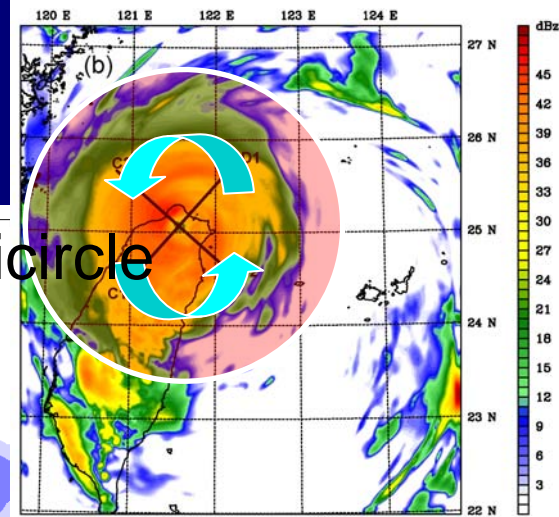




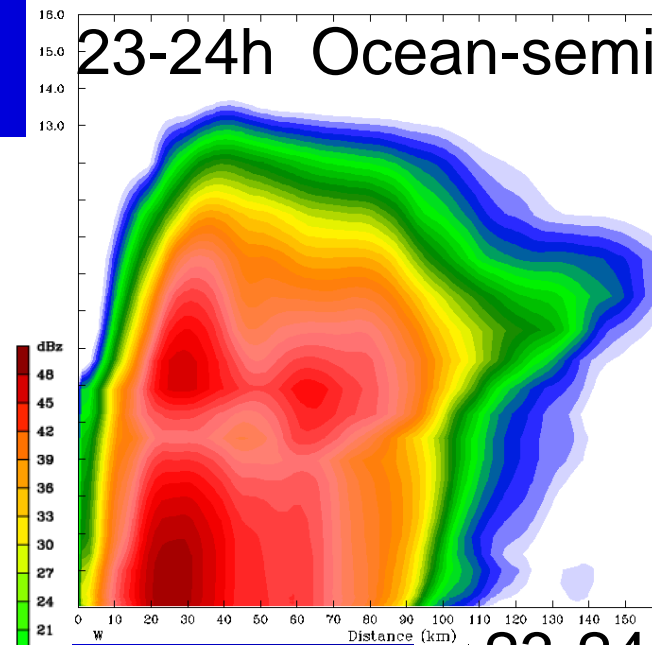


**Landfall Stage
(2 Hours After Landfall)**

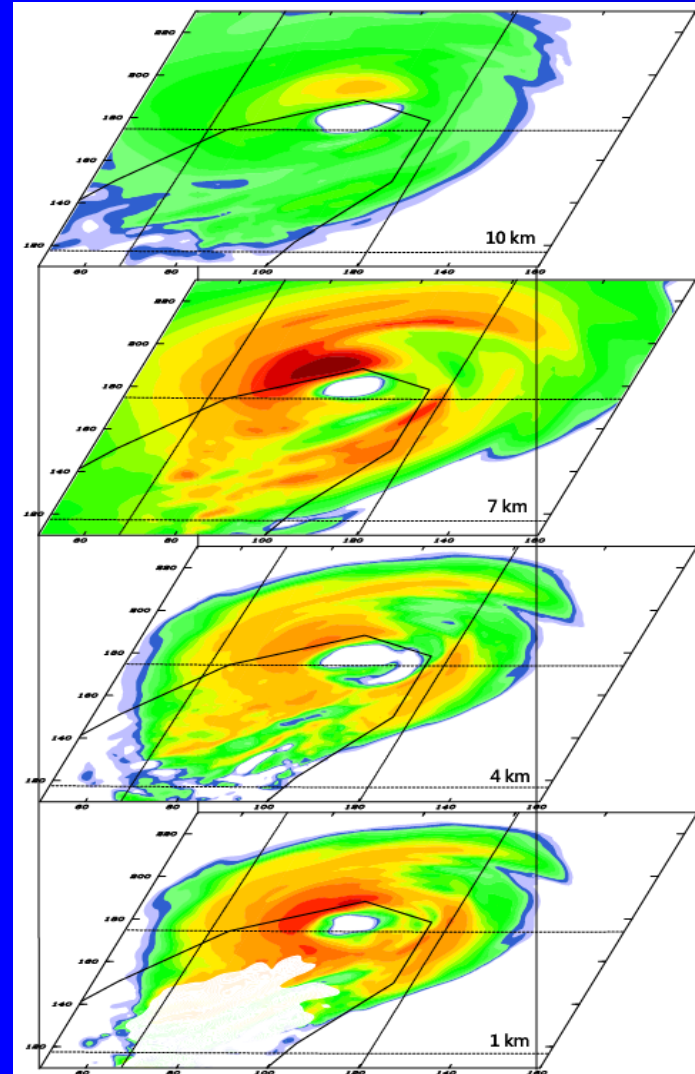
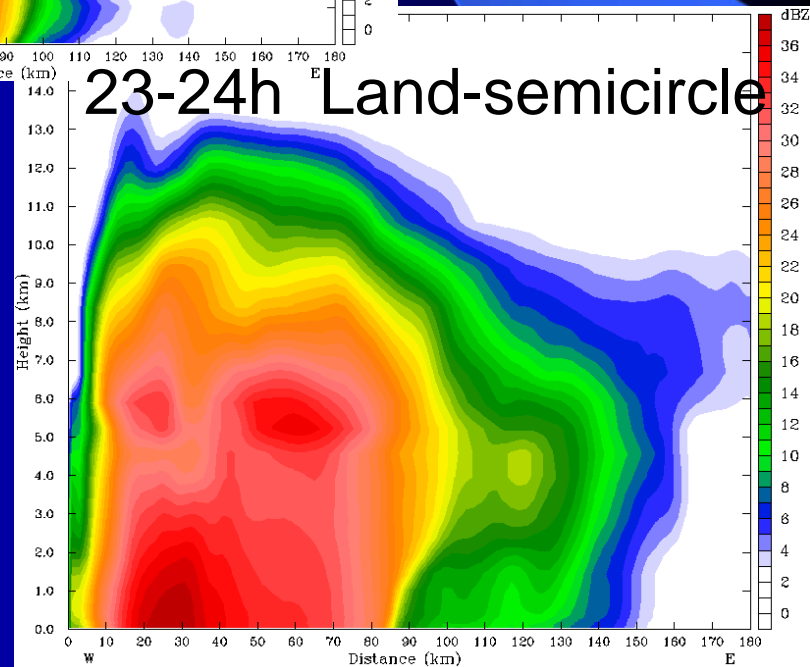
Semicircle-avg. structure ($r=180$ km) Nari is at landfall

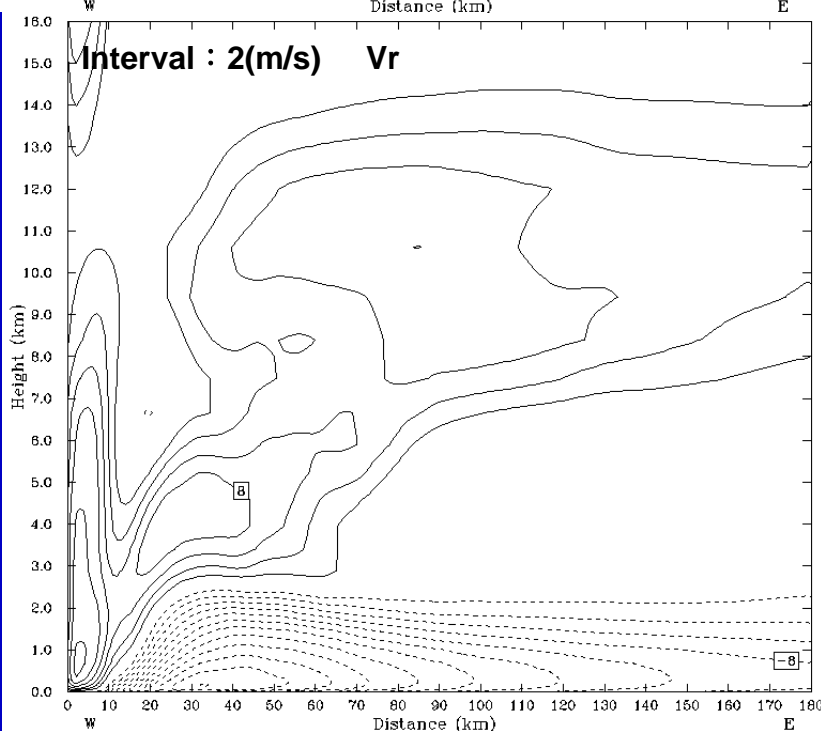
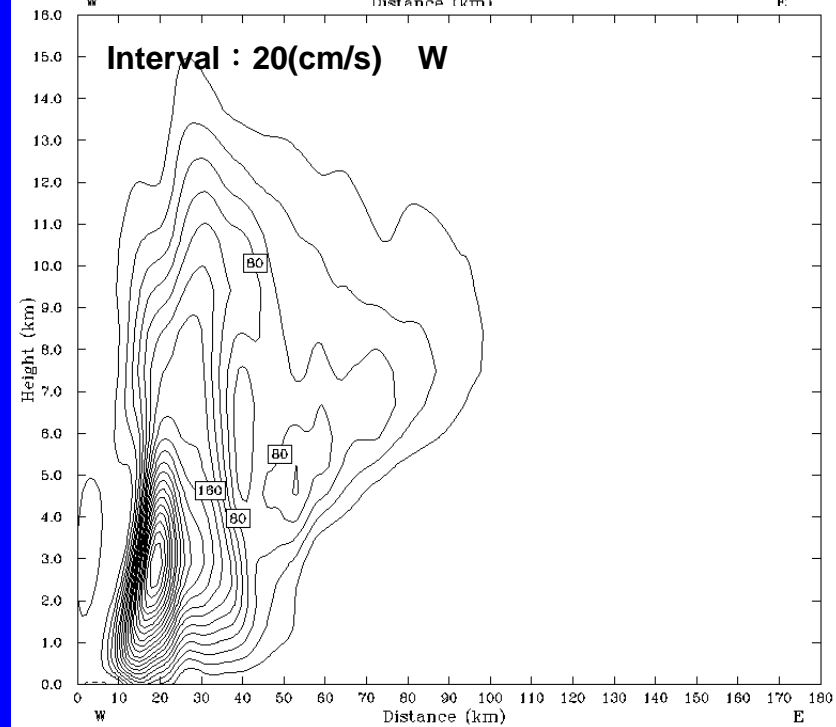
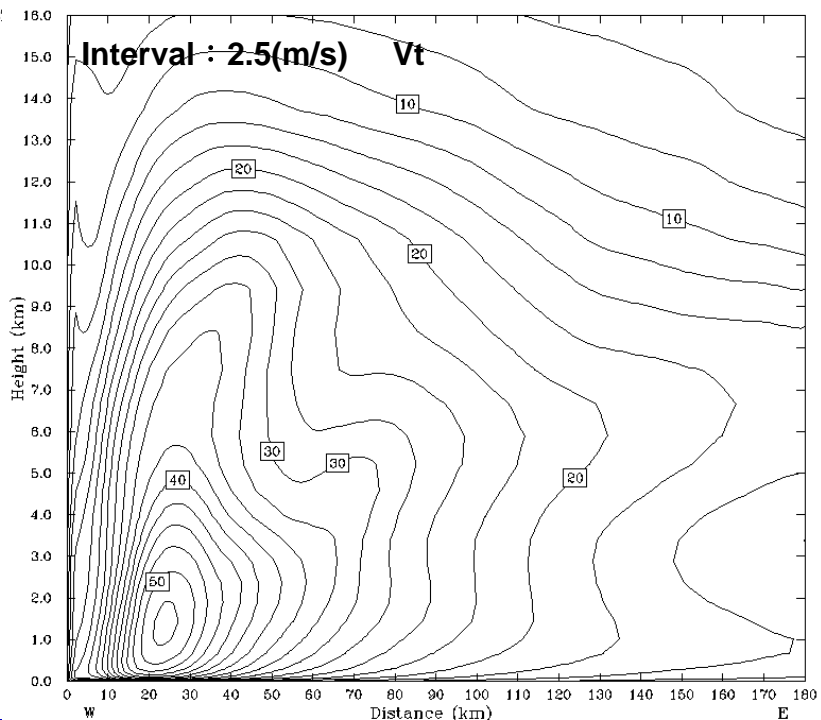
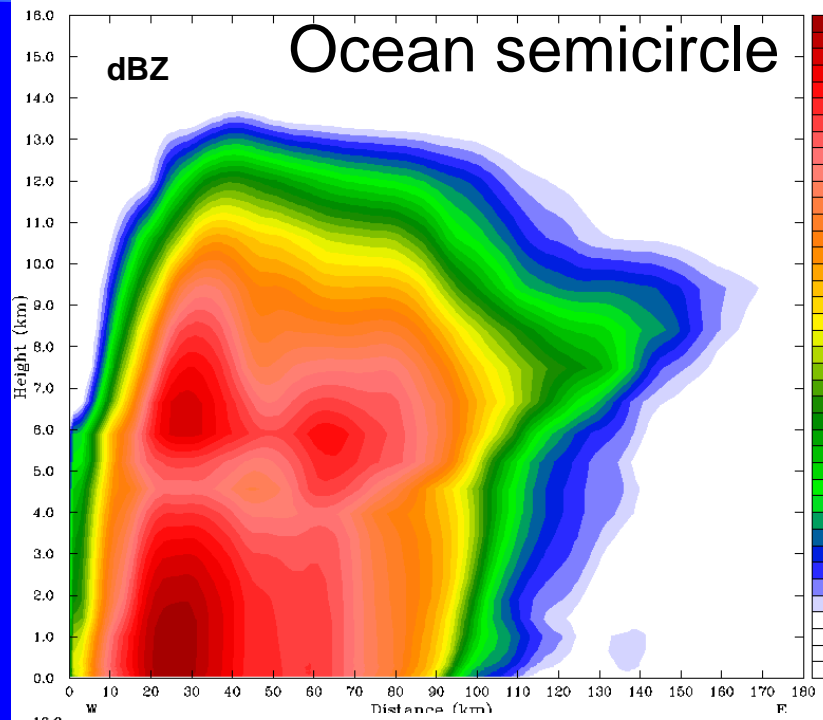


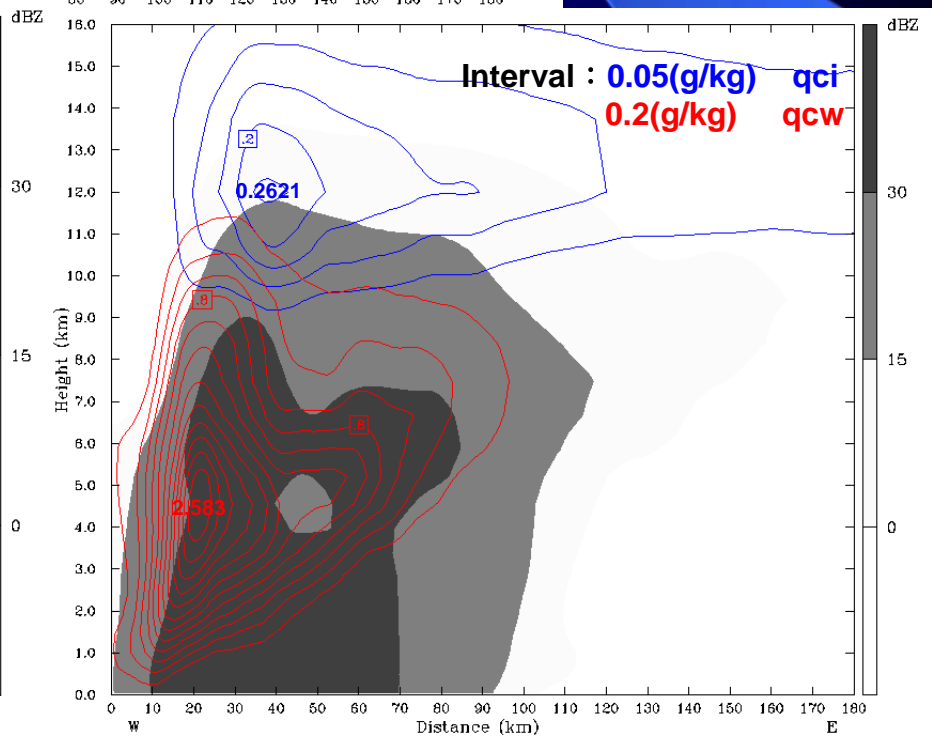
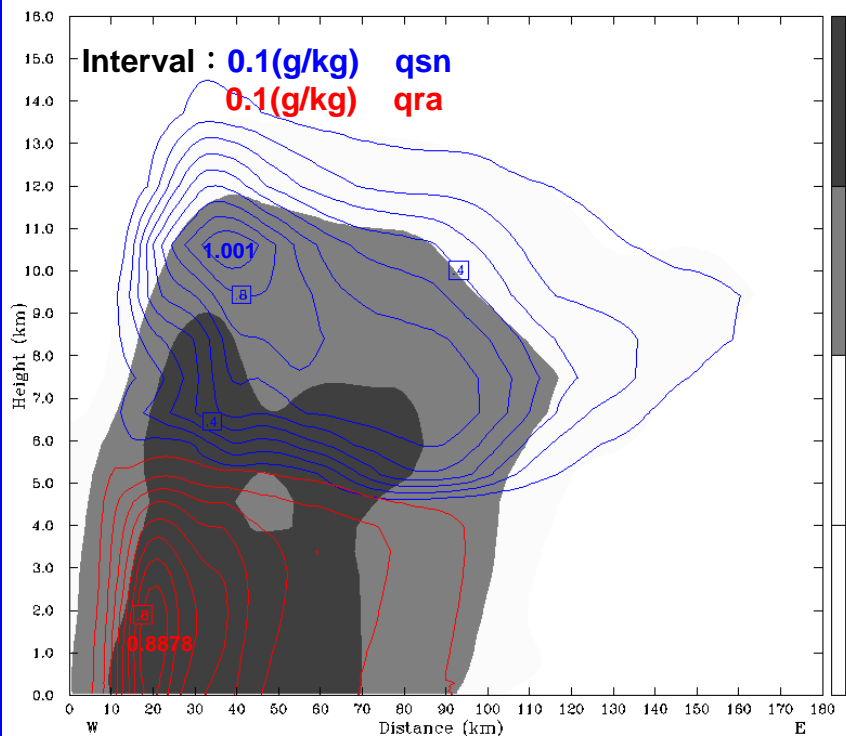
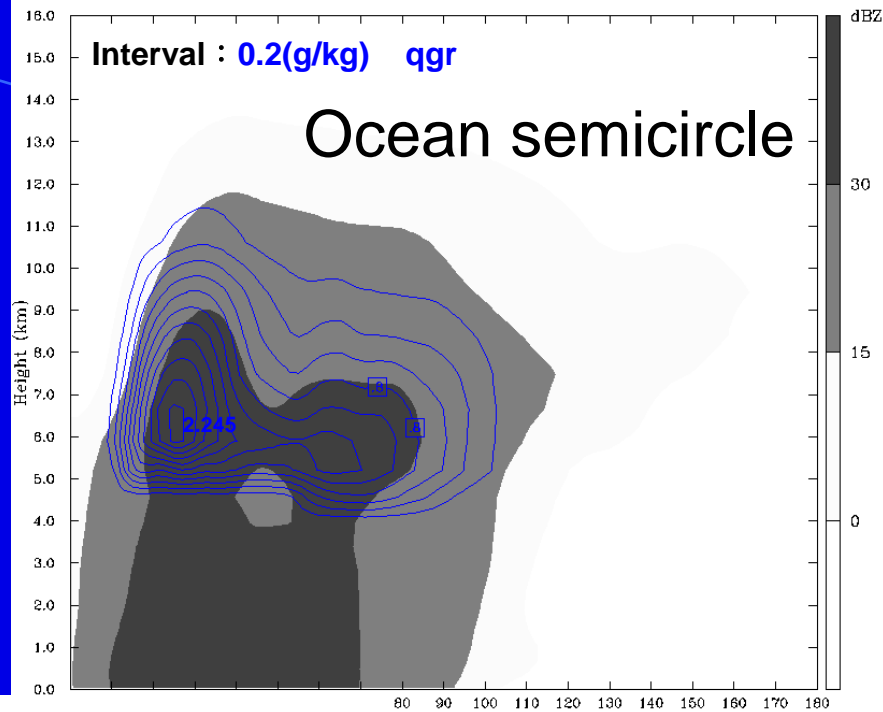
23-24h Ocean-semicircle

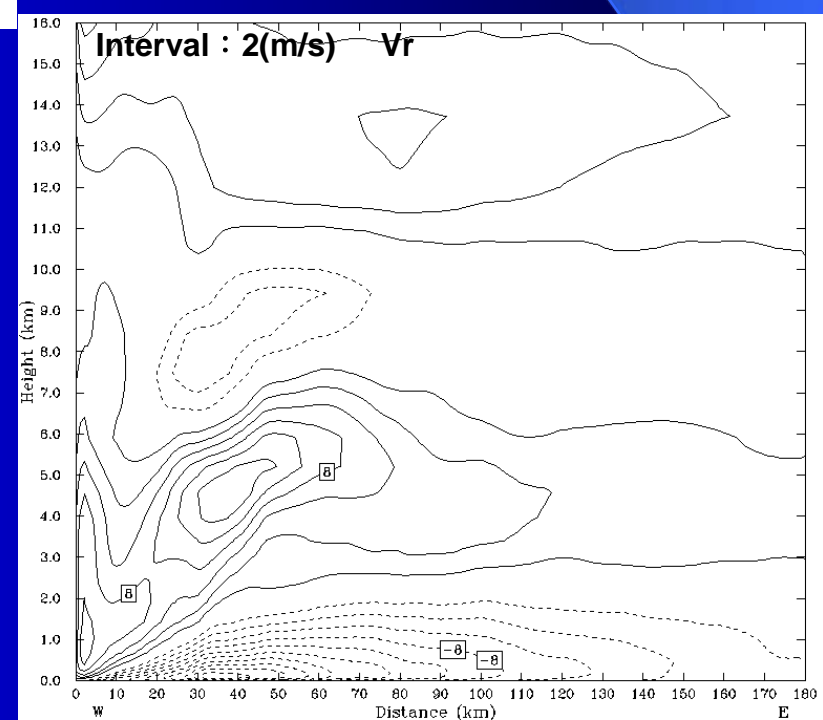
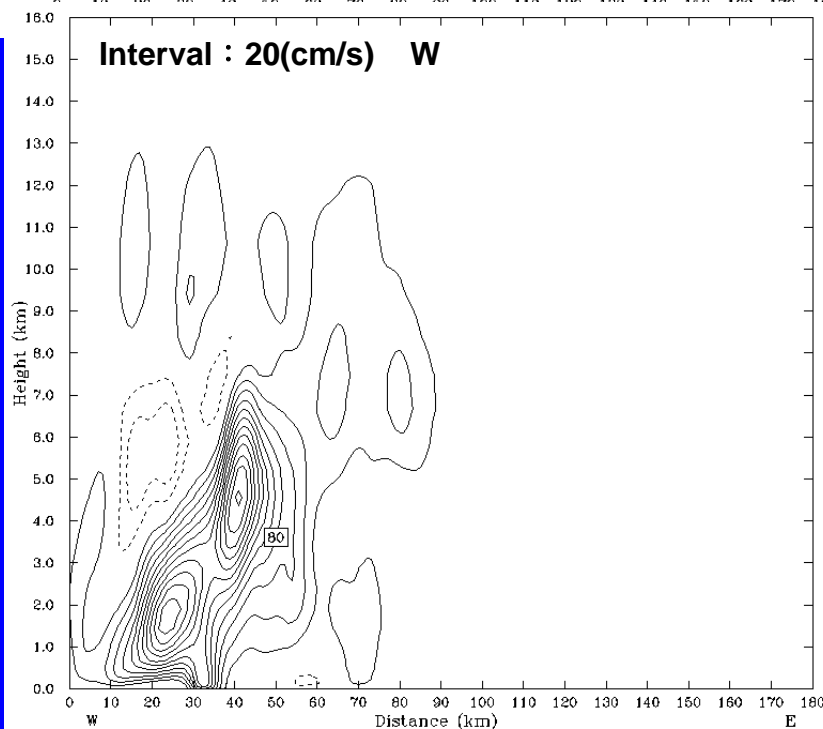
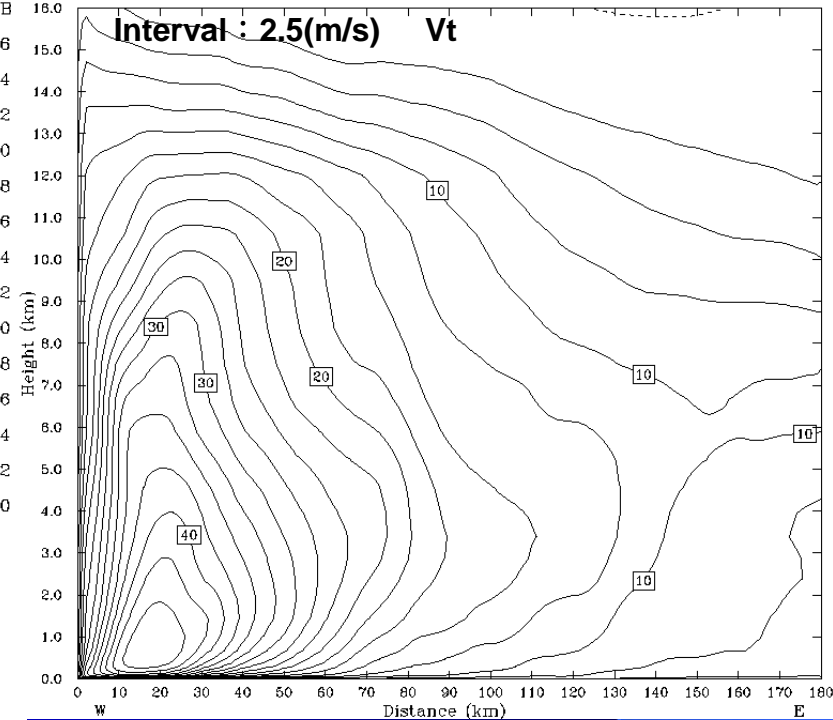
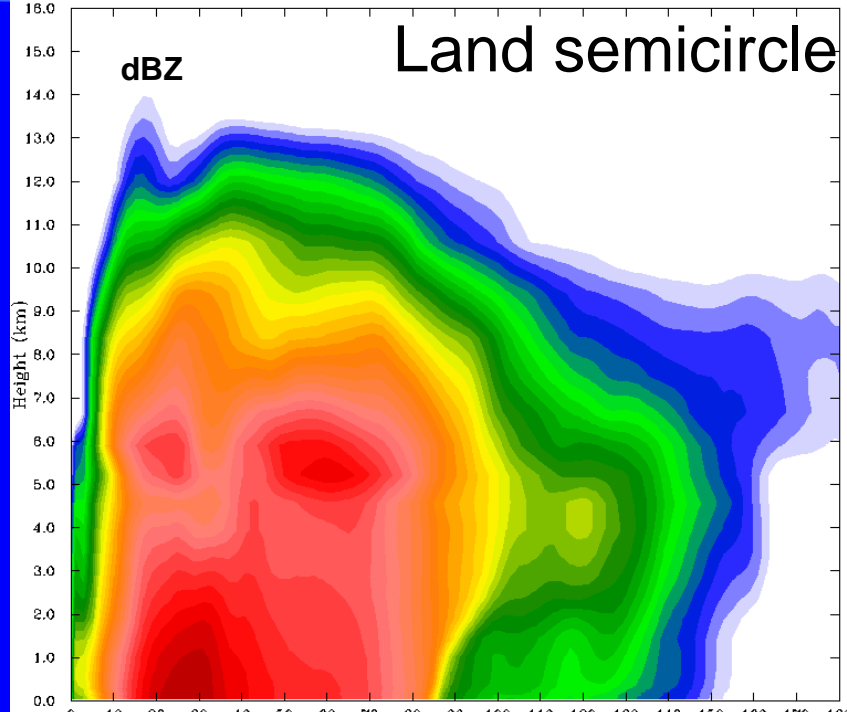


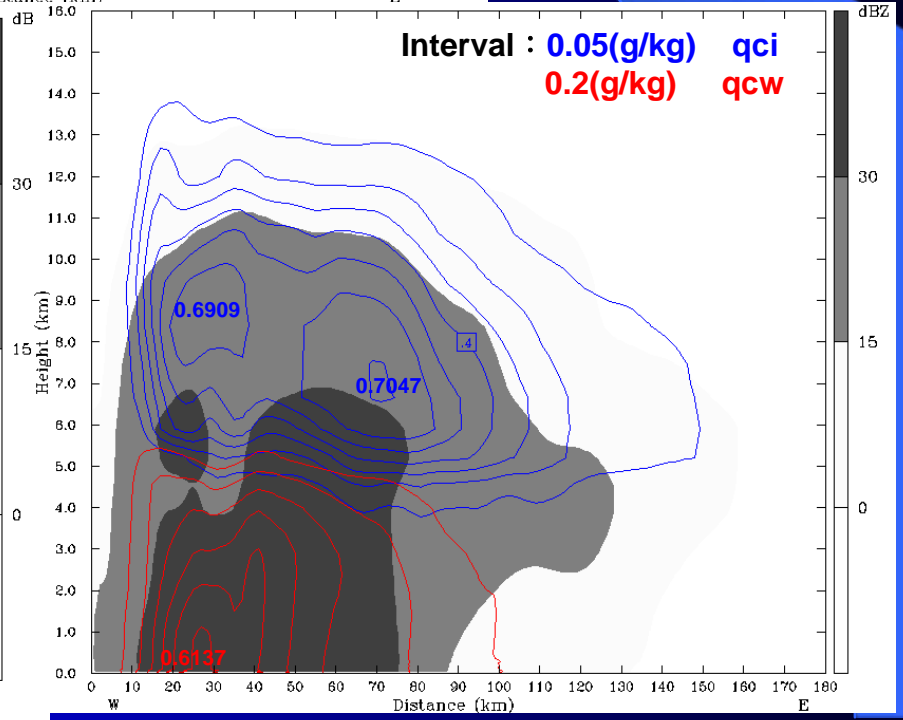
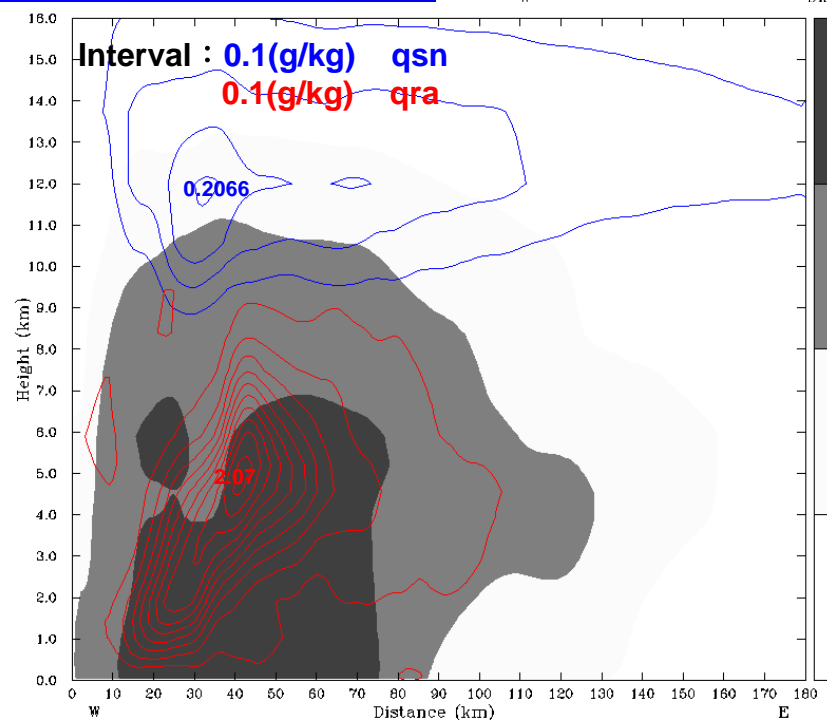
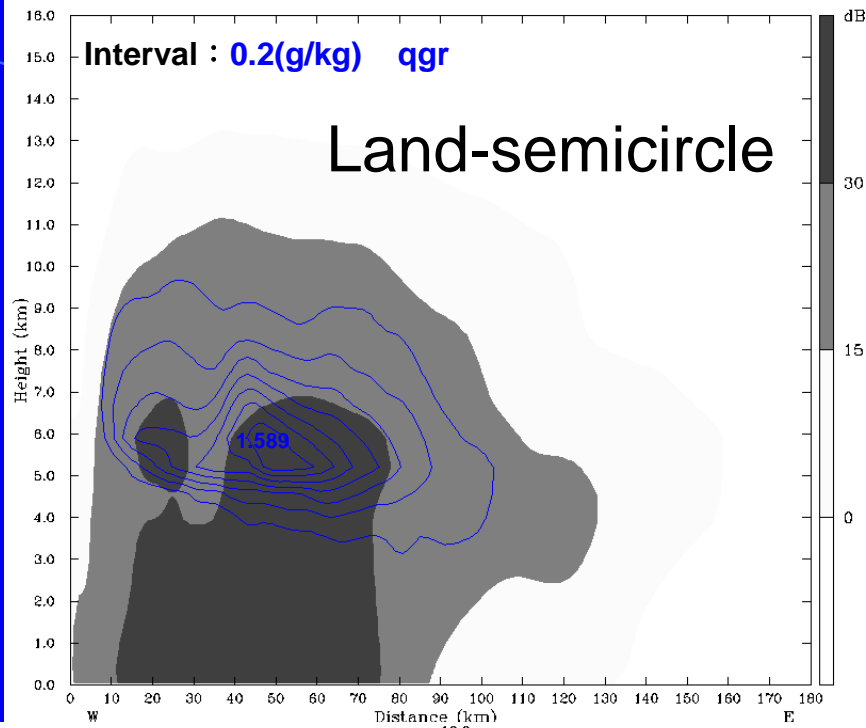
23-24h Land-semicircle

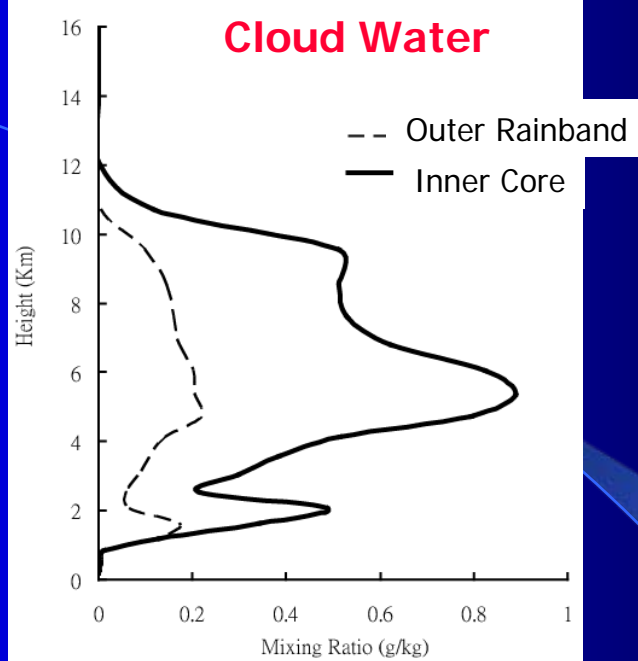
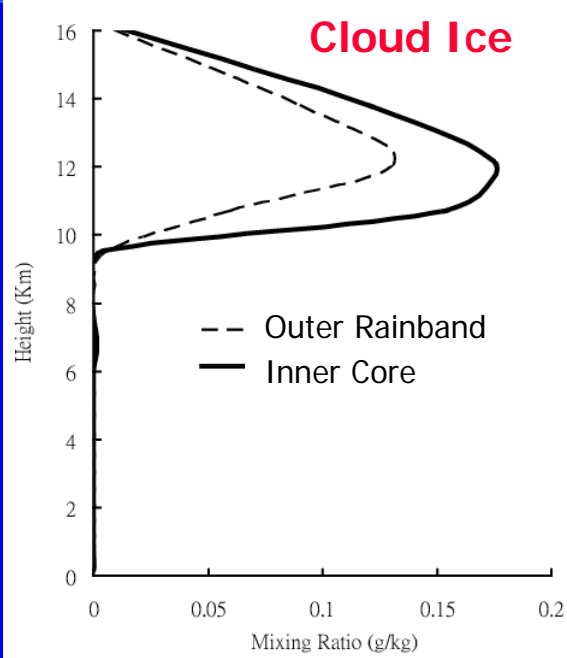




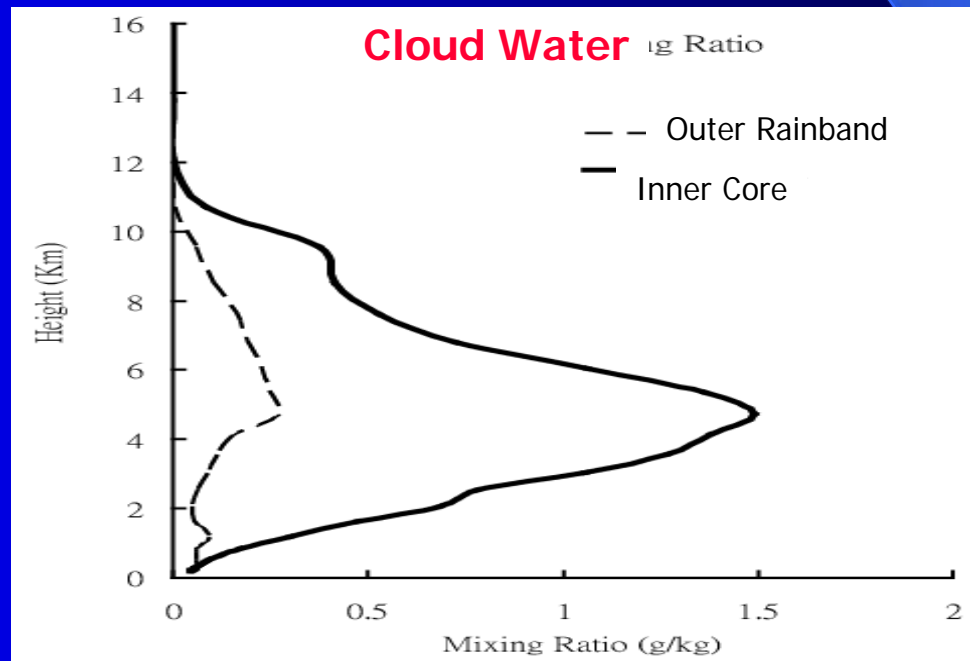
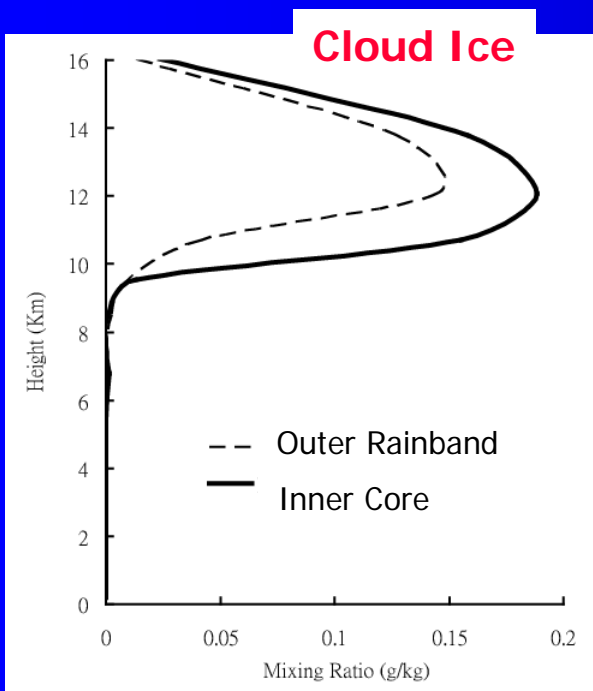




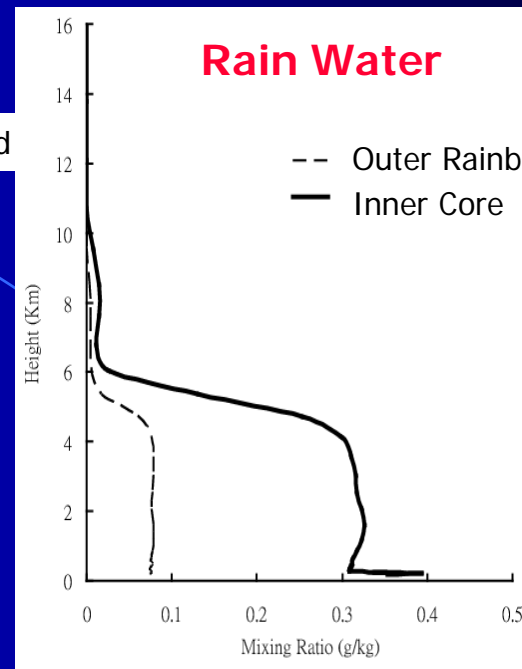
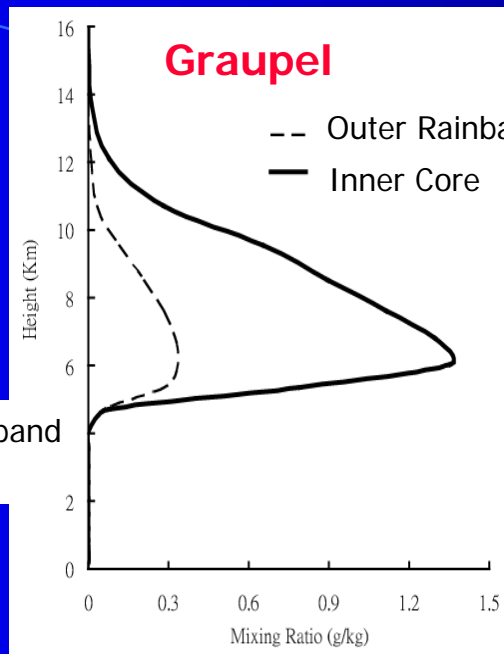
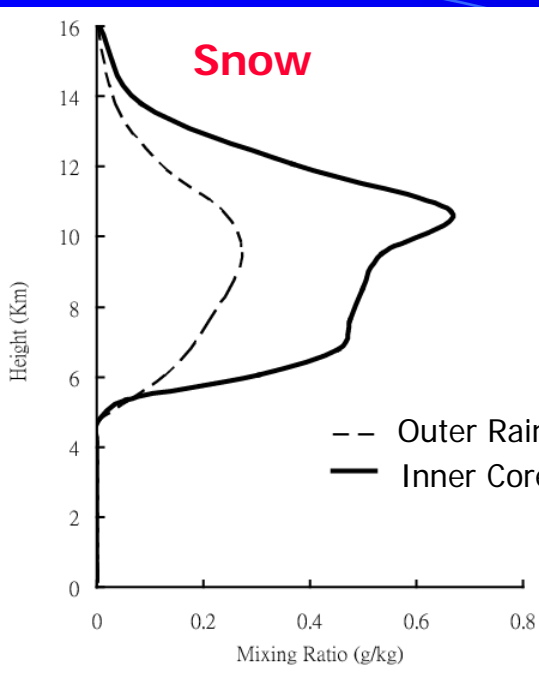




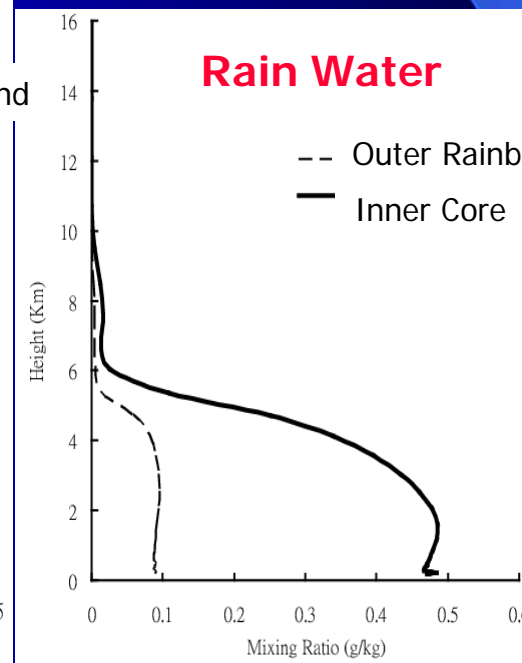
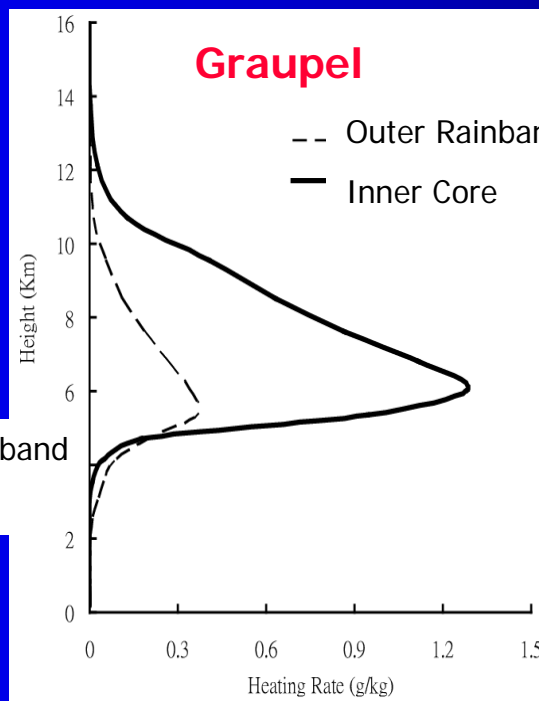
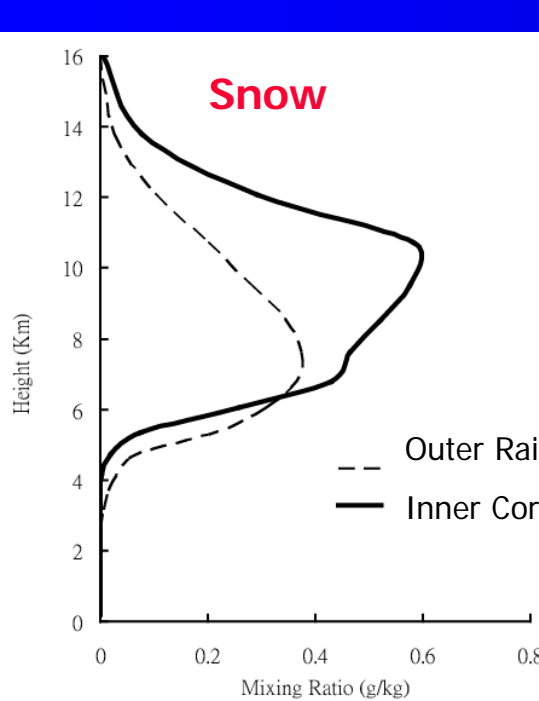
9 h Before
Landfall



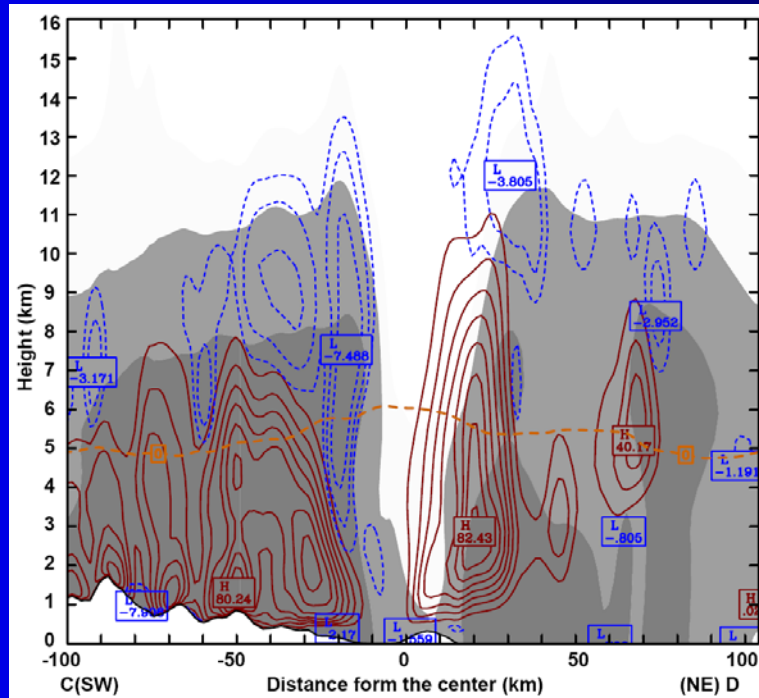
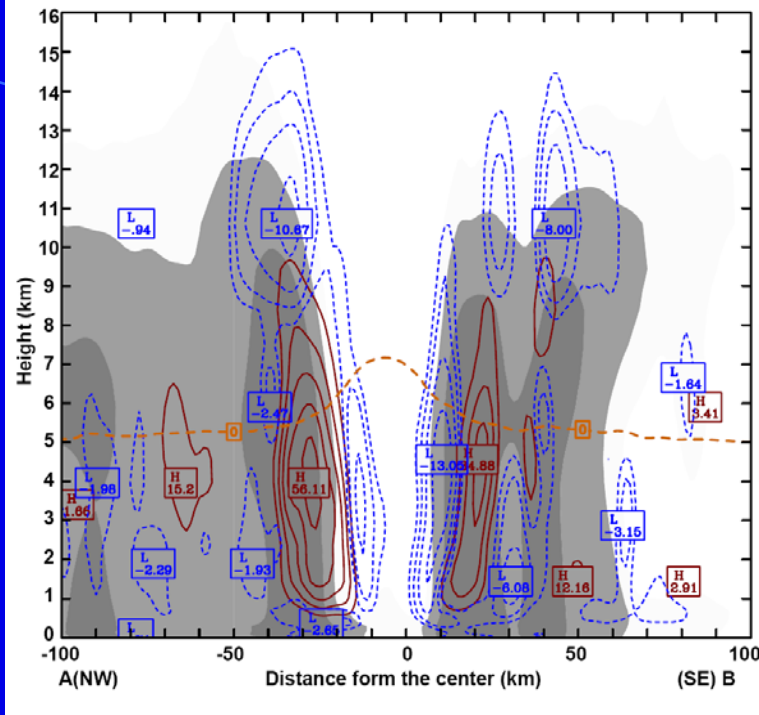
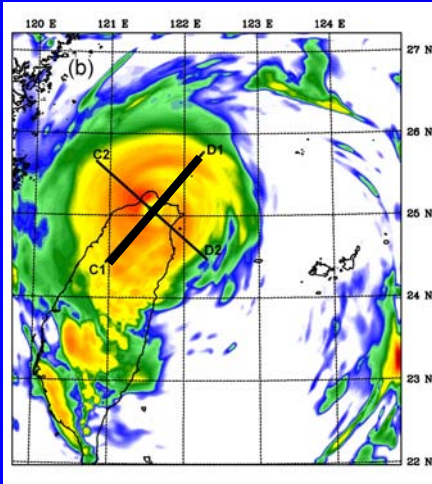
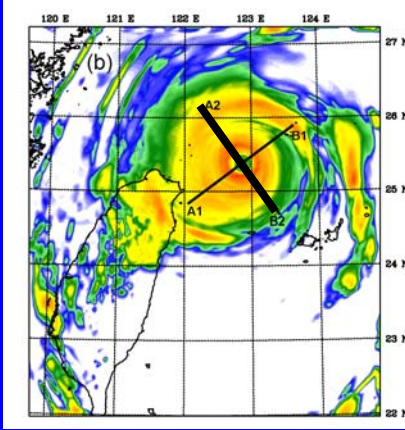
2 h After
Landfall



**Before
Landfall**



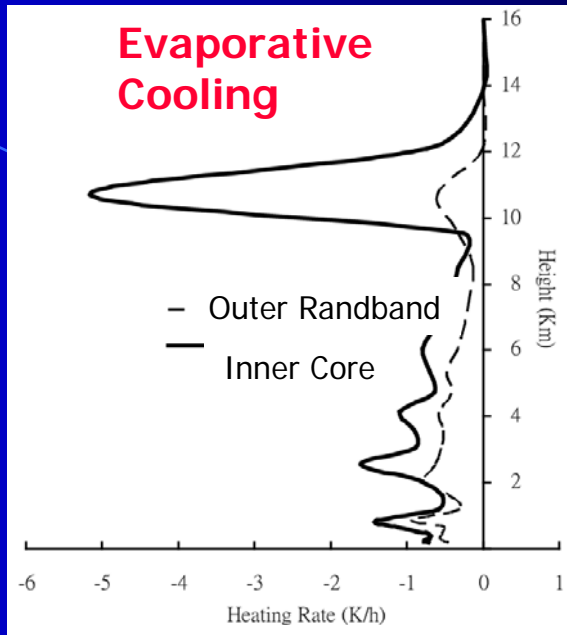
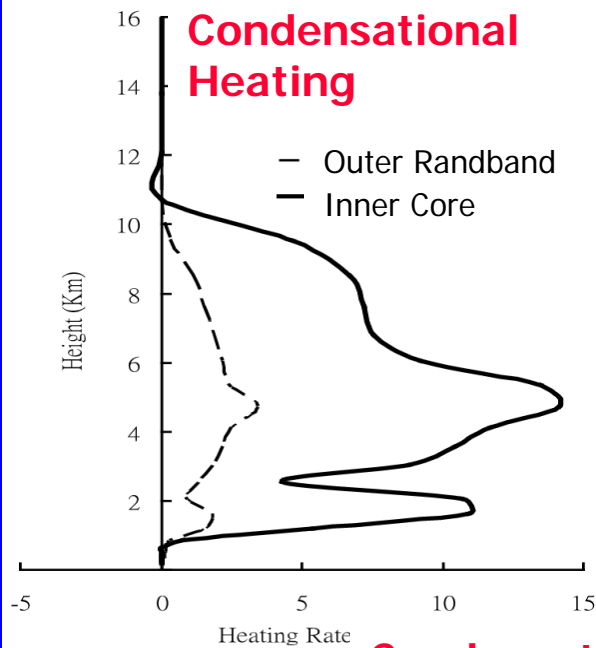
**After
Landfall**



Before Landfall

Condensation Heating
(solid line)
Evaporation Cooling
(dashed line)

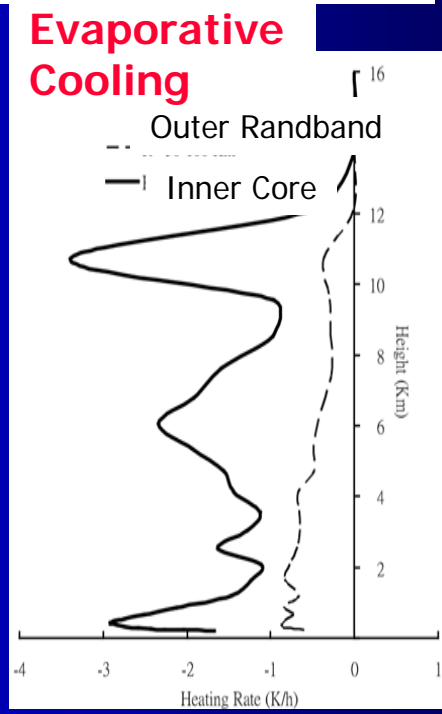
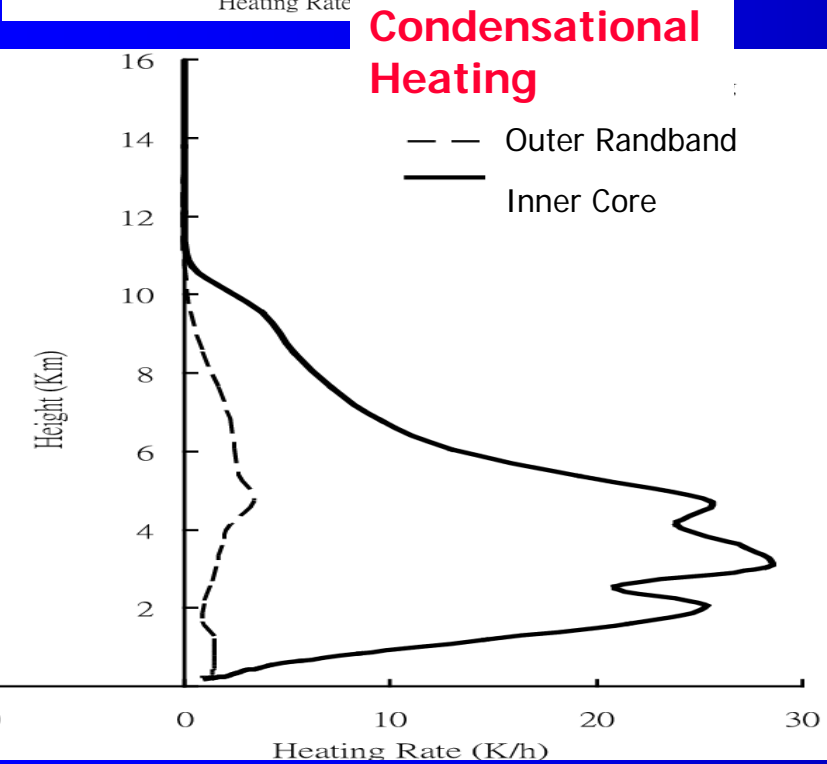
After Landfall



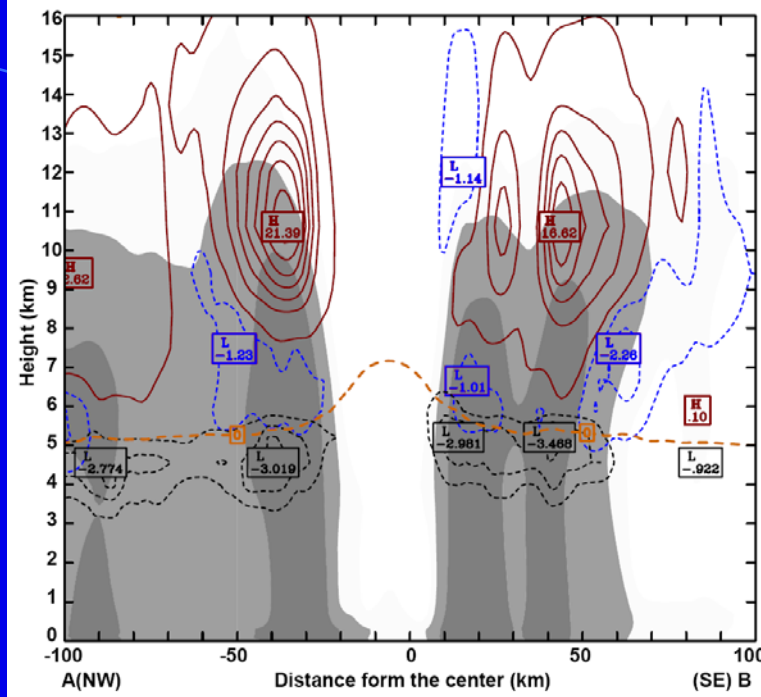
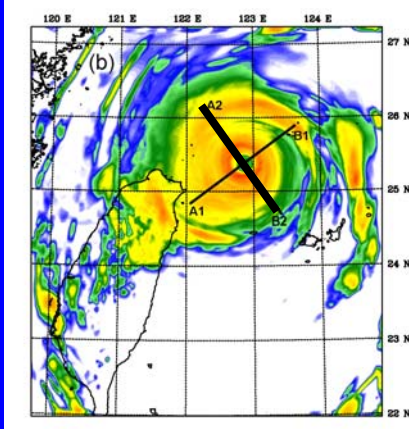
**Before
Landfall**

Condensation
Heating

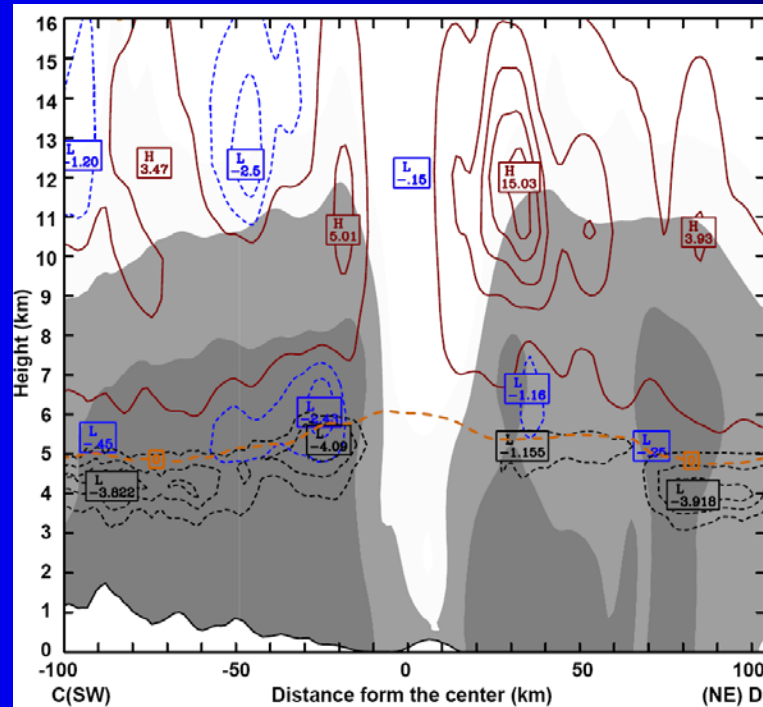
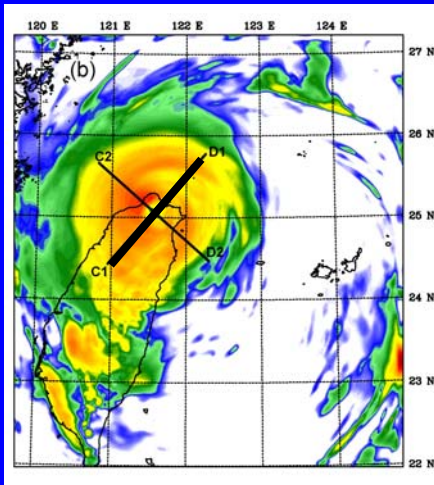
Evaporative
Cooling



**After
Landfall**

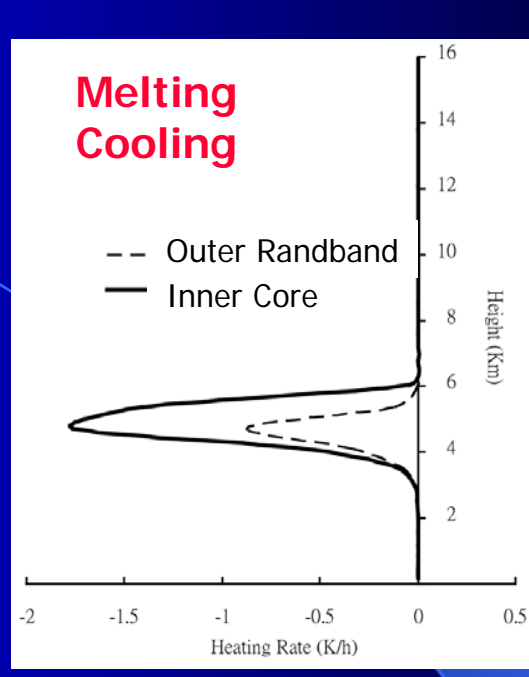
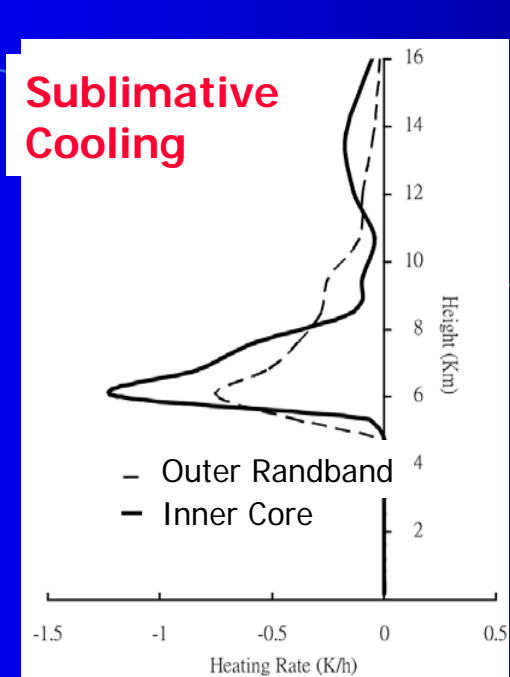
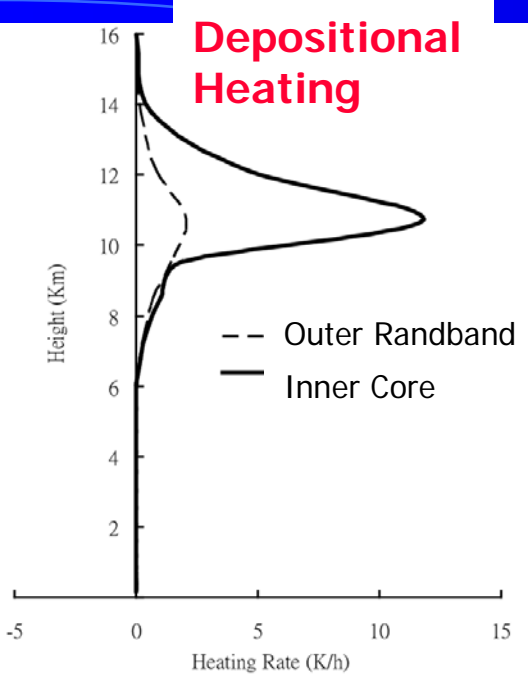


Before Landfall

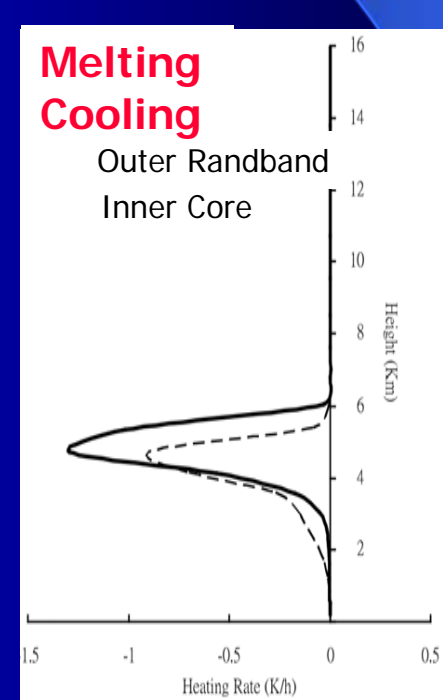
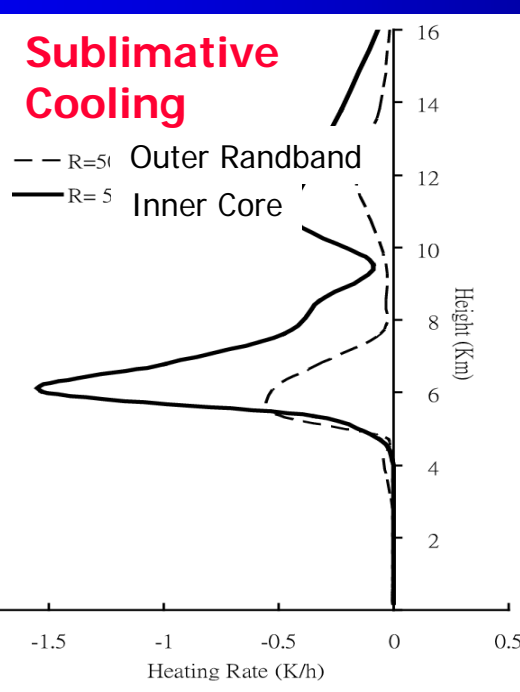
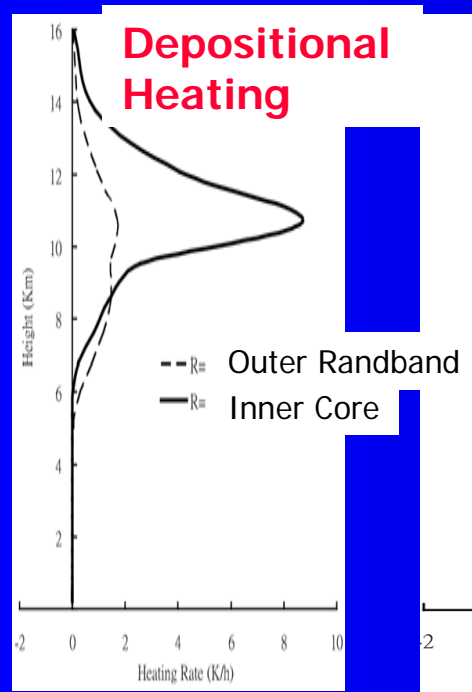


Deposition Heating
(solid black)
Sublimative Cooling
(dashed blue)
Melting Cooling
(dashed black)

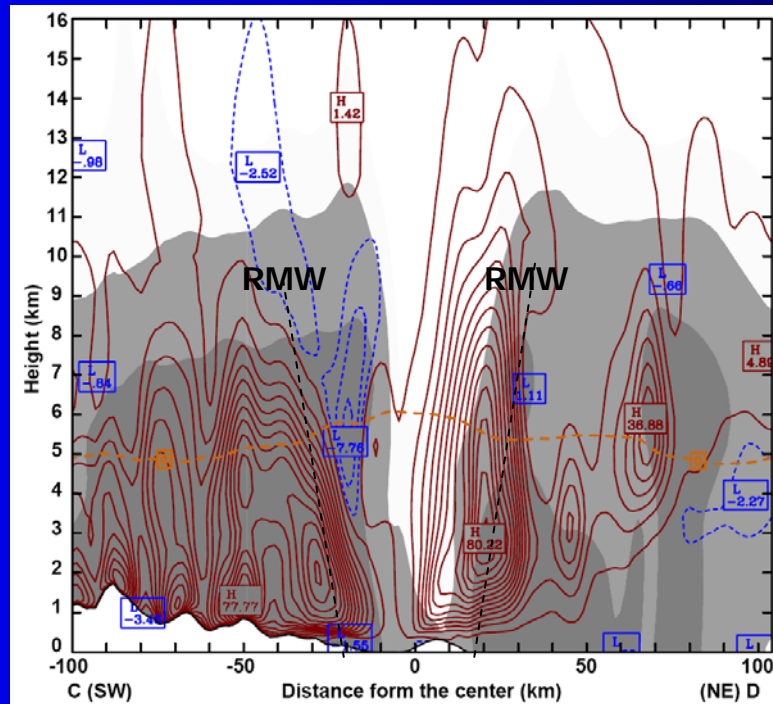
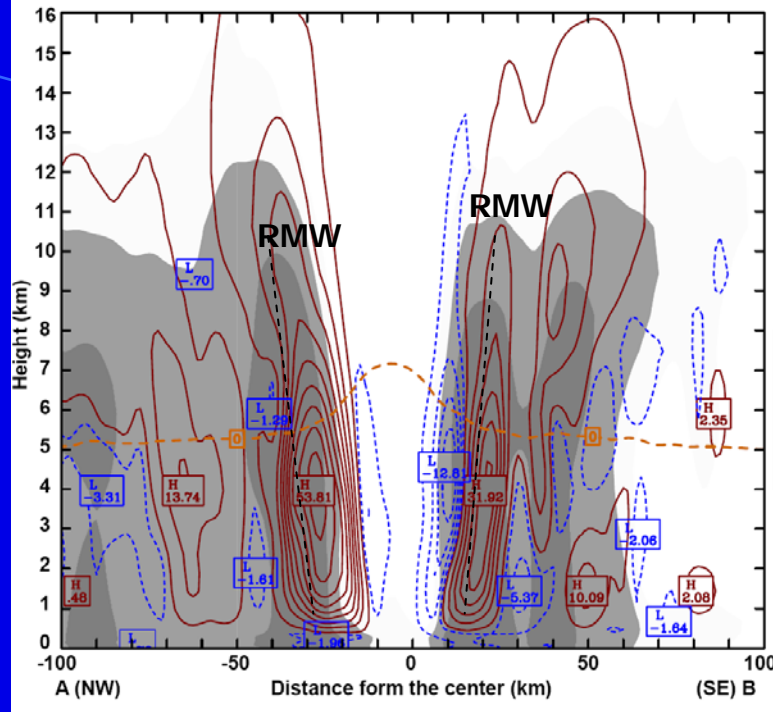
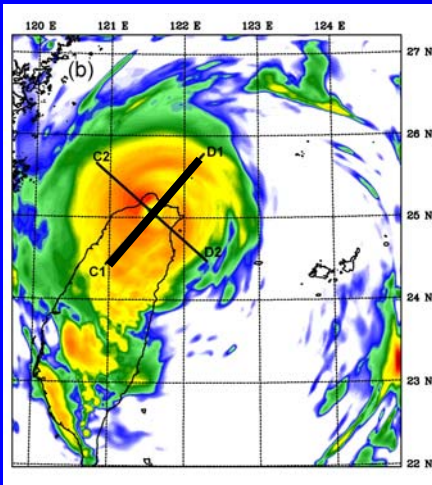
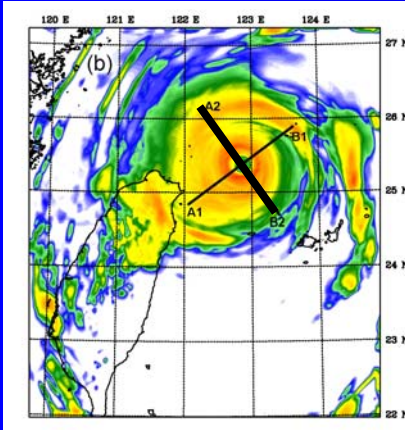
After Landfall



**Before
Landfall**



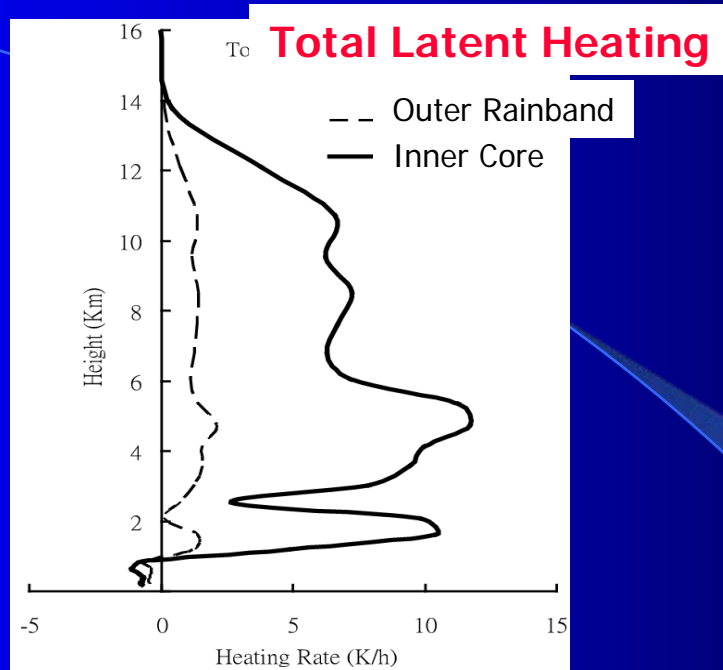
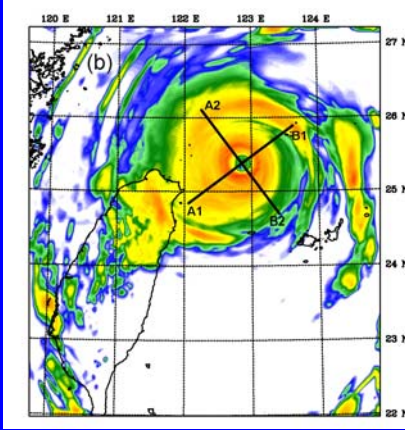
**After
Landfall**



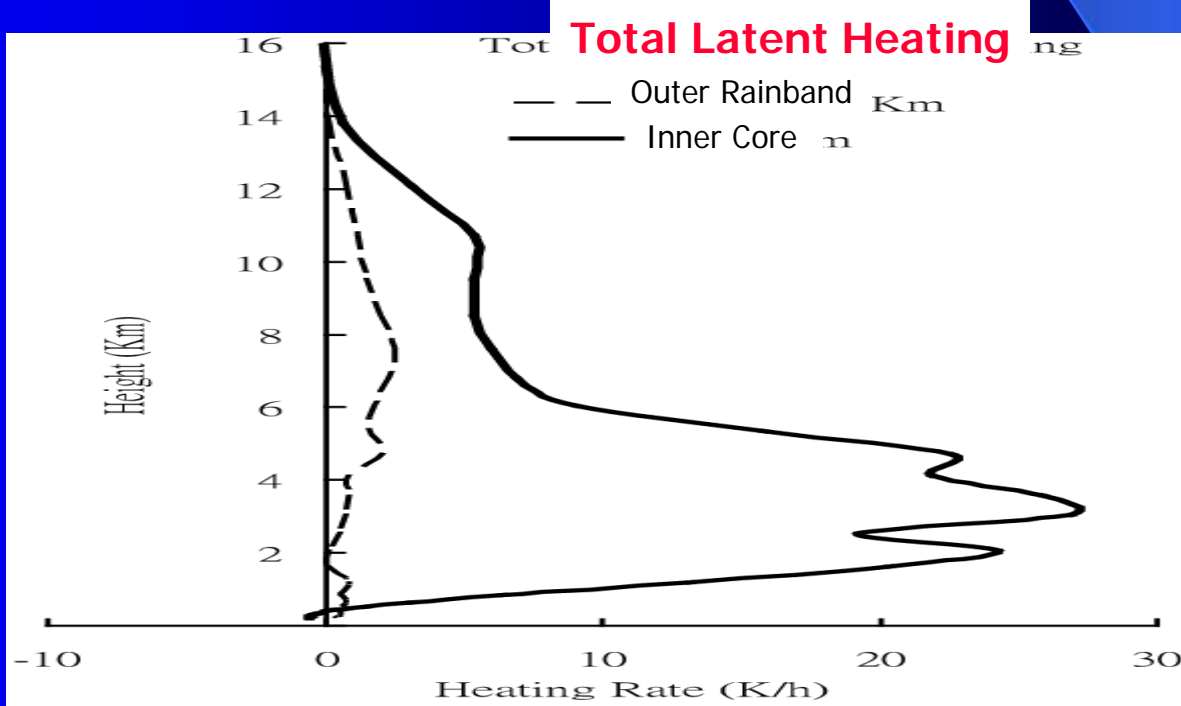
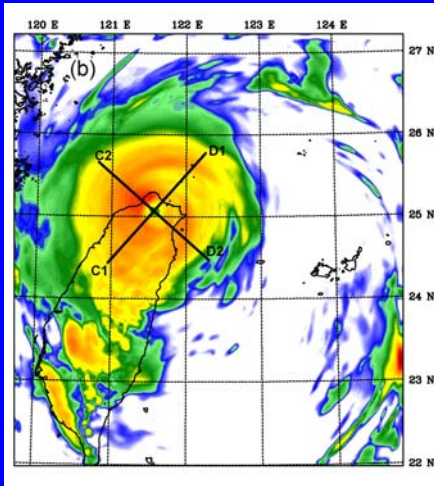
Before Landfall

Total Latent Heating
(solid black)
Total Latent Cooling
(dashed blue)

After Landfall

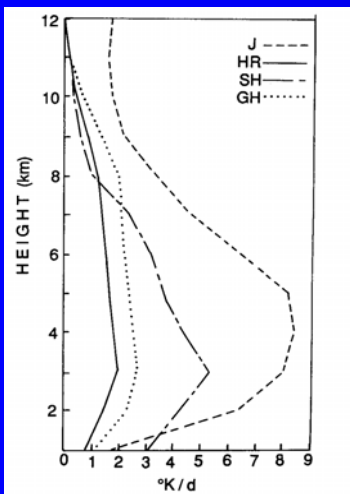


**Before
Landfall**

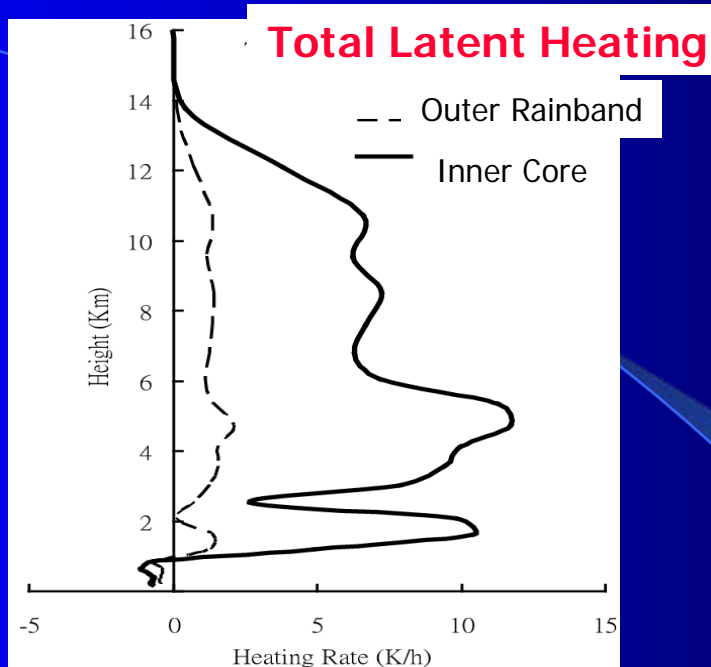


**After
Landfall**

Midlatitude MCS/Cv



Houze (1989; QJRMS)



Before Landfall

Total Latent Heating & Cooling

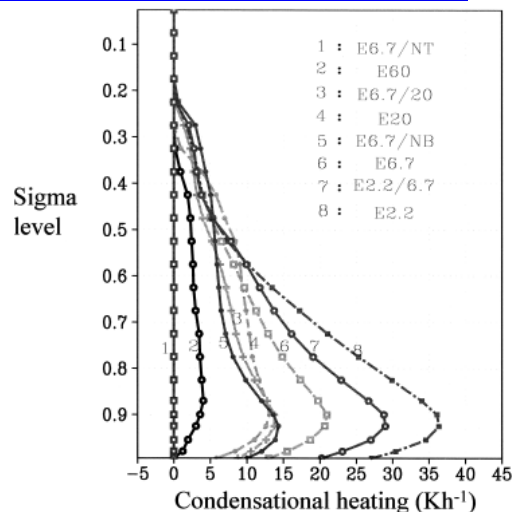
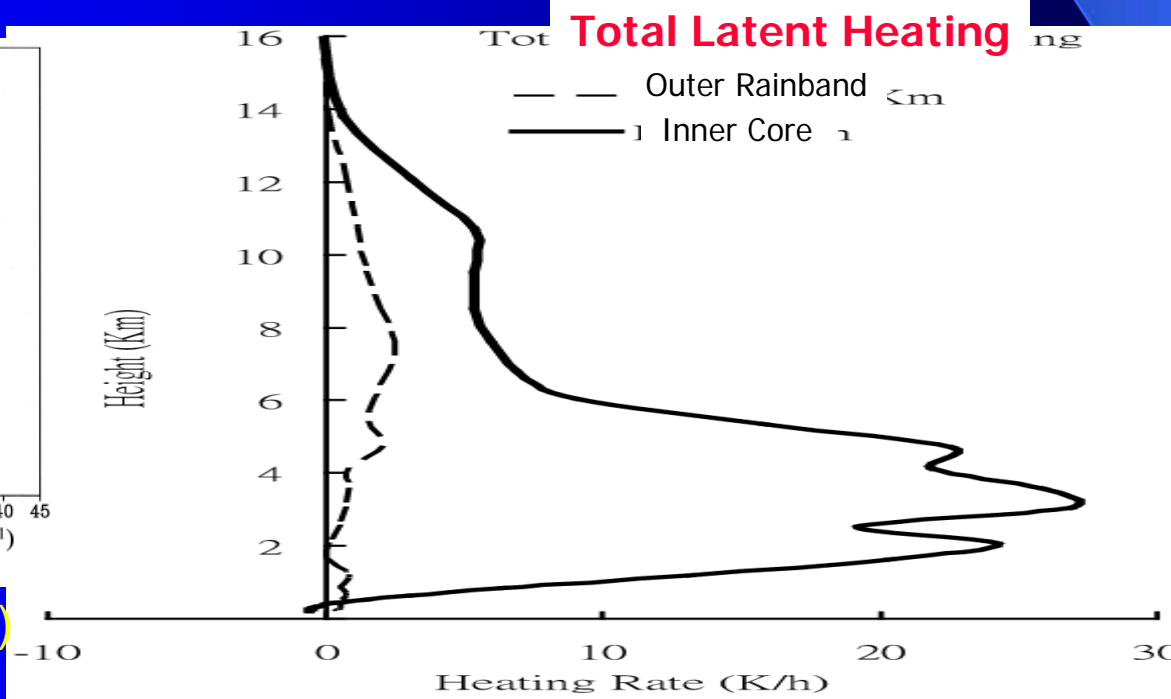


FIG. 12. Same as Fig. 11 but for condensation heating ($K h^{-1}$).

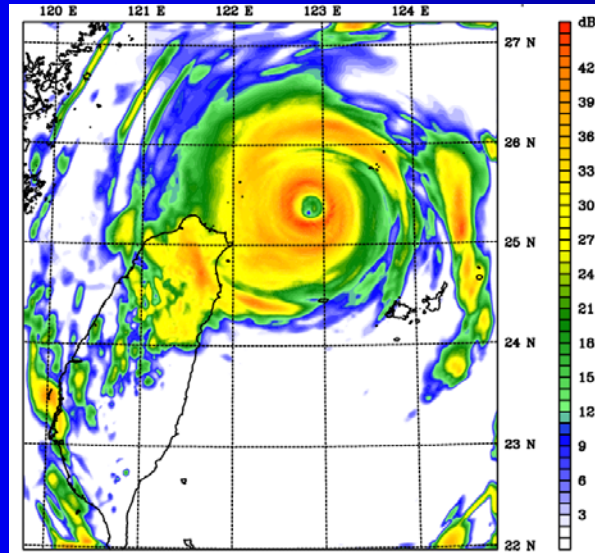
Wu et al. (2002; WAF)



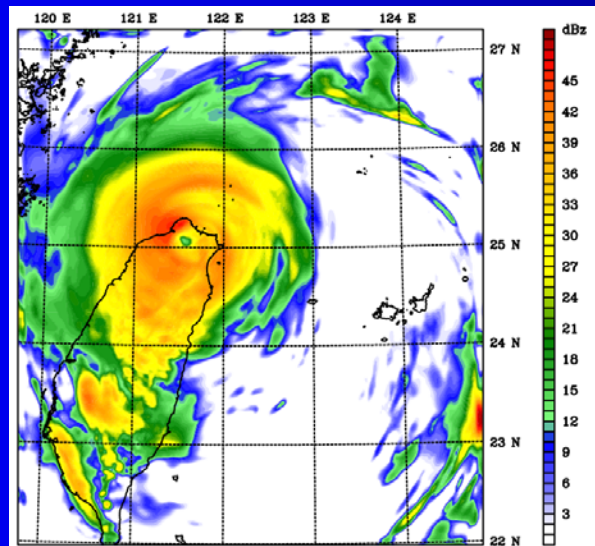
After Landfall

Storm Contraction during Landfall

Shrinkage of Typhoon Eye



**MM5 Radar CV @ 9/16 0130Z
(1-h time averaged)**

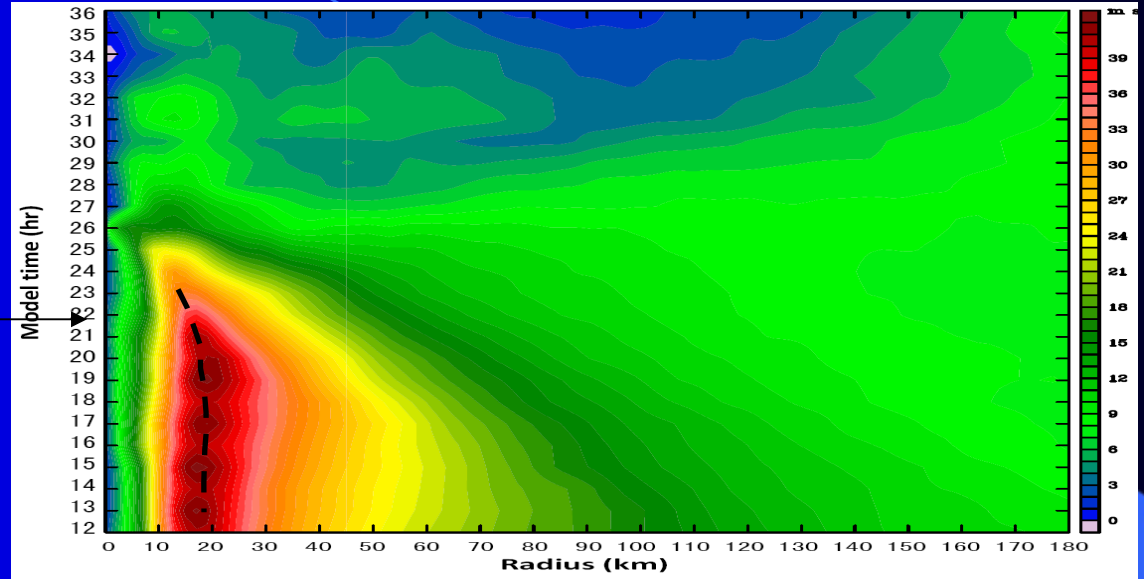


**MM5 Radar CV @ 9/16 1200Z
(1-h time averaged)**

Hovmoller Diagram of Azimuthal-Avg. Wind

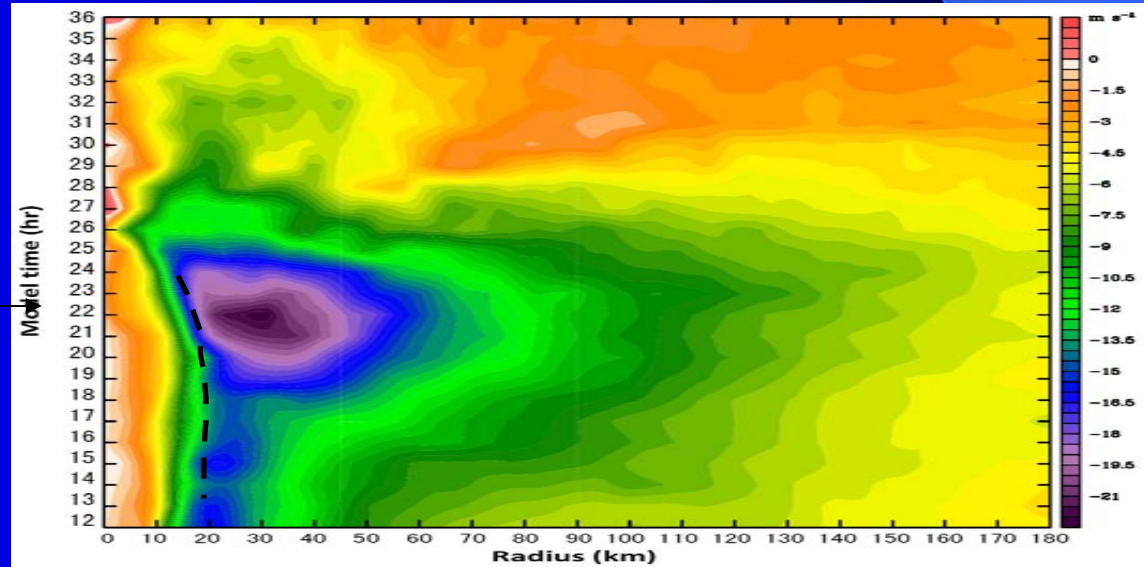
Near-Surface
Tangential Wind (V')

Landfall



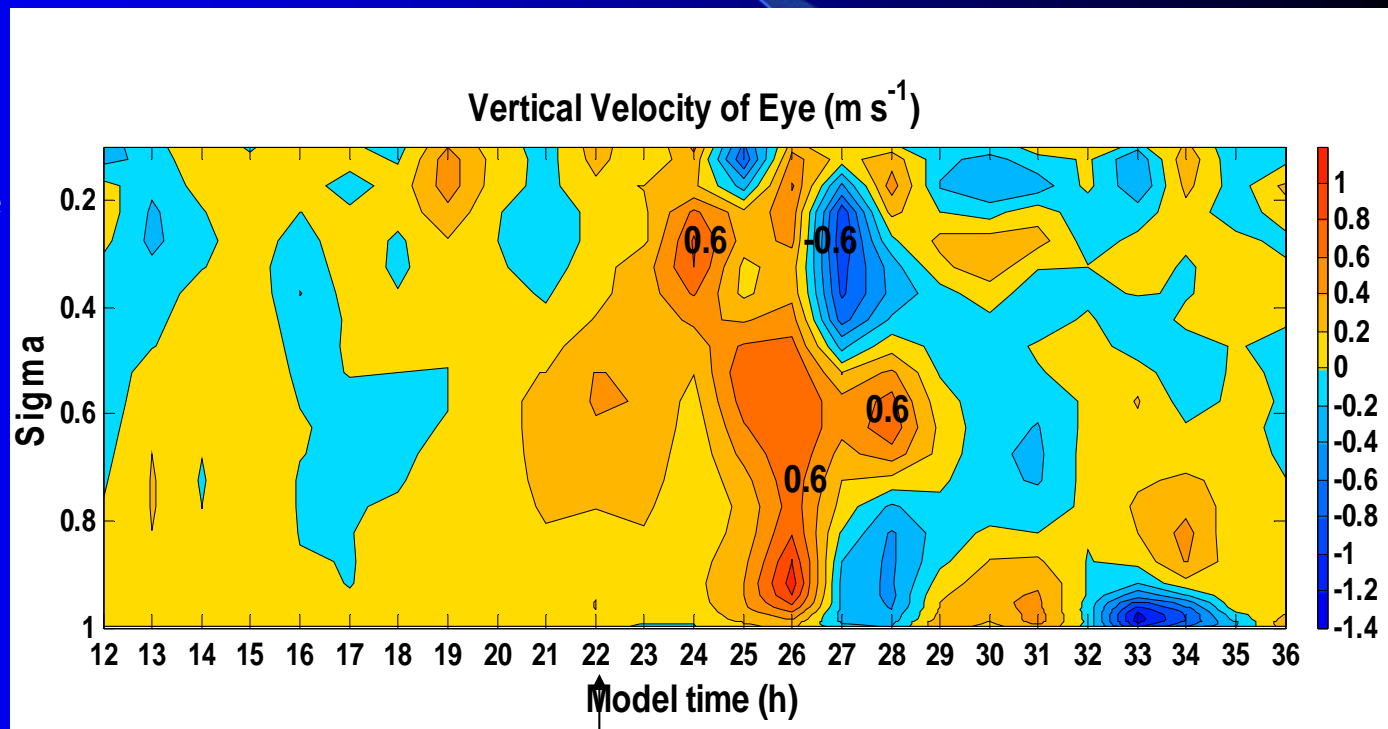
Near-Surface
Radial Wind (U')

Landfall



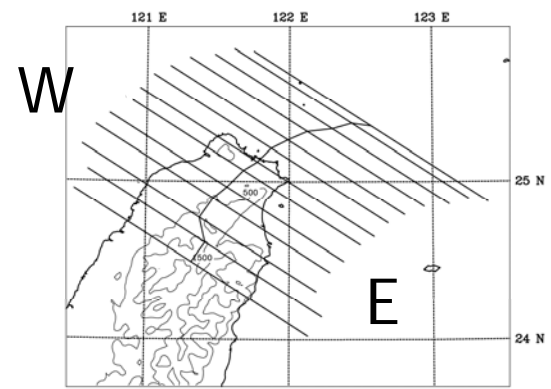
Time Series of Vertical Profiles of Vertical Velocity within the Eye

Averaged over a Square
of 12 km x 12 km
centered on the Eye

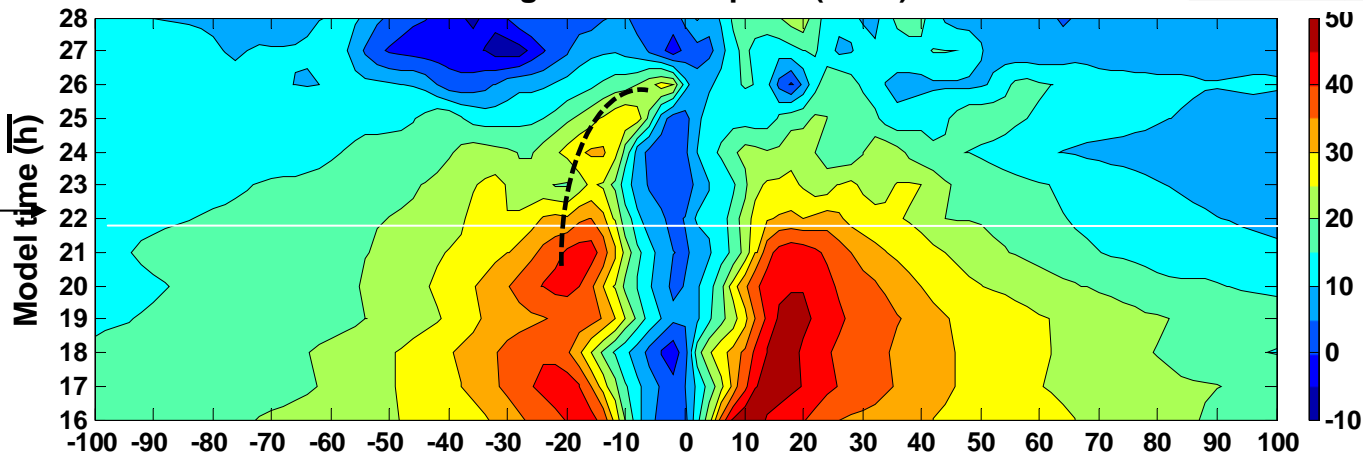


Landfall

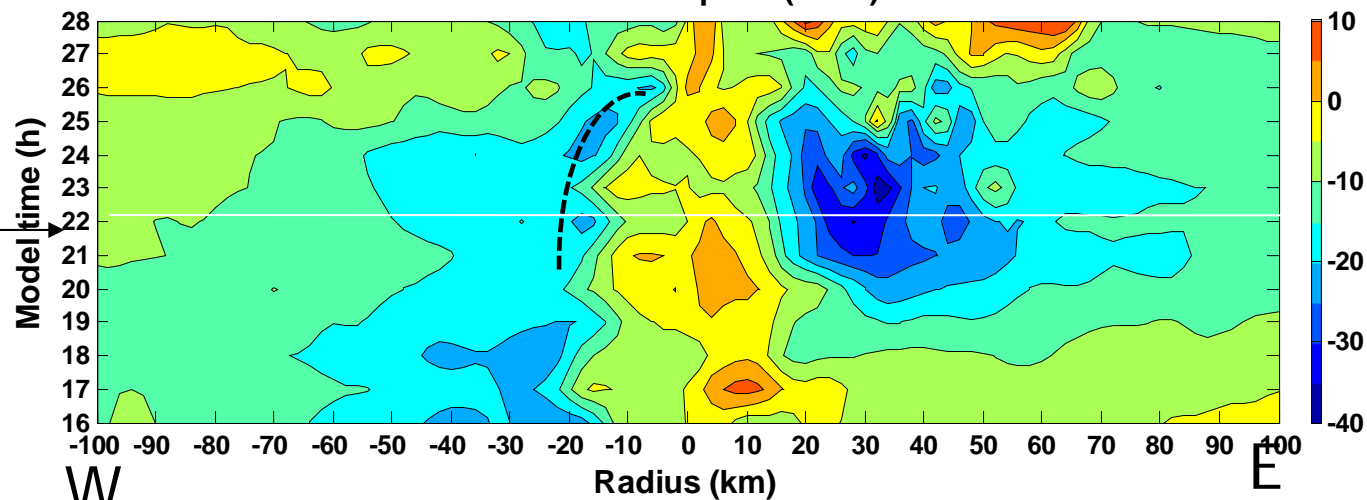
Hovmoller Diagram of Horizontal Wind



Tangential wind speed (m s^{-1})

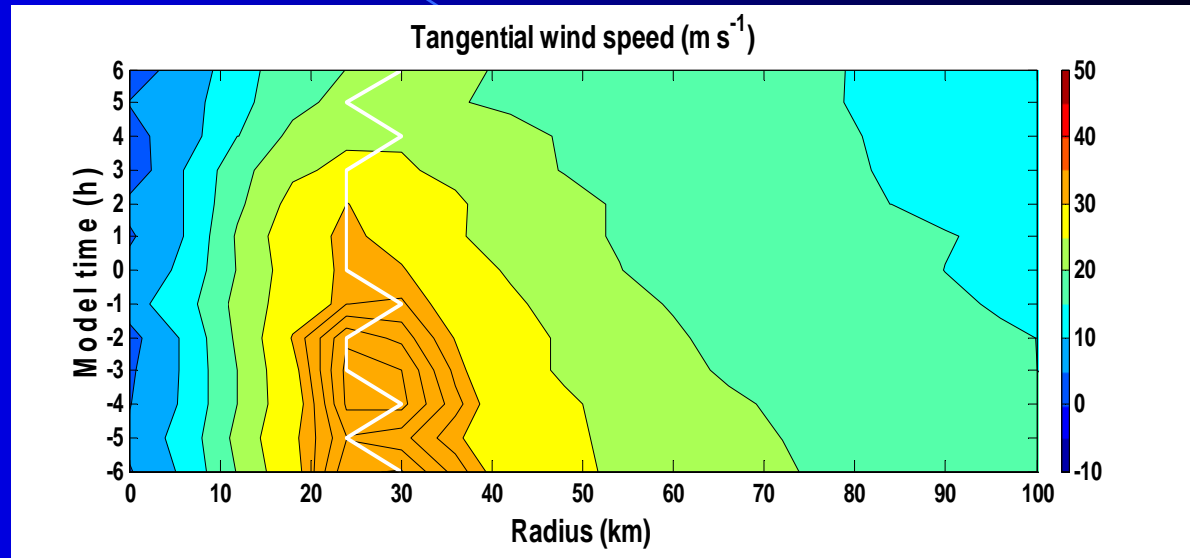


Radial wind speed (m s^{-1})

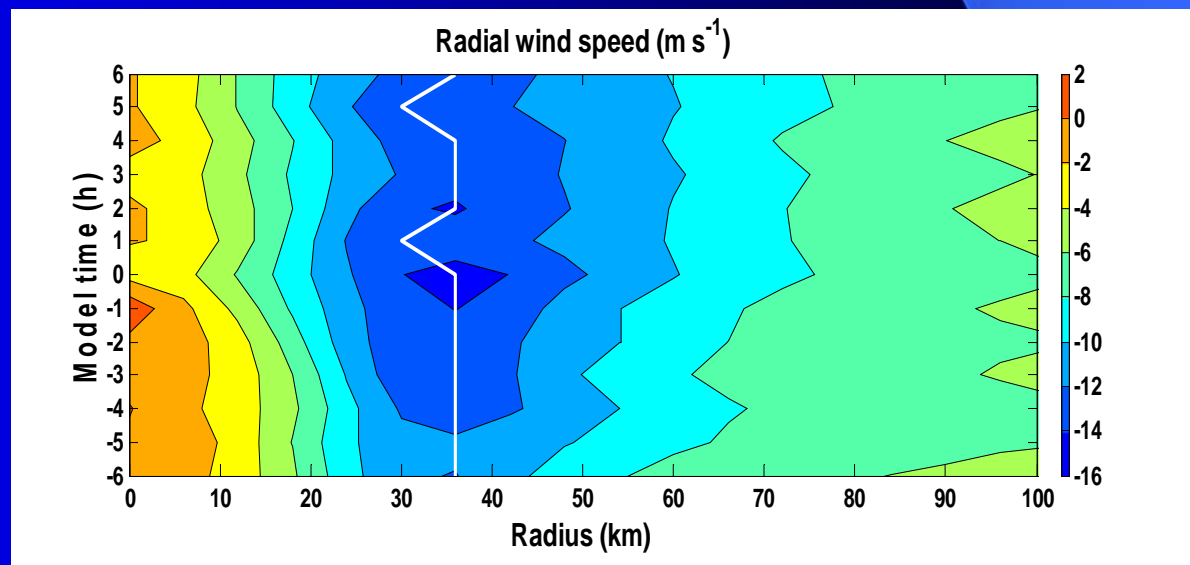


Hovmoller Diagram of Azimuthal-Avg. Wind (No Terrain Experiment)

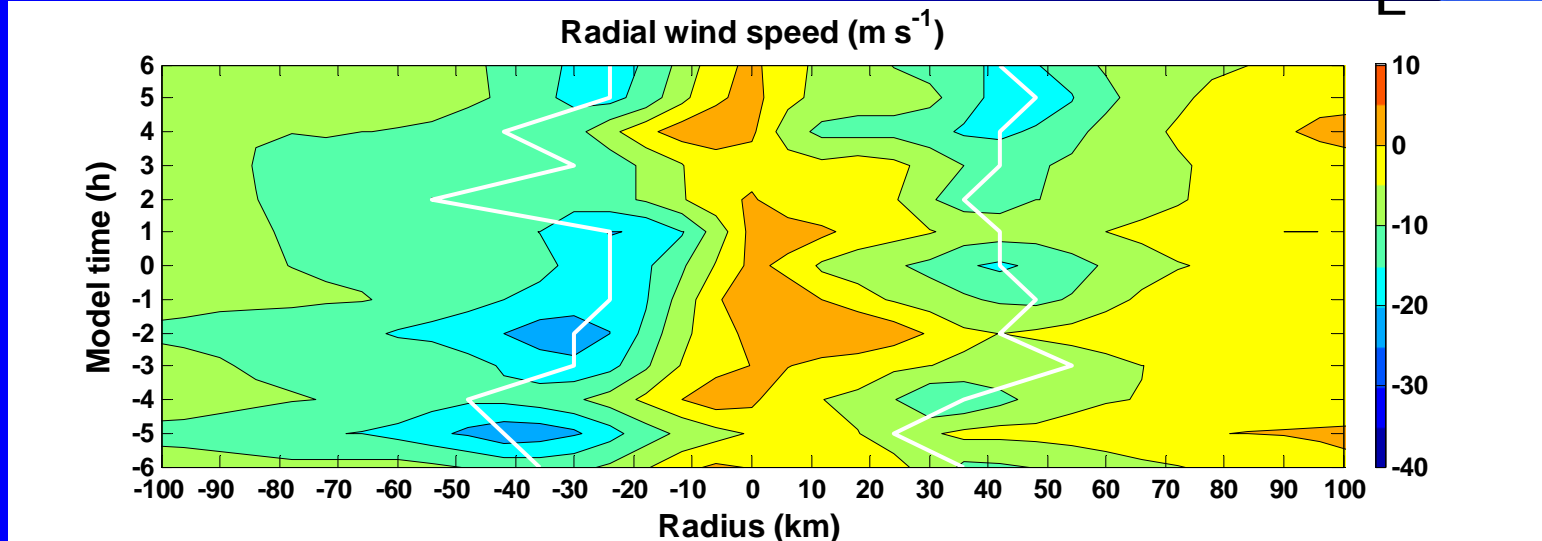
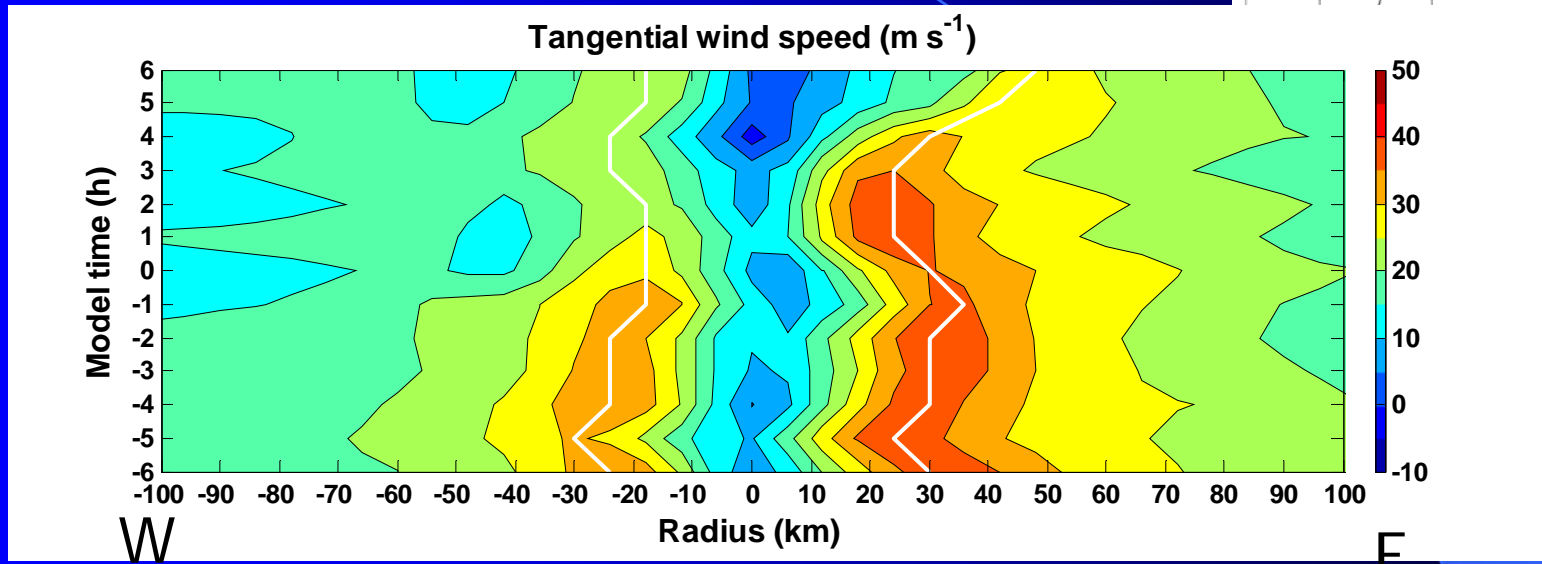
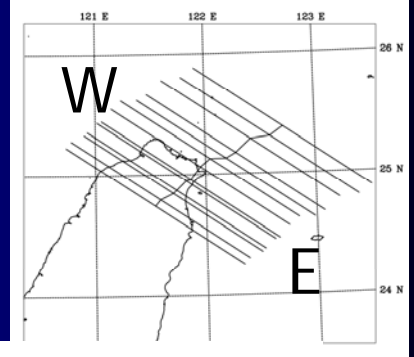
Near-Surface
Tangential Wind (V')



Near-Surface
Radial Wind (U')



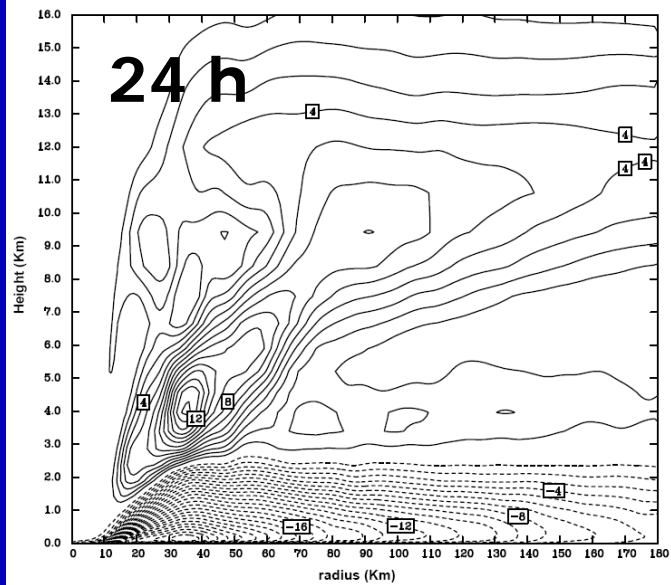
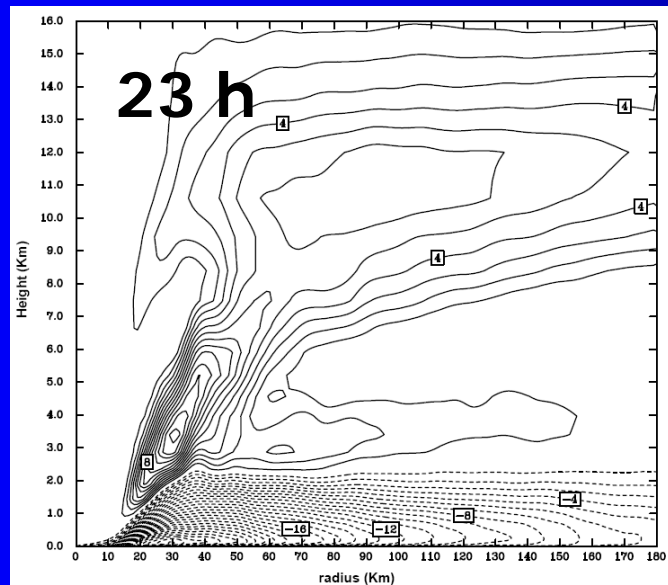
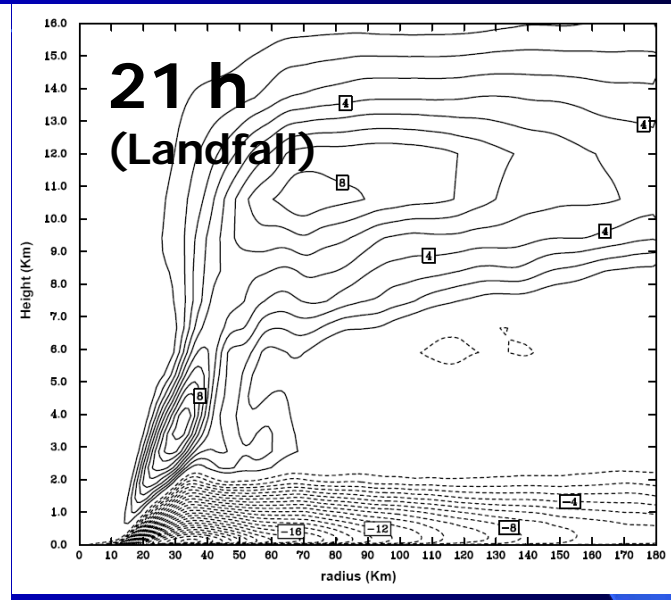
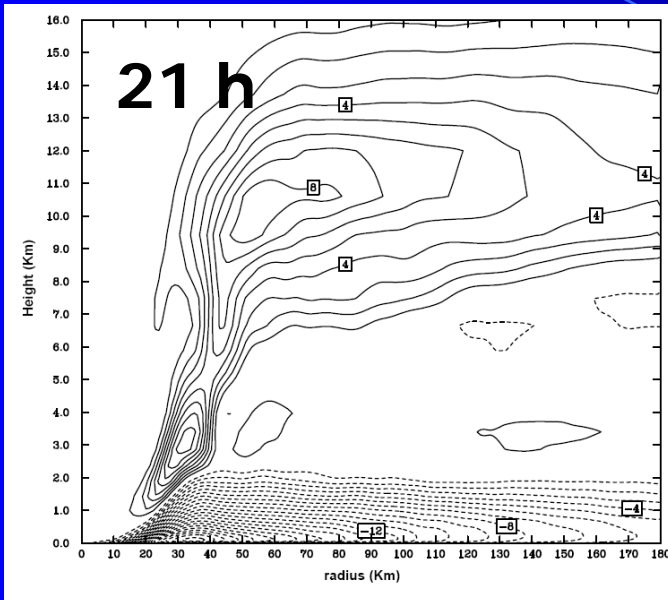
Hovmoller Diagram of Horizontal Wind (No Terrain Experiment)



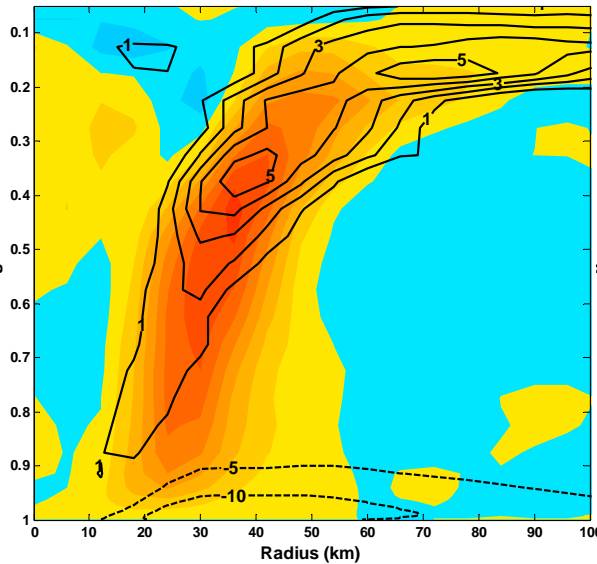
Evolution of Radial Flows

The background is a solid blue color with a subtle gradient. A thin, light blue curved line starts from the top left and arcs towards the right. On the right side, there is a large, light blue triangular shape that points towards the center, partially overlapping the curved line.

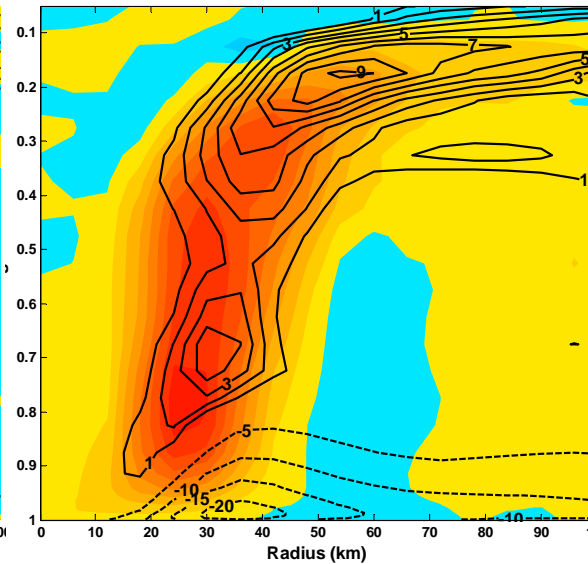
Azimuthal-Avg. Radial Wind ($r=180$ km)



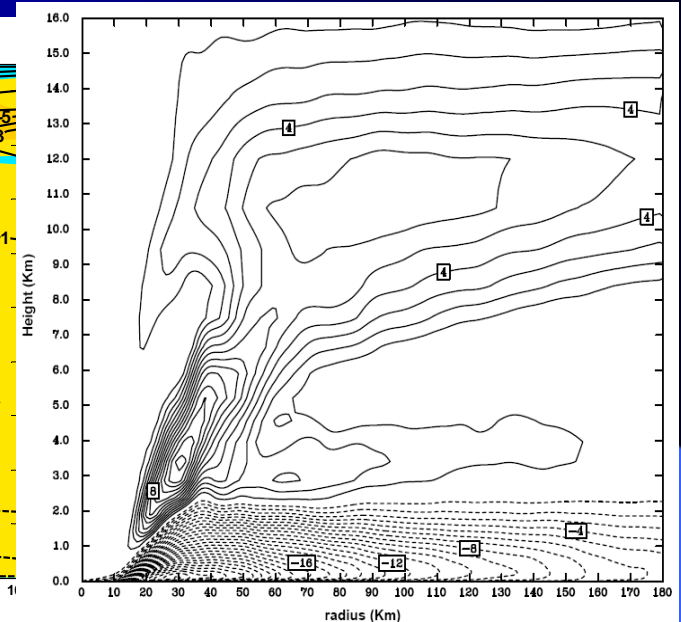
Azimuthal-Avg. Radial Wind (1 h after landfall)



No Terrain
(c.t. = 5 m/s)



50% Terrain
(c.t. = 5 m/s)



Full Terrain
(c.t. = 1 m/s)

Radial Momentum Budget

- Following Zhang et al. (2001), the governing equation for the radial momentum can be written as

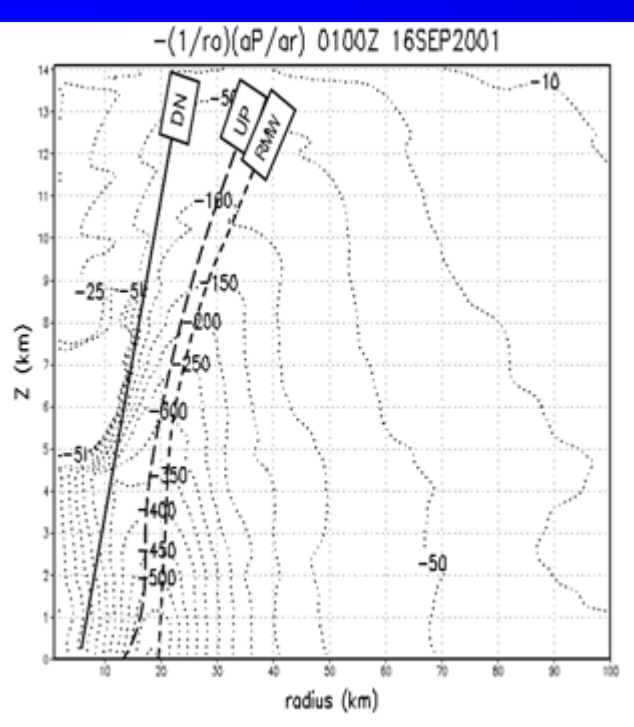
$$\frac{dU}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial r} + \frac{V^2}{r} + fV - 2\Omega \cos \phi W \cos \lambda + U_D$$

- where

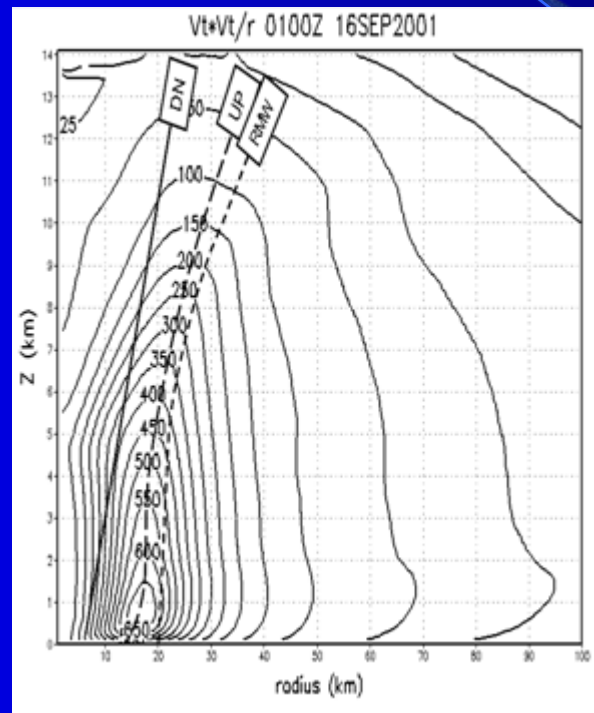
$$\frac{d}{dt} = \frac{\delta}{\delta t} + U' \frac{\partial}{\partial r} + \frac{V'}{r} \frac{\partial}{\partial \lambda} + W \frac{\partial}{\partial z}$$

- and W , U , and V are the vertical, radial and tangential winds; U' and V' are the horizontal wind components relative to the storm.

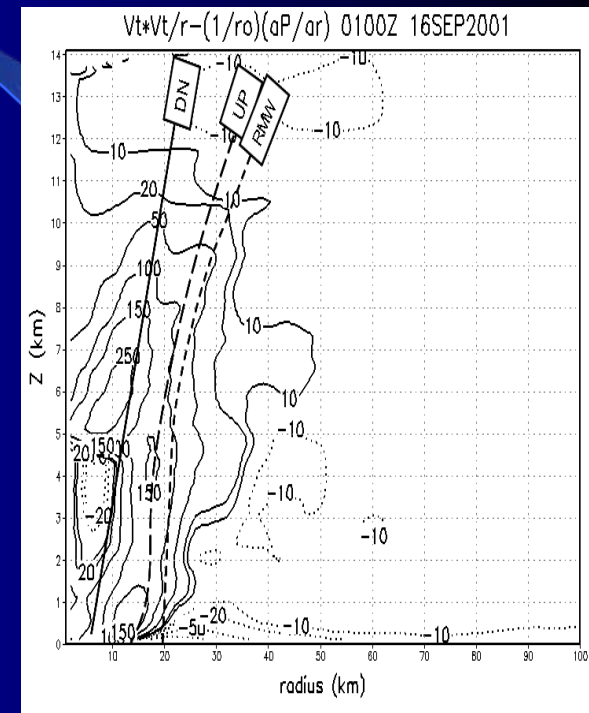
Axisymmetric Radial Momentum Budget 9 h before Landfall



$PGF_R (U_P)$



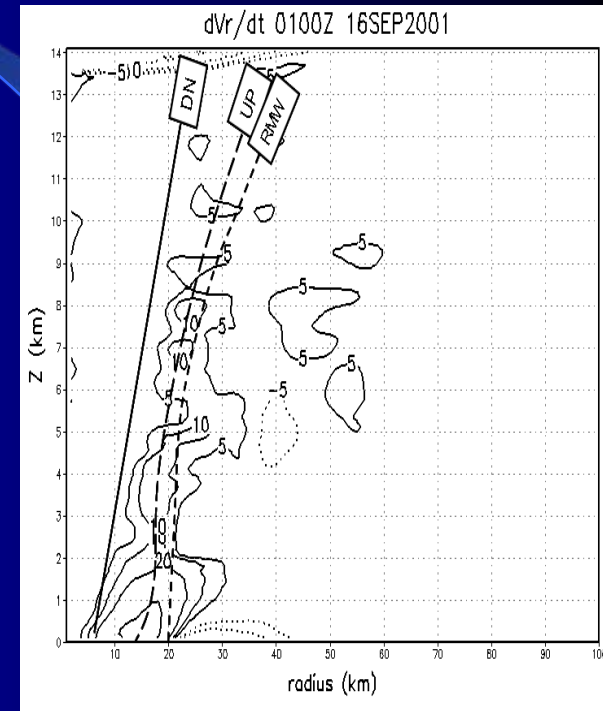
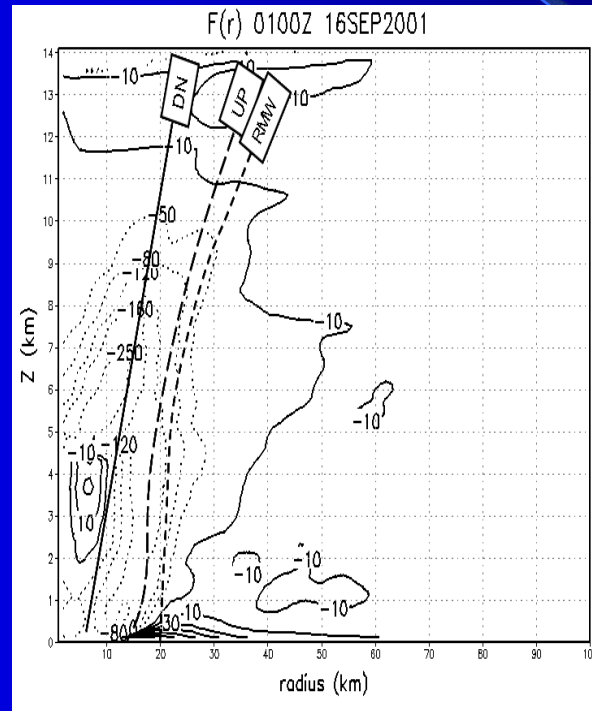
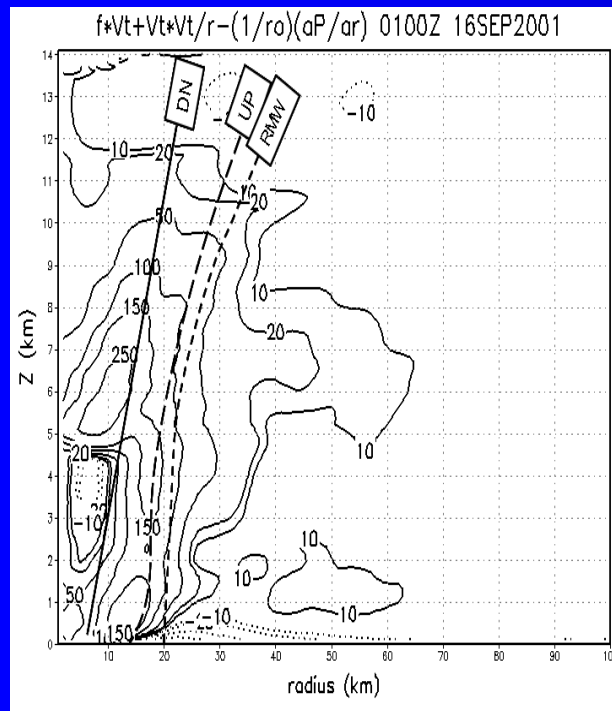
Centrifugal force (U_E)



Cyclostrophic force
imbalance ($U_P + U_E$)

Oceanic Stage @ 13-14 h

Axisymmetric Radial Momentum Budget 9 h before Landfall



Diffusion & MBL (U_B)

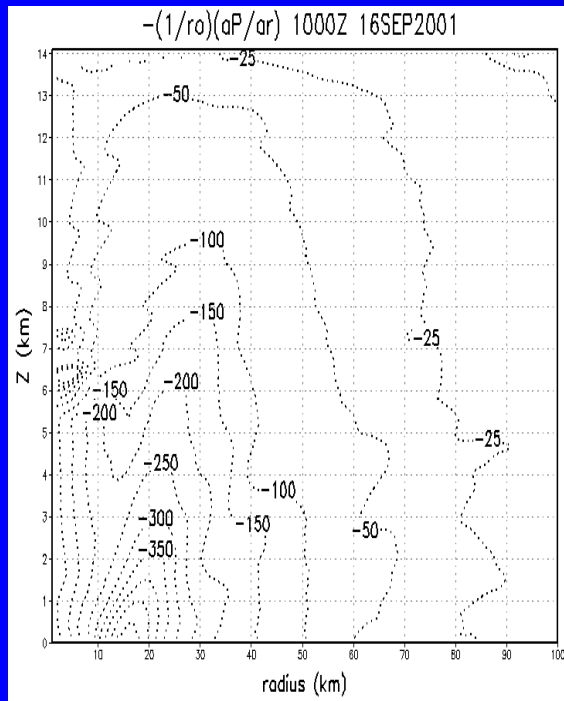
Gradient force imbalance

($U_P + U_E + U_C$)

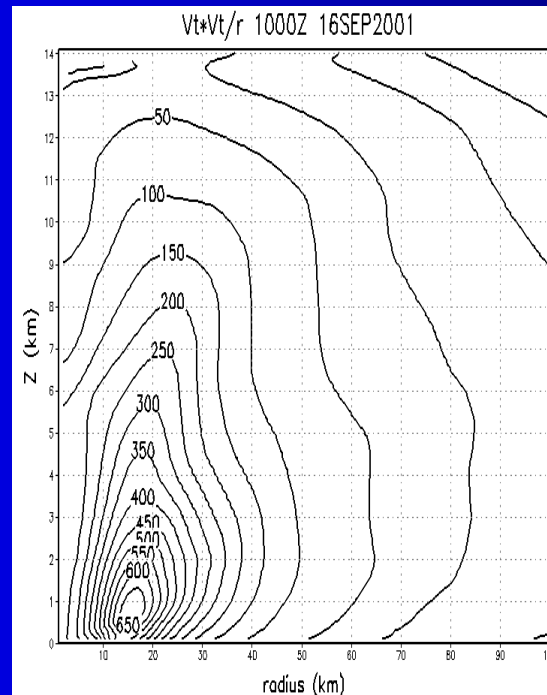
Net Lagrangian
tendency (dU/dt)

Oceanic Stage @ 13-14 h

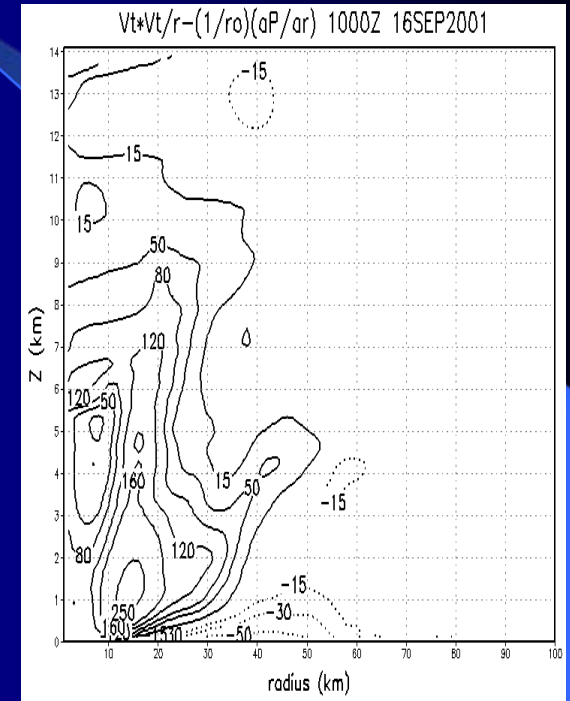
Axisymmetric Radial Momentum Budget 1 h after Landfall



$PGF_R (U_p)$



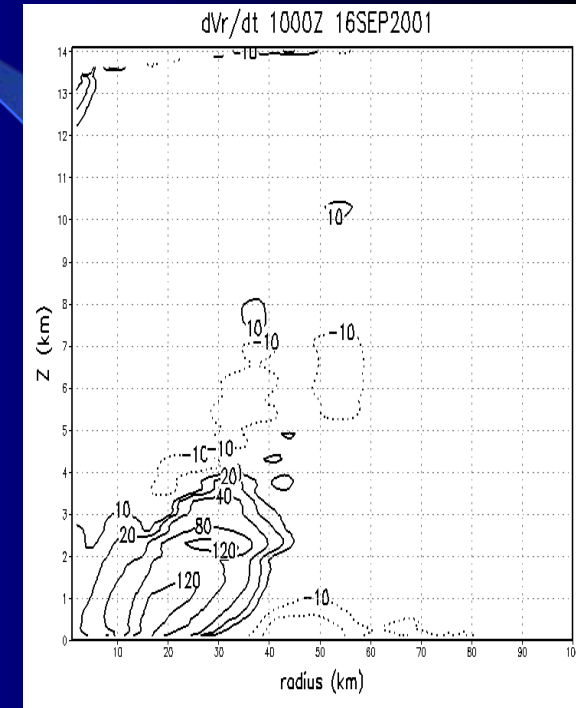
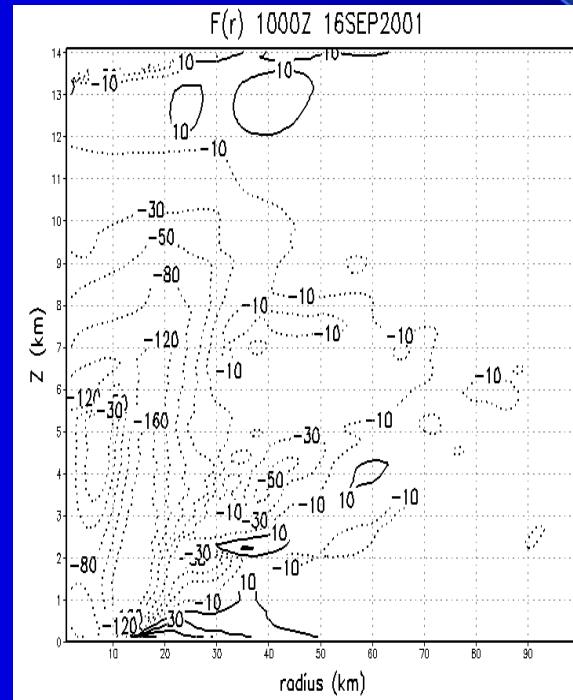
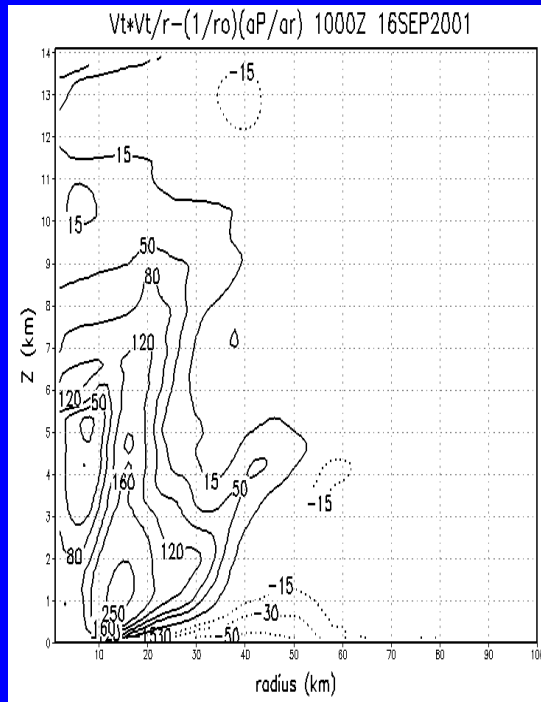
Centrifugal force (U_E)



Cyclotrophic force
imbalance ($U_p + U_E$)

Landfall Stage @ 22-23 h

Axisymmetric Radial Momentum Budget 1 h after Landfall



Diffusion & MBL (U_B)

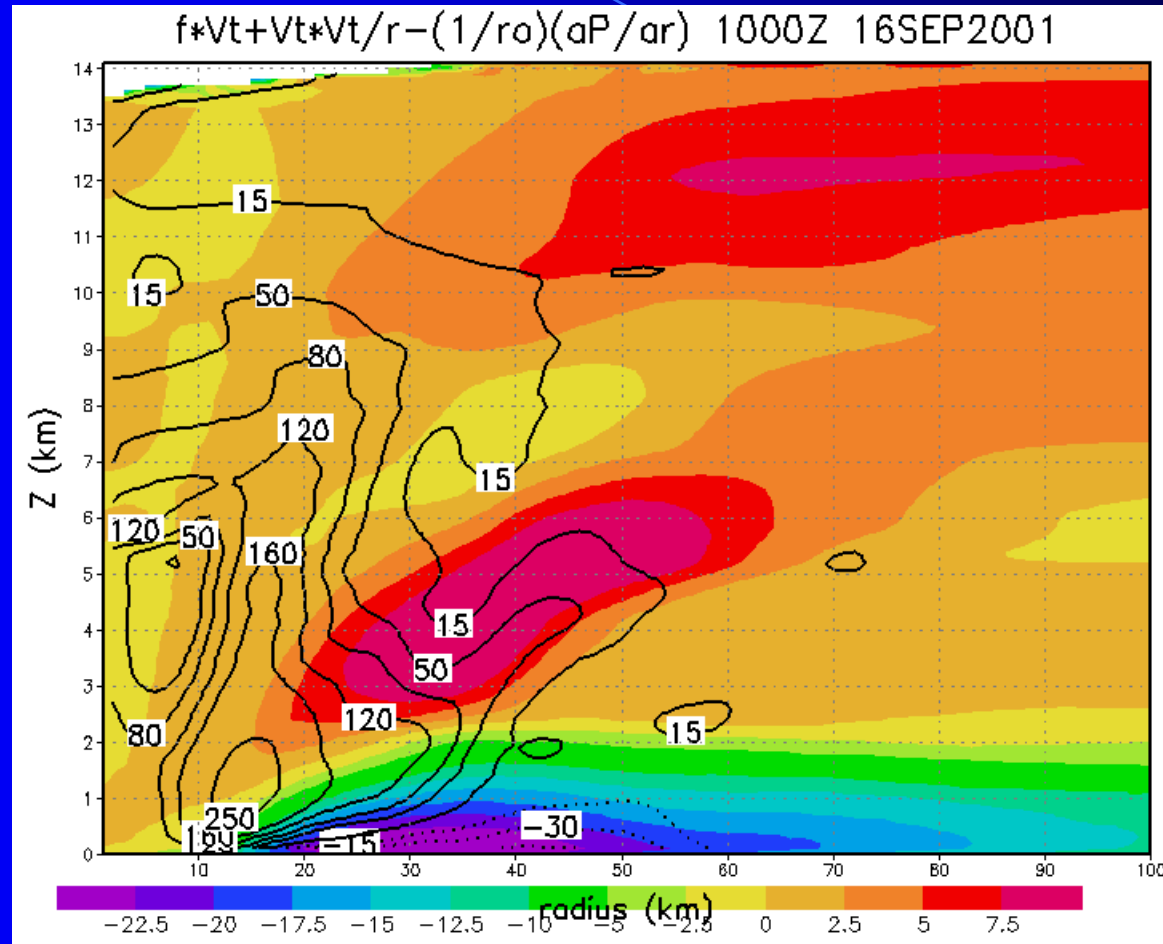
Net Lagrangian
tendency (dU/dt)

Gradient force imbalance

($U_P + U_E + U_C$)

Landfall Stage @ 22-23 h

Axisymmetric Radial Momentum Budget 1 h after Landfall

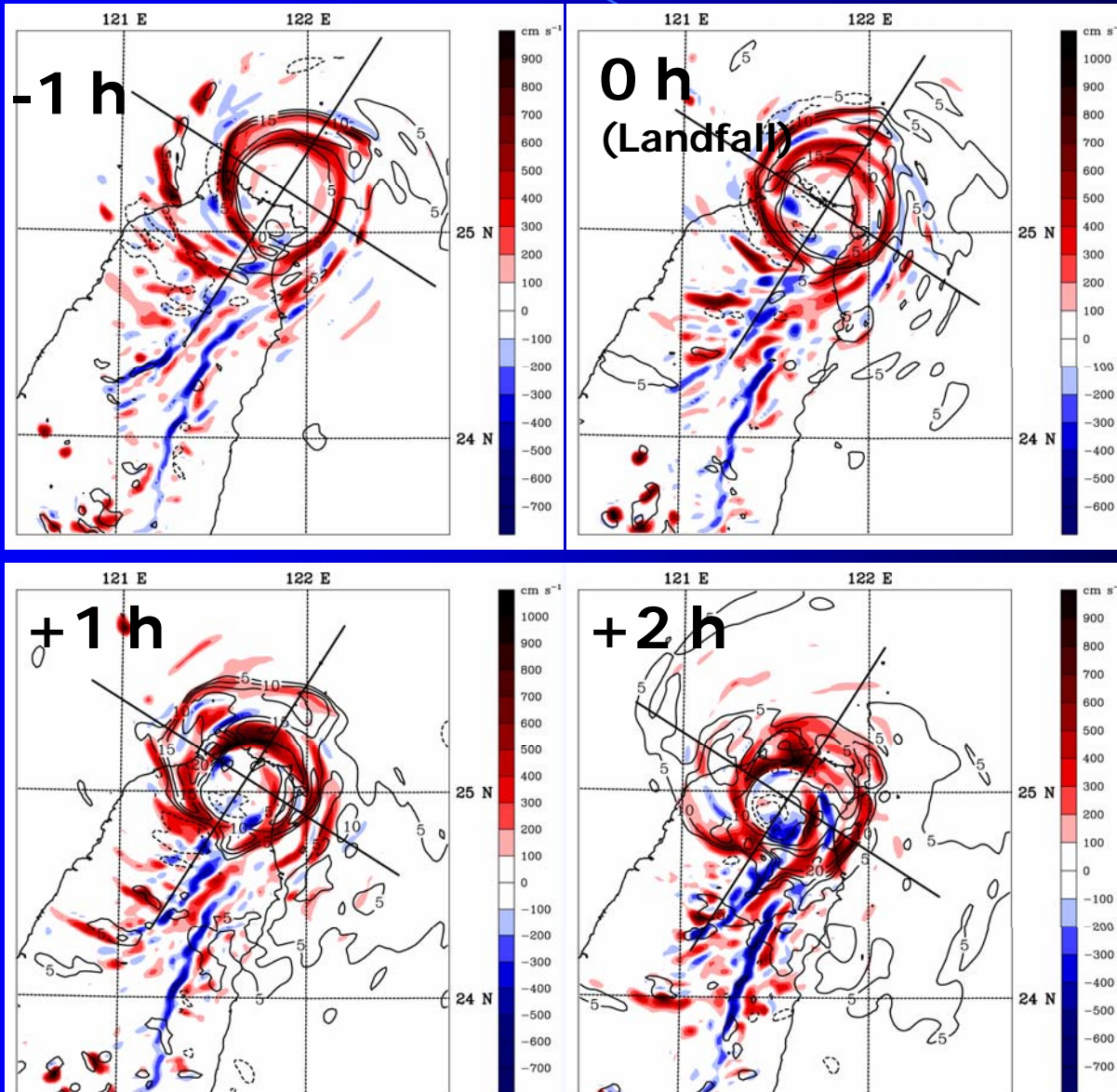


Contour: Gradient force
imbalance

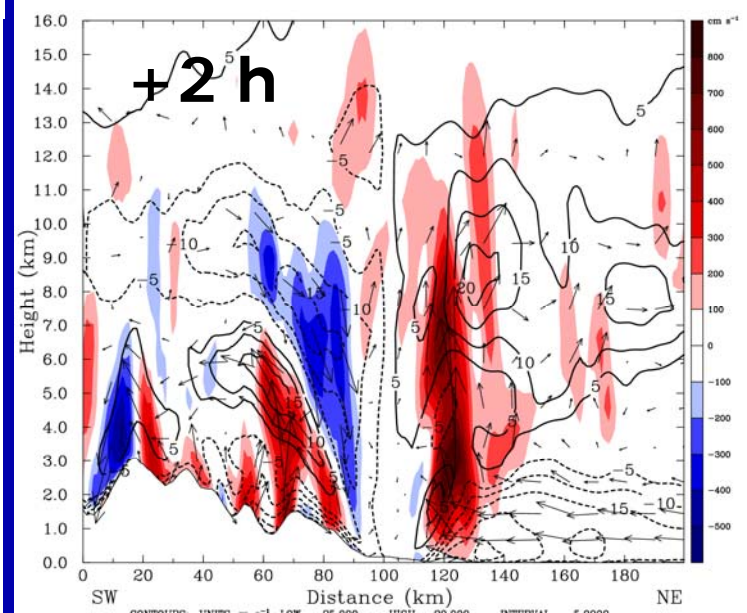
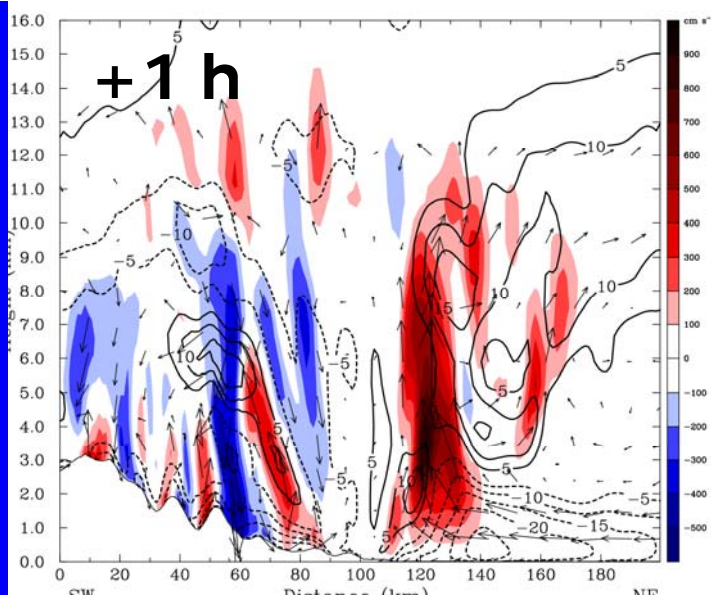
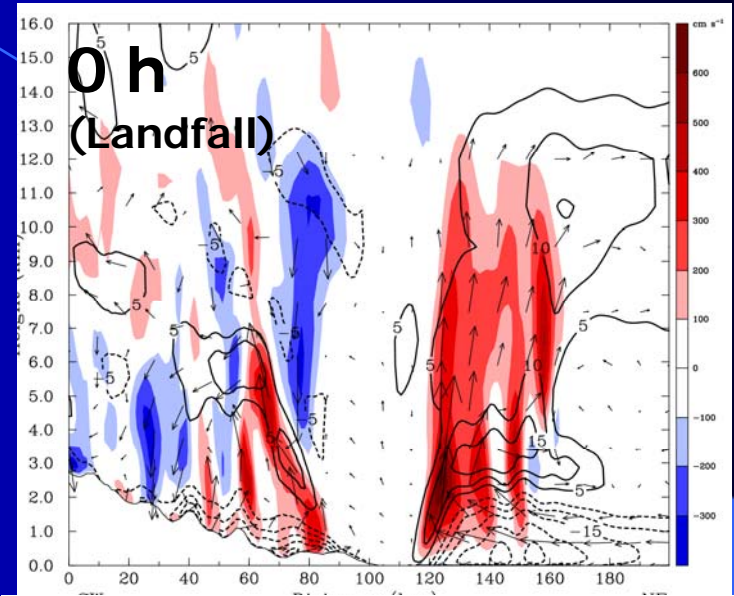
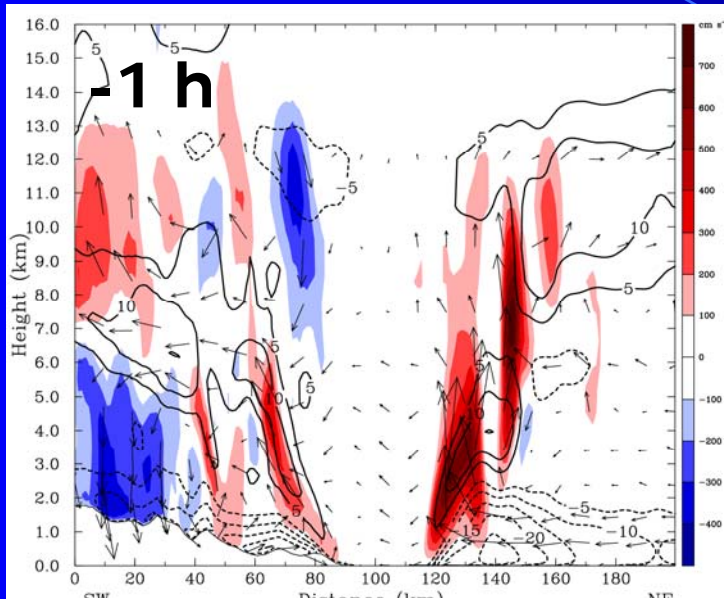
Color: Radial flow (U')

Landfall Stage @ 22-23 h

Horizontal Cross Sections of Vertical Velocity (colored) & Radial Wind (contoured)



SW-NE Vertical Cross Sections of Vertical Velocity (colored) & Radial Wind (contoured)



Conclusions (I)

■ Precipitation structure changes after landfall:

- Precipitation is widely spread over a larger area.
- Cloud water amount averaged within the inner core is nearly doubled and maximized at lower level.
- Rain water amount averaged within the inner core is increased by 50-70%, mainly produced by melting by graupel particles.
- Ice-phase hydrometeors remain similar vertical profiles after landfall.

■ The dominant latent heating (cooling) process within the inner core is **condensational heating (evaporative cooling)**; ice-phase processes are more important in outer rainbands.

Conclusions (II)

■ Latent-heating/cooling structure changes:

- Condensational heating avg. within inner core is almost doubled, and maximized at lower height
- Evaporative cooling avg. within inner core is increased by 50-70%
- Total latent heating within inner core is stronger (almost doubled for peak intensity) and located at a lower height (5 km to 3.5 km) after landfall

■ After Nari's landfall on Taiwan, the **axis of RMW is tilted outward**. **Tangential wind is reduced** by the enhanced surface friction and turbulence mixing over topography.

Conclusions (III)

- After Nari's landfall, the radial inflow at low level becomes stronger and thicker, and strong updrafts occur within the eye, due to enhanced convergence over terrain.
- The centrifugal force (U_E) is greater than the pressure gradient and Coriolis forces ($U_P + U_C$), leading to supergradient winds over Taiwan's topography. It results in the sloping radial outflow jet at middle level over the terrain.
- After landfall, Taiwan topography imposes significant asymmetries on the radial acceleration terms, resulting in strong asymmetries in radial winds.